

# Data-driven models for Monthly Streamflow Time Series Prediction

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## ABSTRACT

Data-driven techniques such as Auto-Regressive Moving Average (ARMA), K-Nearest-Neighbors (KNN), and Artificial Neural Networks (ANN), are widely applied to hydrologic time series prediction. This paper investigates different data-driven models to determine the optimal approach of predicting monthly streamflow time series. Four sets of data from different locations of People's Republic of China (Xiangjiaba, Cuntan, Manwan, and Danjiangkou) are applied for the investigation process. Correlation integral and False Nearest Neighbors (FNN) are first employed for Phase Space Reconstruction (PSR). Four models, ARMA, ANN, KNN, and Phase Space Reconstruction-based Artificial Neural Networks (ANN-PSR) are then compared by one-month-ahead forecast using Cuntan and Danjiangkou data. The KNN model performs the best among the four models, but only exhibits weak superiority to ARMA. Further analysis demonstrates that a low correlation between model inputs and outputs could be the main reason to restrict the power of ANN. A Moving Average Artificial Neural Networks (MA-ANN), using the moving average of streamflow series as inputs, is also proposed in this study. The results show that the MA-ANN has a significant improvement on the forecast accuracy compared with the original four models. This is mainly due to the improvement of correlation between inputs and outputs depending on the moving average operation. The optimal memory lengths of the moving average were three and six for Cuntan and Danjiangkou respectively when the optimal model inputs are recognized as the previous twelve months.

**Keywords:** Hydrologic time series, Auto-Regressive Moving Average, K-Nearest-Neighbors, Artificial Neural Networks, Phase Space Reconstruction, False Nearest Neighbors, dynamics of chaos.

## 1. Introduction

Many data-driven models, including linear, nonparametric or nonlinear approaches, are developed for hydrologic discharge time series prediction in the past decades (*Marques et al., 2006*). Generally, there are two basic assumptions underlay different model techniques. The first assumption suggests that a time series is originated from a stochastic process with an infinite number of degrees of freedom. Under this assumption, linear models such as AutoRegressive (AR), AutoRegressive Moving Average (ARMA), AutoRegressive Integrated Moving Average (ARIMA), and Seasonal ARIMA (SARIMA) had made a great success in river flow prediction (*Carlson et al., 1970; Salas et al., 1985; Haltiner and Salas, 1988; Yu and Tseng, 1996; Kothiyari and Singh, 1999; Huang et al., 2004; María et al., 2004*).

The second assumption is that a random-looking hydrologic time series is derived from a deterministic dynamic system such as chaos. In the past two decades, chaos-based streamflow prediction techniques have been increasingly obtaining interests of the hydrology community (*Jayawardena and Lai, 1994; Jayawardena and Gurung, 2000; Elshorbagy et al., 2002; Wang et al, 2006b*) although some doubts have been raised in terms of the existence of chaos in hydrologic data (*Ghilardi and Rosso, 1990; Koutsoyiannis and Pachakis, 1996; Pasternack, 1999; Schertzer et al., 2002; Wang et al., 2006a*). Generally, the prediction techniques for a dynamic system can be roughly divided into two approaches: local and global. Local approach uses only nearby states to make predictions whereas global approach involves all the states. K-Nearest-Neighbors (KNN) algorithm, Artificial Neural Networks (ANN) and Support Vectors Machine (SVM) are some typical forecast methods for dynamic systems (*Sivapragasam et al., 2001; Laio et al. ,2003; Wang et al., 2006b*). Phase-Space-Reconstruction (PSR) is a

1 precondition before performing any predictions of the dynamic system. Typical methods  
2 involved in PSR are correlation integral, singular-value decomposition of the sample  
3 covariance matrix, False Nearest Neighbors (FNN), and true vector fields (*Grassberger and*  
4 *Procaccia, 1983; Abarbanel et al., 1993*).

5 Comparative studies on the above prediction techniques have been further carried out  
6 by some researchers. Sivakumar et al. (2002) found that the performance of the KNN approach  
7 was consistently better than ANN in short-term river flow prediction. Laio et al. (2003) carried  
8 out a comparison of KNN and ANN for flood predictions and found that KNN performed  
9 slightly better at short forecast time while the situation was reversed for longer time. Similarly,  
10 Yu et al. (2004) proposed that KNN performed worse than ARIMA on the basis of daily  
11 streamflow prediction. The conclusions in literature are very inconsistent. It is difficult to  
12 justify which modeling technique is more suitable for a streamflow forecast.

13 The above two assumptions are in the extremes of a hydrologic streamflow series. Salas  
14 et al. (1985) suggested that a streamflow process should be treated as an integration of  
15 stochastic (or random) and deterministic components. Describing it as either a totally linear  
16 stochastic process or fully nonlinear deterministic chaos is not a practical approach (*Elshorbagy*  
17 *et al. 2002*). Therefore, the model based on either of two assumptions may not be the most  
18 suitable. An investigation on an optimal prediction model is worthy to further study with  
19 different real monthly streamflow data (Xiangjiaba, Cuntan, Manwan, and Danjiangkou).

20 The scope of this study is to compare four forecast models, ARMA, ANN, KNN, and  
21 ANN-PSR and develop an optimal model for monthly streamflow prediction. This paper is  
22 organized in the following manner. Section 2 presents the four sets of streamflow data used in  
23 this study. Section 3 first describes the principles of PSR and then identifies its parameters  
24 using the correlation integral approach and the FNN approach. The implementation of the  
25 forecast models, including data preparation and selection of parameters, is discussed in Section  
26 4. Forecast results are described in Section 5 and conclusions of the study are presented in  
27 Section 6.

## 28 **2. Streamflow Data**

29 Monthly streamflow series of three watersheds and one river, i.e. Xiangjiaba, Manwan,  
30 Danjiangkou, and Yangtze River, were analyzed in this study.

31 The largest watershed, Xiangjiaba, is at the upstream of Yangtze river with average  
32 yearly discharge of  $4538 \text{ m}^3/\text{s}$ . Monthly streamflow series were taken from the hydrological  
33 station near the Xiangjiaba Dam site located in Sichuan Province. The basin area contributed to  
34 the streamflow series is around  $45.88 \times 10^4 \text{ km}^2$ . The period of the data was from January 1940  
35 to December 1997.

36 The medium watershed, Manwan, is located in the Lancang River which originates  
37 from the Qinghai-Tibet Plateau. Monthly streamflow series were taken from the hydrological  
38 station near the Manwan Dam site located in Sichuan Province. The catchment area controlled  
39 by the station is  $11.45 \times 10^4 \text{ km}^2$ , and the average yearly discharge is  $1230 \text{ m}^3/\text{s}$  based on a  
40 statistic of 30-year data (January 1974 to December 2003).

41 The smallest watershed, Danjiangkou, lies at the upstream of Han river with average  
42 yearly discharge of  $1203 \text{ m}^3/\text{s}$ . Monthly streamflow data came from the hydrology station at the  
43 Danjiangkou Dam site which is located in Hubei Province. The catchment area at the dam site  
44 is around  $9.5 \times 10^4 \text{ km}^2$ . The data range was from January 1930 to December 1981.

1 The last streamflow series is Yangtze River, the largest river in China. The selected  
2 monthly streamflow data were from the hydrology station of Cuntan located in the middle  
3 stream of the river. The stream flow series spanned from January 1893 to December 2007.

4 Four monthly streamflow series are shown in Fig. 1. Monthly streamflow data in  
5 Xiangjiaba, Manwan, and Cuntan are characterized by a smooth process whereas monthly  
6 streamflow data in Danjiangkou exhibits complex oscillations. The linear fits (dotted lines in  
7 Fig. 1) verify the consistency of the streamflow series. All series exhibit good consistency  
8 because the linear fits are closed to horizontal. Since there was no large-scale hydraulic works  
9 such as dams built during the data collection period, the streamflow process is fairly pristine in  
10 each case.

### 11 3. Reconstruction of Dynamics

#### 12 3.1. Phase Space Reconstruction

13 To describe the temporal evolution of a dynamical system in a multi-dimensional phase  
14 space with a scale time series, it is essential to employ some techniques to unfold the multi-  
15 dimensional structure using the available data (*Wang et al., 2006a*). The most frequently used  
16 reconstruction method for a univariate or multivariate time series is the delay-time method  
17 (*Takens, 1981; Farmer and Sidorowich, 1987; Sauer et al., 1991; Jayawardena and Lai, 1994*).

18 A dynamic univariate time series  $\{x_1, x_2, \dots, x_N\}$ , may be reconstructed into a series of delay  
19 vectors of the type  $\mathbf{Y}_t = \{x_t, x_{t+\tau}, x_{t+2\tau}, \dots, x_{t+(m-1)\tau}\}$ ,  $t = 1, 2, \dots, N - (m-1)\tau$ , where  $\mathbf{Y}_t \in \mathbb{R}^m$ ,  $\tau$  is the  
20 delay time as a multiple of the sampling period and  $m$  is the embedding dimension. Under ideal  
21 conditions of time series of infinite length, all the reconstructions would be analogous and  
22 topologically equivalent to the real system. Owing to the shortness of real time series and the  
23 inevitable presence of dynamical noise, optimal reconstruction is involved in the choice of  
24  $m$  and  $\tau$  (*Laio et al., 2003*).

25 The time evolution of the dynamic system is given as a mapping  $\mathbf{Y}(t) \mapsto \mathbf{Y}(t+T)$   
26 (or  $\mathbf{Y}_t \mapsto \mathbf{Y}_{t+T}$ ). The function relationship between the current state  $\mathbf{Y}(t)$  at time  $t$  and the  
27 predicted state  $\mathbf{Y}^F(t+T)$  at time  $t+T$  can be written as follows:

$$28 \quad \mathbf{Y}^F(t+T) = f(\mathbf{Y}(t)) + e_t \quad (1)$$

29 where  $e_t$  is a typical noise term. In the form of time series, it can be expressed  
30 as  $[x_{t+T}^F, x_{t+T+\tau}^F, \dots, x_{t+T+(m-2)\tau}^F, x_{t+T+(m-1)\tau}^F] = f([x_t, x_{t+\tau}, \dots, x_{t+(m-2)\tau}, x_{t+(m-1)\tau}]) + e_t$ . Therefore,  
31 predicting future trajectory by current trajectory becomes viable once the function  $f(\bullet)$  is  
32 determined. In practice, the expression is often defined as:

$$33 \quad x_{t+T+(m-1)\tau}^F = f(\mathbf{Y}(t)) + e_t \quad (2)$$

34 where only the last component in  $\mathbf{Y}^F(t+T)$  is indicated since normally the prediction of this last  
35 component is of concern (*Laio et al., 2003*). For more details and examples of phase space  
36 reconstruction, please refer to *Laio et al. (2003)*.

37 Both global and local methods can be applied to estimate the function of  $f(\bullet)$  in Eq. (2).  
38 The global approach depends on the observations at all points whereas the local approach  
39 depends on observations that are in some finite neighborhood of the point of estimate. The  
40 KNN algorithm is a widely-used local method, which was originally developed by Farmer and  
41 Sidorowich (1987). The basic idea behind KNN is that only nearby states are used to make

1 predictions for the local approximation. This method is similar to the Nearest Neighbor Method  
2 (NNM) (Yakowitz and Karlsson, 1987; Toth et al., 2000; Solomatine et al., 2008) once an  
3 attractor is correctly unfolded or the phase space of the dynamical system is correctly  
4 reconstructed. Attractor reconstruction refers to those methods for inference of geometrical and  
5 topological information about a dynamical attractor from observations. An attractor is the point  
6 where the dynamics are discontinuous, through which the minimum embedding dimension can  
7 be determined. For the purpose of comparisons, a local approach of KNN algorithm (Farmer  
8 and Sidorowich, 1987; Jayawardena and Lai, 1994) and a global approach of ANN were  
9 discussed in this study.

### 10 3.2. Determination of Parameters ( $\tau, m$ )

11 Many methods, as mentioned in Section 1, are able to reconstruct the phase space by  
12 identifying the parameter pair ( $\tau, m$ ). The correlation integral and FNN were employed to the  
13 following discussion. The correlation integral identifies the two parameters for the perspective  
14 of verifying the existence of chaos.

#### 15 (1) Correlation exponent

16 The correlation exponent method is commonly used to investigate the existence of  
17 chaos in hydrology community (Jayawardena and Lai, 1994; Sivakumar et al., 1998;  
18 Sivakumar et al., 2001). Correlation integral is generally applied to verify the existence of  
19 chaotic dynamics using the saturation value of the correlation dimension. The diagnosis of the  
20 existence of chaos can begin if the phase space has been reconstructed. The PSR requires two  
21 parameters ( $\tau, m$ ), where  $m$  can be identified from the plot of the correlation dimension ( $d_2$ )  
22 versus  $m$  when it reaches a saturation value ( $D_2$ ) (Tsonis, 1992). In other words, the saturation  
23 value must be obtained prior the achievement of  $m$ . However,  $m$  is subject to  $\tau$  (often called  
24 decorrelation time) and should be first determined. Also, some researchers thought that  $\tau$  and  $m$   
25 should not be determined separately. The length of the embedding window,  $\tau_w = (m-1)\tau$   
26 (Broomhead and King, 1986; Mees et al., 1987; Martinerie et al., 1992) and optimizing the  
27 triplet ( $m, \tau, k$ ) using Genetic Algorithm (GA) (Liong et al., 2002) are typical examples of  
28 determining  $\tau$  and  $m$  together. In the present study, we tend to adopt a widely accepted  
29 method to obtain  $\tau$ .

30 Determination of  $\tau$  and  $m$  have been reported in numerous works (Grassberger and  
31 Procaccia, 1983; Fraser and Swinney, 1986; Casdagli et al., 1991; Tsonis, 1992; Abarbanel et  
32 al., 1993; Kugiumtzis, 1996; Hegger et al., 1999; and Kantz and Schreiber, 2004). Generally,  
33  $\tau$  can be defined when the AutoCorrelation Function (ACF) attains the value of zero or below a  
34 small value, or the Average Mutual Information (AMI) reaches the first minimum. The  
35 calculations of ACF and AMI are discussed in detail in the works of Fraser and Swinney (1986),  
36 Tsonis (1992), and Abarbanel et al. (1993). Fig. 2 displays the ACF and AMI for the four  
37 monthly streamflow series. The ACF for all series are first attained zeros at the same lag time  
38 of 3. Since the AMI gives the same estimates of  $\tau$  as the ACF,  $\tau$  was consistently chosen at the  
39 lag time of 3 for all cases.

40 It is also suggested that a relationship of  $\tau = T/m$  (where  $T$  represents the dominant  
41 periodicity of the original series as revealed by Fourier analysis) exists if ACF is periodic  
42 (Tsonis, 1992). Fig. 2 also demonstrates that the present cases suffice the condition. If all  
43 streamflow series are chaotic, the potential  $m$  are therefore around 4 (i.e.  $m = 12/3$ ) with all the  
44 predominant periodicity of 12 months and  $\tau$  of 3.

1 After the primary determination of  $\tau$ , the correlation dimension can be computed by the  
 2 correlation integral according the formula of Grassberger-Procaccia algorithm (*Grassberger  
 3 and Procaccia,1983*). This original formula was modified by Theiler (*1986*) for the estimation  
 4 of the correlation integral in a time series which poses serious problems of temporal  
 5 correlations. Thus, the modified correlation integral  $C(r)$  for a  $m$ -dimension phase space is  
 6 defined as:

$$7 \quad C(r) = \frac{2}{N_{\text{pairs}}} \sum_{i=1}^N \sum_{j=i+w+1}^{N-i} H(r - \|\mathbf{Y}_i - \mathbf{Y}_j\|) \quad (3)$$

8 where  $N_{\text{pairs}} = (N - w + 1)(N - w)$ ,  $w$  is the Theiler window excluding those points which are  
 9 temporally correlated,  $\mathbf{Y}_i$  and  $N$  are already discussed in Section 3.1,  $r$  is the radius of a ball  
 10 centered on  $\mathbf{Y}_i$ ,  $H$  is the Heaviside step function with  $H(u) = 1$  if  $u > 0$  and  $H(u) = 0$  if  $u \leq 0$ .  
 11 The correlation integral only counts the pairs  $(\mathbf{Y}_i, \mathbf{Y}_j)$  whose distance, in a Euclidean sense, is  
 12 smaller than  $r$ . In the limit of an infinite amount of data ( $N \rightarrow \infty$ ) and sufficiently small  $r$ , the  
 13 relation of  $C(r) \propto r^{D_2}$  between  $C(r)$  and  $r$  is expected when  $m$  exceeds the correlation  
 14 dimension of the chaos system. The correlation exponent  $\nu$  and the correlation dimension  $D_2$   
 15 can be defined as:

$$16 \quad \nu = \frac{\partial \ln C(r)}{\partial \ln r} \quad (4)$$

$$17 \quad D_2 = \lim_{\substack{r \rightarrow 0 \\ N \rightarrow \infty}} \nu \quad (5)$$

18 Since  $D_2$  is unknown before conducting the computation, the convergence of the correlation  
 19 dimension  $D_2$  in  $m$  must be examined.

20 The procedure of the computation is first to plot  $\ln C(r)$  versus  $\ln r$  with a given  $m$ .  
 21 Then, the potential scaling region is determined wherever the slope (i.e. the correlation  
 22 exponent  $\nu$ ) of the curve for the given  $m$  is approximately constant. The constant slope can be  
 23 estimated by a straight line fitting of the scaling region. In general, the best way to define the  
 24 scaling region is to produce another figure which demonstrates the slope of the  $\ln C(r)$  as a  
 25 function of  $\ln r$ . If a scaling region exists, a plateau should be shown in the figure. This plateau  
 26 provides an estimate for  $d_2$ , a correlation dimension of the possible attractor for the present  $m$ .  
 27 If  $d_2$  converges to a finite value  $D_2$  (i.e. saturation value) after repeating the above procedure for  
 28 successively higher  $m$ , a true attractor of dimension  $D_2$  is formed and the system may be  
 29 considered as chaos. Meanwhile,  $m$  can be identified as the value that corresponds to the first  
 30 occurrence of the saturation value  $D_2$  in the plot of  $d_2$  versus  $m$ .

31 For a finite dataset, there is a radius  $r$  below which there are no pairs of points (i.e.  
 32 depopulation zone). Conversely, when  $r$  approaches the diameter of the cloud of points, the  
 33 number of pairs will increase no further as  $r$  increases (i.e. saturation zone). The scaling region  
 34 would be found somewhere between the depopulation and the saturation zones. In view  
 35 of  $C(r) \propto r^{D_2}$ , the population of pairs of points for a finite data set on small scales is smaller  
 36 than the population of pairs on large scales. This leads to poor statistics at small  $r$  and the  
 37 function  $C(r)$  may be distorted. Nevertheless, the scaling region over large  $r$ 's should remain  
 38 unchanged if there are sufficient points available. On the contrary, the scaling region may be  
 39 completely masked if there are inadequate points. Also, the scaling region can be "lost"  
 40 between the depopulation and the saturation zones by increasing the  $m$  while the number of

1 points remain the same positions. Thus, an accurate estimation of  $d_2$  requires a minimum  
2 number of points.

3 Many literatures focus on the quantity of data required to determine the reliable value of  
4  $d_2$  (Abarbanel *et al.*, 1993 and Wang *et al.*, 2006a). Some researches claim that the size should  
5 be  $10^A$  (Procaccia, 1988) or  $10^{2+0.4m}$  (Tsonis, 1992), where  $A$  is the greatest integer smaller than  
6  $d_2$  and  $m$  ( $m < 20$ ) is the embedding dimension used for estimating  $d_2$  with an error less than  
7 5%. Whereas other researches found that smaller data size is needed. For instance, the  
8 minimum data points for reliable  $d_2$  is  $10^{d_2/2}$  (Ruelle, 1990; Essex and Nerenberg, 1991), or  
9  $\sqrt{27.5}^{d_2}$  (Hong and Hong, 1994) and empirical results of dimension calculations are not  
10 substantially altered by going from 3000 or 6000 points to subsets of 500 points (Abraham *et*  
11 *al.*, 1986).

12 The  $\ln C(r)$  versus  $\ln r$  graphs and the correlation exponent  $\nu$  versus  $\ln r$  for the four  
13 streamflow series are shown in Figs. 4 and 5 respectively. Fig. 5 demonstrates that the scaling  
14 region cannot be identified for any  $m$  ( $m < 20$ ) in Danjiangkou catchment whereas the scaling  
15 region can be determined in other three cases. Obviously, Danjiangkou series cannot be totally  
16 defined as non-chaotic or random process because its number of data is as small as 624. A fact  
17 should be noted for the other three streamflow series is that the identified scaling regions  
18 become narrow or even “lost” as the saturation zone occurs over larger scales with the increase  
19 of  $m$ . If the number of data in the streamflow series decreases, the speed for the scaling region  
20 to narrow or “lost” with the increase of  $m$  will become faster. A representative detail of these  
21 curves in Fig. 5 is presented in Fig. 6 by taking curves at  $m = 1, 4, 8$  and 20. The scaling region  
22 becomes ambiguous at  $m = 8$  for Manwan and Xiangjiaba whereas a narrow scaling region can  
23 still be defined even at  $m = 20$  for Cuntan which has 1380 points from data of 114 years.  
24 However, the length of data size in Danjiangkou, Cuntan and Manwan is adequate in defining  
25 the scaling region when  $m \leq 4$ . Based on the results of Fig. 5, the relationship between  
26 correlation dimension  $d_2$  and embedding dimension  $m$  is depicted in Fig. 7. The saturation  
27 values  $D_2$  for the three streamflow series are at interval of (1.5, 2). Generally, a sufficient  
28 condition for the smallest  $m$  is that  $m$  is an integer larger than  $2D_2$ . The associated  $m$  is  
29 therefore set the value of 4 for the three series. A test for the robustness of  $m$  with variable  
30  $\tau$  from 1 to 5 was performed, which implied that  $m$  is insensitive to  $\tau$ .

31 With the potential values on  $D_2$  or  $m$ , some criteria such as  $10^A$  can be satisfied whereas  
32 other criteria such as  $10^{2+0.4m}$  cannot be satisfied. The latter criteria means that few hydrologic  
33 records can be assessed for  $m > 5$  attractors since as many as 10,000 points require 27 years of  
34 daily records or around 900 years of monthly records. Thus, the three monthly series of  
35 Danjiangkou, Cuntan and Manwan may be treated as chaotic with suggested variable  $m = 4$ .  
36 Furthermore, the phase portraits of four streamflow series are portrayed in Fig. 8 where  $(\tau, m)$  is  
37 (3,3). Obviously, the state spaces in the 3-dimensional maps are clearly unfolded for Xiangjiaba,  
38 Cuntan, and Manwan whereas no clear trajectory is revealed for Danjiangkou.

### 39 (2) False Nearest Neighbors(FNN)

40 The correlation integral method appears to be data intensive and certainly subjective.  
41 For simplicity, the FNN method is commonly employed for the PSR of a hydrologic  
42 streamflow series (Wang *et al.*, 2006b). The FNN algorithm was originally developed for  
43 determining the number of time-delay coordinates needed to recreate autonomous dynamics  
44 directly from properties of the data itself (Kennel *et al.*, 1992; Abarbanel *et al.*, 1993). The  
45 following discussion outlines the basic concepts of the FNN algorithm. Suppose the point

1  $\mathbf{Y}_i = \{x_i, x_{i+\tau}, x_{i+2\tau}, \dots, x_{i+(m-1)\tau}\}$  has a neighbor  $\mathbf{Y}_j = \{x_j, x_{j+\tau}, x_{j+2\tau}, \dots, x_{j+(m-1)\tau}\}$ , the criterion that  
 2  $\mathbf{Y}_j$  is viewed as a false neighbor of  $\mathbf{Y}_i$  is:

$$3 \quad \frac{|x_{i+m\tau} - x_{j+m\tau}|}{\|\mathbf{Y}_i - \mathbf{Y}_j\|} > R_{\text{tol}} \quad (6)$$

4 where  $\|\cdot\|$  stands for the distance in a Euclidean sense,  $R_{\text{tol}}$  is some threshold with the common  
 5 range of 10 to 30 (Wang *et al.*, 2006b). For all points  $i$  in the vector state space, Eq. (6) is  
 6 performed and then the percentage of points which have FNNs is calculated. The algorithm is  
 7 repeated for increasing  $m$  until the percentage of FNNs drops to zero, or some acceptable small  
 8 number such as 1%, where  $m$  is the target  $m$  (Abarbanel *et al.*, 1993).

9 The sensitivity analysis of the percentage of FNNs (FNNP) on  $R_{\text{tol}}$  is respectively  
 10 demonstrated in Fig. 9, where  $\tau$  is set as value of 3 and  $R_{\text{tol}}$  is from 10 to 30 with a step size of  
 11 5. Fig. 9 shows that FNNP is stable when  $R_{\text{tol}} \geq 15$ . Fig. 10 demonstrates FNNPs for the four  
 12 streamflow series with  $R_{\text{tol}} = 15$ . The identified  $m$  is 4 for Manwan and 5 for other three cases.  
 13 Two random series with the same data size as Cuntan and Danjiakou were respectively  
 14 generated by the random function. Their FNNPs with  $\tau$  of 3 and  $R_{\text{tol}}$  of 15 are demonstrated in  
 15 Fig. 11 where a similar  $m$  was found. The FNN technique may not be able to distinguish  
 16 random process from deterministic system. Therefore, the phase space reconstructed by the  
 17 FNN may not reveal the true trajectory hidden in the dynamic system. In other words, the FNN  
 18 technique is not concerned with a dynamic system being deterministic or not. Thus, the  
 19 correlation integral method can be more reliable for unfolding a dynamical system. The  
 20 preliminary parameter pair  $(\tau, m)$  in PSR is (3, 4) for Xiangjiaba, Cuntan, and Manwan. For the  
 21 purpose of the symmetry in models comparison, the parameter pair  $(\tau, m)$  for Danjiangkou is (3,  
 22 5) from Fig. 10. Certainly, these parameters pair need to be checked further for their robustness  
 23 when they are used in KNN model.

## 24 4. Implementation of prediction models

### 25 4.1 Data preparation

26 Streamflow series data were divided into three parts: model training, cross-validation  
 27 and validation. The last ten years' streamflow data, called validation set, were set aside for  
 28 validation. This validation set would not be used until all model development and training was  
 29 finished completely. Of the remaining data, the first two-thirds called training set was for  
 30 model training, and the other one-third called cross-validation set was for the purpose of  
 31 confirming and validating the initial analysis. The ANN model has difficulties to extrapolate  
 32 beyond the range of the data used for training. As a consequence, poor predictions can be  
 33 expected when the validation data contains values outside the range of those used for training.  
 34 It is also imperative that the training and validation sets are representative of the same  
 35 population. Statistical properties (mean, deviation, range) from them are compared in order to  
 36 measure the representative. Table 1 displays the statistical properties of these data sets.  
 37 Generally, the division of data is satisfied for ANN since the statistical parameters of the  
 38 training sets for all the streamflow series are closed to the cross-validation sets and the testing  
 39 sets.

## 4.2 Evaluation of model performances

The Pearson's correlation coefficient ( $r$ ) or the coefficient of determination ( $R^2=r^2$ ), have been identified as inappropriate measures in hydrologic model evaluation by Legates and McCabe (1999). The coefficient of efficiency (CE) (Nash and Sutcliffe, 1970) is a good alternative to  $r$  or  $R^2$  as a "goodness-of-fit" or relative error measure in that it is sensitive to differences in the observed and forecasted means and variances. Legates and McCabe (1999) also suggested that a complete assessment of model performance should include at least one absolute error measure (e.g., Root Mean Square Error) as necessary supplement to a relative error measure. Besides, the Persistence Index (PI) (Kitanidis And Bras, 1980) is able to check the prediction lag effect.

Thus, three measures of evaluating model performance are used in the present paper comprising Root Mean Square Error (RMSE), the Nash-Sutcliffe Coefficient of Efficiency (CE) (Nash and Sutcliffe, 1970), and the PI (Kitanidis And Bras, 1980). They are respectively

formulated as:  $RMSE = \sqrt{\frac{1}{N} \sum_{i=1}^N (T_i - \hat{T}_i)^2}$ ,  $CE = 1 - \frac{\sum_{i=1}^N (T_i - \hat{T}_i)^2}{\sum_{i=1}^N (T_i - \bar{T})^2}$  and

$PI = 1 - \frac{\sum_{i=1}^N (T_i - \hat{T}_i)^2}{\sum_{i=1}^N (T_i - T_{i-L})^2}$ . In these equations,  $N$ =number of observations,  $\hat{T}_i$ =predicted

streamflow,  $T_i$ =observed streamflow,  $\bar{T}$ =average observed streamflow, and  $T_{i-L}$  is the streamflow estimated from a so-call persistence model (or called naïve model) that basically takes the last streamflow observation (at time  $i$  minus the lead time  $L$ ) as a prediction. CE and PI values of 1 indicate perfect fits.

## 4.3 Configurations of models and parameters optimization

Cuntan streamflow series was used as a representative case due to the similarity among Xiangjiaba, Cuntan, and Manwan. Two cases, Cuntan and Danjiangkou, are therefore analyzed in the rest of this article.

Autoregressive-moving-average (ARMA) models are mathematical models of the persistence, or autocorrelation, in a time series. ARMA models can contribute to understanding the physical hydrological system by disclosing some information about the physical process that builds persistence into the series. Models of ARMA and ANN were used for the purpose of comparison with models based on PSR. The order of the ARMA model is included in parentheses as ARMA( $p,q$ ), where  $p$  is the autoregressive order and  $q$  the moving-average order. Parameters ( $p, q$ ) in ARMA were estimated by trial and error with each of them varied over the range of 0 to 12, excluding the unfeasible case where both of them are simultaneously equal to 0, in view of the predominant periodicity of 12 months (Figs. 3 and 4). The best pair of ( $p, q$ ) is associated with the minimal value of Akaike Information Criterion (AIC), which is a commonly adopted statistical measures of goodness-of-fit of an ARMA( $p,q$ ) model. Goodness of fit might be expected to be measured by some function of the variance of the model residuals, i.e., the fit improves as the residuals become smaller.

ANNs mimic the functioning of the human brain and nervous systems in a simplified computational form. They are constituted by highly interconnected simple elements (termed artificial neurons or sometimes termed nodes) which receive information, elaborate them through mathematical functions and pass them to other artificial neurons (Alvisi et al., 2006). They acquire knowledge through a learning process that involves finding an optimal set of weights for the connections and threshold values for the neurons. ANNs can be categorized as single layer, bilayer and multilayer according to the number of layers, and as feed-forward,

1 recurrent, and self-organizing according to the direction of information flow and processing  
2 ([ASCE, 2000a](#)). With respect to ANN model, the type of ANN and its architecture must be  
3 determined. The ANN type used in this study was the static multilayer feed-forward network.  
4 The static ANN is able to capture the dynamics of a system in the network model by using  
5 delay time inputs. The architecture design of the ANN comprises the number of hidden layers  
6 and the number of neurons in input layer, hidden layers and output layer. ANNs with one  
7 hidden layer are commonly used in hydrologic modeling ([Dawson and Wilby, 2001](#); [de Vos  
8 and Rientjes, 2005](#)) since these networks are considered to provide enough complexity to  
9 accurately simulate the dynamic and nonlinear-properties of the hydrologic process. Therefore,  
10 a three-layer static ANN was finally employed.

11 One output neuron was selected whilst the input neurons were determined by following  
12 the popular approach of examining the dependence between the input and output time series.  
13 The model inputs are originally considered to take the previous three sequential data because  
14 the ACF first attains zero and AMI achieves the first local minima at the lag time of 3  
15 simultaneously (Fig. 3). A trial and error method was then performed to check the three inputs  
16 by systematically increasing the number of inputs from the latest month to the past twelve  
17 months. The test results showed that the model with the 12-month data as inputs was optimal  
18 for Cuntan and Danjiangkou.

19 The ensuing task is to optimize the size of the hidden layer with the chosen three inputs  
20 and one output. The optimal size of the hidden layer was found by systematically increasing the  
21 number of hidden neurons from 1 to 20 until the network performance on the cross-validation  
22 set was no longer improved significantly. The optimal number of hidden neurons may be  
23 affected by the use of a training algorithm. The most popular supervised training algorithms are  
24 gradient descent techniques (e.g. BackPropagation (BP) algorithm) and Newton or quasi-  
25 Newton optimization techniques (e.g., Levenberg-Marquart (LM) algorithm, and Broyden-  
26 Fletcher-Goldfarb-Shanno (BFGS) algorithm). They are called local optimization techniques.  
27 Alternatives to these methods are global optimization techniques such as the Particle Swarm  
28 Optimization (PSO), the Genetic algorithm (GA), and the Shuffled Complex Evolution (SCE-  
29 UA). Generally, these local optimization methods have fast convergence but are susceptible to  
30 local minima and unstable. The global optimization methods, on the other hand, are able to  
31 overcome local minima and obtain more stable solutions but are slow convergence and tend to  
32 find out a relative optimal solution. As an attempt to combine their merits of the two  
33 optimization techniques, an integrated training algorithm combining LM with GA is adopted  
34 ([Chau et al., 2005](#)). The basic idea behind this algorithm is that GA is first utilized in search of  
35 a set of weights and biases for ANN. These chosen weights and biases are then used as initial  
36 values for a further optimization via the LM algorithm. Consider Danjiangkou as a pilot case,  
37 Table 2 demonstrates the one-step-ahead forecast performance statistics with 30 time runs in  
38 terms of RMSE, CE and PI along with training time of three training algorithms of LM, PSO  
39 and integrated algorithm of LM and GA. As expected, the PSO has better stability of  
40 performance at the expense of training time whereas the LM has faster convergence but  
41 exhibits a weak stability of performance in terms of the performance values in the Range  
42 column (last column of Table 2). The hybrid algorithm indeed combines their merits with both  
43 fast convergence and good stability of performance. Thus, the LM-GA algorithm was used for  
44 training in the following study. The PSO is also acceptable because the time-consuming in  
45 training is not too long. It was found that the optimal size of the hidden layer was 5 for all three  
46 algorithms. Finally, a 12-5-1 configuration of the ANN model with the LM-GA algorithm was  
47 designed for Danjiangkou. Similarly, a 12-13-1 configuration of ANN was applied to Cuntan.

1 The hyperbolic tangent functions were used as a transfer function in the hidden layer and output  
2 layer. All input and output data were linearly scaled to the range of [-1, 1] because the output of  
3 this transfer function is bounded to the range of [-1, 1].

4 The KNN method is based on local approximation which makes use of only nearby  
5 observations of the point of estimate. It has its origins as a non-parametric statistical pattern  
6 recognition procedure, aiming at distinguishing between different patterns according to chosen  
7 criteria ([Mack and Rosenblatt, 1979](#); [Aha et al., 1991](#)). Amongst various non-parametric kernel  
8 methods, the K-NN method seems more intuitive, but nevertheless possesses powerful  
9 statistical properties ([Karlsson and Yakowitz, 1987a, b](#)). Yakowitz ([1987](#)) and Karlsson and  
10 Yakowitz ([1987a, b](#)) did considerable work in extending the KNN method to time-series and  
11 forecasting problems, obtaining satisfactory results and constructing a robust theoretical base  
12 for the KNN method. Please refer to [Karlsson and Yakowitz \(1987a\)](#) for details.

13 Two PSR-based models, KNN and ANN (hereafter referred to as ANN-PSR), were  
14 applied to the following study. ANN-PSR is the phase space reconstruction-based artificial  
15 neural networks integrating the techniques of phase space reconstruction and ANN. Both the  
16 above two models have the same inputs which were the  $m$  coordinates of PSR. The number of  
17  $k$  was used as  $k = m + 1$  ([Sugihara and Mary, 1990](#)), but  $k > m + 1$  ([Farmer and Sidorowich,  
18 1987](#)) was applied to ensure the stability of the solution. In order to verify the robustness of  
19  $m$  and  $\tau$ , a trial-and-error procedure (i.e. testing the optimal combinations of  $\tau$  and  $m$ ) still  
20 have to be performed although the preliminary parameter values of them have been  
21 recommended previously. With all thirty-five combinations of  $\tau$  (1 to 5) and  $m$  (2 to 8), the  
22 identified optimal parameter pair ( $\tau, m$ ) in terms of RMSE were (5, 7) for Cuntan and (5, 6) for  
23 Danjiangkou. Correspondingly, the fixed  $k = 8$  and  $k = 7$  were respectively adopted in Cuntan  
24 and Danjiangkou. The ANN-PSR also employed the LM-GA algorithm for training. The  
25 optimal size of the hidden layer was determined via the same method mentioned above. Finally,  
26 the architectures of the ANN-PSR are 5-9-1 for both Cuntan and Danjiangkou. A summary on  
27 relevant model parameters for one-step-ahead prediction were displayed in Table 3.

## 28 5. Results and discussion

### 29 5.1. Main results

30 Table 3 shows the one-month-ahead forecast performance of various models in terms of  
31 the RMSE, CE, and PI for Cuntan and Danjiangkou. The KNN outperformed other models in  
32 terms of the three performance indexes. However, the advantage of KNN over ARMA is very  
33 weak. The comparison between ANN and ARMA demonstrates that the ANN cannot perfectly  
34 capture the autocorrelation relationship in each streamflow series. ANN-PSR model displays  
35 worse performance than the KNN in each case, which suggests that the local approach may be  
36 more suitable for one-month-ahead forecast. For the perspective of different cases, the  
37 performance indexes from each model are very low for Danjiangkou. Actually, Danjiangkou  
38 series was characterized by very complex oscillations as depicted in Fig. 1. Its main signal was  
39 heavily contaminated by noises as an evidence of appearance of many high frequency signals  
40 (Fig. 2).

41 Figs. 12 and 13 present the detail of the observed and predicted validation data series,  
42 their errors (predicted-observed), and relative errors (hereafter referred to REs) of forecasts  
43 from four prediction models for Cuntan. These models exhibit good overall match between the  
44 observed and predicted data series. The plots of errors (middle graphs of Figs. 12 and 13)  
45 illustrates that each model underestimates or overestimates quite a number of peaks of

1 streamflows although the grand mean of REs is around 20%. The plots of REs (bottom graphs  
2 of Figs. 12 and 13) suggest that the low streamflows are mostly well forecasted. This is due to  
3 the frequent occurrences of the low streamflow patterns which allow the trained model to have  
4 better generalization of these regions. Figs. 12 and 13 also demonstrate that the timing of the  
5 peaks for these models is quite good. The results are consistent with the high PI in Table 3.

6 Figs. 14 and 15 show the detail of the observed and predicted validation data series,  
7 their errors, and REs from four prediction models for Danjiangkou. The forecasts of the peaks  
8 of streamflows are mostly underestimated although these models can capture the whole trend of  
9 validation data series. The grand mean of REs are about 50% but the maximum RE is up to  
10 300% as shown in the plots of REs. There is a lag effect in the timing of the peaks for these  
11 models by visual inspection, which is reflected by the low PI in Table 3.

12 Histograms of forecast errors of one-month-ahead prediction for Cuntan are presented  
13 in Fig. 16, and corresponding Auto-Correlation Functions (ACFs) of forecast errors are  
14 depicted in Fig. 17. The histograms are quasi-normal distributions and errors satisfy random  
15 processes in terms of ACFs. The results suggest that employed four models were indeed trained  
16 fully and forecasts from them are reliable. Similar results also can be found for Dangjiangkou.

## 17 5.2. Discussion

18 The parameter of  $k$  in KNN model poses a great impact on the performance of KNN.  
19 As mentioned previously, the choice of  $k$  should still ensure the stability of the solution  
20 although a preliminary value of  $k$  is based on  $k = m + 1$ . The verification of stability of  $k$  in  
21 terms of RMSE is presented in Fig. 18. The new value of  $k$  was 9 for both Cuntan and  
22 Danjiangkou (recall that the original  $k$  was 8 and 7 respectively). The performance of KNN  
23 with new  $k$  is presented in Table 4. Comparison of performance between the new KNN and the  
24 original KNN (Table 3) suggests that taking  $k = m + 1$  tends to be reasonable because the  
25 new  $k$  is nearby  $m + 1$  and there is only a small improvement of performance for the KNN with  
26 the new  $k$ .

27 The issue of lagged predictions in the ANN model has been mentioned by some  
28 researchers (*Dawson and Wilby, 1999; Jian and Srinivasulu, 2004; de Vos and Rientjes, 2005;*  
29 *Muttill and Chau, 2006*). The work of de Vos and Rientjes (2005) in the context of rainfall-  
30 runoff modeling investigated sources of the prediction lag effect and suggested two types of  
31 methods to resolve the issue. One of identified sources is due to the use of previous observed  
32 streamflow data as ANN inputs. Thus, the ANN tends to give the most weight to the latest  
33 streamflow input if the model inputs are highly auto-correlative. As a consequence, ANN  
34 models underrate the information contained in other input signals such as precipitation. An  
35 effective solution is to decrease the weight of streamflow components as model inputs by using  
36 a moving average time series of original streamflows as streamflow input components of the  
37 rainfall-runoff ANN.

38 Fig. 19 investigates the existence of the forecast lag effect by Cross-Correlation  
39 Functions (CCFs) between the predicted and observed validation data series. As expected by  
40 visual inspection previously, there is no lag effect for Cuntan whereas there is obvious lag  
41 effect in all models except for the KNN for Danjiangkou. The moving average operation was  
42 used to generate new monthly streamflow time series to construct the model inputs of ANN.  
43 Based on trial and error method, the optimal memory lengths of three months and six months  
44 were respectively used in the moving average of the streamflow series of Cuntan and  
45 Danjiangkou. With the same architecture as the original ANN, the new moving average ANN  
46 (hereafter referred to MA-ANN) achieved a significant improvement in performance of

1 predictions. Table 4 presents the forecast results using the MA-ANN for Cuntan and  
2 Danjiangkou. The high PI (over 0.9) indicates the elimination of timing error. Fig. 20 depicts  
3 the detail of forecasts of MA-ANN for Cuntan and Danjiangkou. The observed data was fitted  
4 perfectly in particular for Cuntan. The errors and REs also significantly decreased. Fig. 21  
5 compares the CCFs between inputs and outputs in the MA-ANN model and the ANN model.  
6 The results show that absolute CCFs in the first twelve lags from the former are mostly larger  
7 than those from the latter, in particular for Danjiangkou. Therefore, the improved performance  
8 from MA-ANN is due to the increase of the correlation between inputs and outputs by moving  
9 average operation on original streamflow time series. This conclusion was different from the  
10 effect of the moving average method employed in the work of de Vos and Rientjes (2005). The  
11 moving average operation is however beneficial to the performance of ANN for one-step-ahead  
12 streamflow prediction indeed.

## 13 6. Conclusions

14 The purpose of this study attempts to determine a relative optimal forecast model for  
15 monthly streamflow data. Two methods of the PSR, the correlation dimension and the FNN,  
16 were employed. The results show that three cases, Xiangjiaba, Cuntan, and Manwan, exhibit  
17 chaotic characteristics on the basis of the correlation integral technique. The parameter pair  
18  $(\tau, m)$  identified by the correlation integral and FNN is only preliminary for the KNN and  
19 ANN-PSR models, and they need to be further chosen by trial and error method. The final  
20  $(\tau, m)$  is (5,7) and (5,6) for Cuntan and Danjiangkou respectively. Comparison of the  
21 performance for different models indicates that the KNN exhibited no obvious superiority to  
22 traditional ARMA. With respect to the PSR-based prediction model, the KNN outperformed  
23 the ANN-PSR. This implies that the local approximation technique is better than the global  
24 approximation technique for one-month-ahead discharge prediction (at least in the current  
25 cases). Additionally, the common three-layer-feed-forward ANN did not expose its powerful  
26 ability of mapping any complex function because its prediction performance was worse than  
27 ARMA. The CCFs between the observed and predicted data series show that the forecast lag  
28 effect exists for Danjiangkou. The operation of moving average on original streamflow series  
29 can significantly improve the performance of ANN and eliminate the timing error by increasing  
30 the correlation between inputs and outputs of ANN. For instance, the PI for Danjiangkou is  
31 from 0.39 in original ANN to 0.93 in MA-ANN and its corresponding CE is from 0.47 to 0.93.  
32 Based on the findings of this paper, the MA-ANN can be proposed as the optimal model for  
33 one-month-ahead forecast.

34

35

## 36 Abbreviation

37 ACF	AutoCorrelation Function
38 AIC	Akaike Information Criterion
39 AMI	Average Mutual Information
40 ANN	Artificial Neural Networks
41 AR	Auto-Regressive
42 ARMA	Auto-Regressive Moving Average
43 ARIMA	Auto-Regressive Integrated Moving Average
44 BFGS	Broyden-Fletcher-Goldfarb-Shanno
45 BP	BackPropagation
46 CCF	Cross-Correlation Function

1	CE	Coefficient of Efficiency
2	CT	Cuntan
3	DJK	Danjiangkou
4	FNN	False Nearest Neighbors
5	FNNP	Percentage of FNN
6	GA	Genetic algorithm
7	KNN	K-Nearest-Neighbors
8	LM	Levenberg-Marquart
9	LM-GA	Levenberg-Marquart and Genetic algorithm
10	MA-ANN	Moving Average Artificial Neural Networks
11	MW	Manwan
12	NNM	Nearest Neighbor Method
13	PI	Persistence Index
14	RE	Relative Error
15	PSO	Particle Swarm Optimization
16	PSR	Phase Space Reconstruction
17	ANN-PSR	Phase Space Reconstruction-based Artificial Neural Networks
18	RMSE	Root Mean Square Error
19	SARIMA	Seasonal ARIMA
20	SEC-UA	Shuffled Complex Evolution
21	SVM	Support Vectors Machine
22	XJB	Xiangjiaba
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## 1 **References**

- 2 Abraham, N.B., Albano, A.M., Das, B., de Guzman, G., Yong, S., Gioggia, R.S., Puccioni, G.P., and Tredicce, J.R.  
3 (1986), Calculating the dimension of attractors from small data. *Phys. Lett. A* 114, 217–221.
- 4 Abarbanel, H. D. I., Brown, R., Sidorowich, J. J., and Tsimring, L.S. (1993), The analysis of observed chaotic data  
5 in physical systems. *Reviews of Modern Physics*, 65(4), 1331-1392.
- 6 Aha, D., Kibler, D., and Albert, M. (1991), Instance-based learning algorithms. *Machine Learning*, 6, 37-66.
- 7 Alvisi, S., Mascellani, G., Franchini, M., Bárdossy, A. (2006), Water level forecasting through fuzzy logic and  
8 artificial neural network approaches. *Hydrology and Earth System Sciences*, 10, 1-17.
- 9 ASCE. (2000a), Artificial neural networks in hydrology 1: Preliminary concepts. *Journal of Hydrologic*  
10 *Engineering*, 5(2), 115-123.
- 11 Broomhead, D.S., and King, G.P., (1986), Extracting qualitative dynamics from experimental data, *Physica D*, 20,  
12 217.
- 13 Carlson, R.F., MacCormick, A.J.A. and Watts D.G., (1970), Application of linear models to four annual  
14 streamflow series, *Water Resour. Res.*, 6(4), 1070-1078.
- 15 Casdagli, M., Eubank, S., Farmer, J.D., and Gibson, J. (1991), State space reconstruction in the presence of noise,  
16 *Physica D*, 51, 52-98.
- 17 Chau, K.W., Wu, C.L., Li, Y.S., (2005). Comparison of several flood forecasting models in Yangtze River. *J.*  
18 *Hydrol. Eng.* 10 (6), 485–491.
- 19 Dawson, C. W., and Wilby, R. L., (2001), Hydrological Modeling Using Artificial Neural Networks. *Progress in*  
20 *Physical Geography*, 25(1), 80-108.
- 21 De Vos, N.J. and Rientjes, T.H.M., (2005), Constraints of artificial neural networks for rainfall -runoff modeling:  
22 trade-offs in hydrological state representation and model evaluation. *Hydrology and Earth System Sciences*, 9,  
23 111-126.
- 24 Elshorbagy, A., Simonovic, S.P., and Panu, U.S., (2002), Estimation of missing stream flow data using principles  
25 of chaos theory, *Journal of Hydrology*, 255, 123–133.
- 26 Elsner, J. B. and Tsonis, A. A., (1997), *Singular Spectrum Analysis: A New Tool in Time Series Analysis*. Plenum  
27 Press, New York.
- 28 Essex, C., and Nerenberg, M.A.H., (1991), *Proc. R. Soc. London, Series A* 435, 287.
- 29 Farmer, J. D., and Sidorowich, J. J., (1987), Predicting chaotic time series, *Phys. Rev. Lett.*, 59(4), 845–848.
- 30 Fraser, A.M. and Swinney, H.L., (1986), Independent coordinates for strange attractors from mutual information,  
31 *Physical Review A*, 33(2), 1134-1140.
- 32 Ghilardi, P., and Rosso, R., (1990), Comment on chaos in rainfall. *Water Resour. Res.* 26 (8), 1837–1839.
- 33 Grassberger, P., and Procaccia, I., (1983), Measuring the strangeness of strange attractors, *Physica 9D*, 189-208.
- 34 Haltiner, J.P., and Salas, J.D., (1988), Short-term forecasting of snowmelt discharge using ARMAX models,  
35 *Water Resources Bulletin*, 24(5), 1083-1089.

- 1 Hegger, R., Kantz, H. and Schreiber, T., (1999), Practical implementation of nonlinear time series methods: The  
2 TISEAN package. *CHAOS*, 9(2), 413-435.
- 3 Hong, S.Z., and Hong,S.M., (1994), An amendment to the fundamental limits on dimension calculations. *Fractals*  
4 2 (1), 123–125.
- 5 Huang, W.R., Xu, B., and Hilton, A. (2004), Forecasting Flows in Apalachicola River Using Neural Networks,  
6 *Hydrological Processes*, 18, 2545-2564.
- 7 Jain, A., and Srinivasulu, S., (2004), Development of effective and efficient rainfall-runoff models using  
8 integration of deterministic, real-coded genetic algorithms and artificial neural network techniques, *Water Resour.*  
9 *Res.*, 40, W04302.
- 10 Jayawardena, A.W., and Gurung, A.B., (2000), Noise reduction and prediction of hydro-meteorological time series:  
11 dynamical systems approach vs. stochastic approach. *Journal of Hydrology*, 228, 242–264.
- 12 Jayawardena, A.W., and Lai, F., (1994), Analysis and prediction of chaos in rainfall and stream flow time series,  
13 *Journal of hydrology*, 153, 23-52.
- 14 Kantz, H., and Schreiber, T., (2004), *Nonlinear time series analysis* (2<sup>nd</sup> edition). Springer.
- 15 Karlsson, M., and Yakowitz, S. (1987a), Nearest-neighbor methods for nonparametric rainfall–runoff forecasting.  
16 *Water Resources Research*, 23 (7), 1300-1308.
- 17 Karlsson, M., and Yakowitz, S., (1987b), Rainfall–runoff forecasting methods, old and new. *Stochastic Hydrology*  
18 *and Hydraulics*, 1, 303-318.
- 19 Kennel, M. B., Brown, R., Abarbanel, H. D. I., (1992), Determining embedding dimension for phase space  
20 reconstruction using geometrical construction. *Phy. Rev. A.*, 45(6), 3403-3411.
- 21 Kitanidis, P. K. and Bras, R. L., (1980), Real-time forecasting with a conceptual hydrologic model, 2, applications  
22 and results, *Wat. Resour. Res.*, 16 (6), 1034–1044.
- 23 Kothyari, U.C., and Singh, V.P., (1999), A multiple-input single-output model for flow forecasting, *Journal of*  
24 *Hydrology*, 220, 12–26.
- 25 Koutsoyiannis, D., and Pachakis, D., (1996). Deterministic chaos versus stochasticity in analysis and modeling of  
26 point rainfall series. *J. Geophys. Res.* 101 (D21), 26441–26451.
- 27 Kugiumtzis, D., (1996), State space reconstruction parameters in the analysis of chaotic time series - the role of the  
28 time window length, *Physica D* 95, 13-28.
- 29 Laio, F., Porporato, A. Revelli, R. and Ridolfi, L., (2003), A comparison of nonlinear flood forecasting methods,  
30 *Water Resour. Res.*, 39(5), 1129, doi:10.1029/2002WR001551.
- 31 Legates, D. R., and McCabe, Jr, G. J. (1999), Evaluating the use of goodness-of-fit measures in hydrologic and  
32 hydroclimatic model validation, *Water Resources. Research*, 35(1), 233- 241.
- 33 Liong, S.Y., and Sivapragasam, C., (2002), Flood stage forecasting with support vector machines, *Journal of*  
34 *American Water Resour*, 38(1),173 -186.
- 35 Mack, Y. P. and Rosenblatt, M. (1979), Multivariate k-nearest neighbor density estimates. *Journal of Multivariate*  
36 *Analysis*, 9, 1-15.

- 1 María, C.M., Wenceslao, G.M., Manuel, F.B., José, M.P.S., and Román, L.C., (2004), Modelling of the monthly  
2 and daily behaviour of the discharge of the Xallas river using Box–Jenkins and neural networks methods, J.  
3 Hydrol., 296,38–58.
- 4 Marques, C.A.F., Ferreira, J. Rocha, A. Castanheira, J. P. Gonçalves, Vaz., N., and Dias, J.M., (2006), Singular  
5 spectral analysis and forecasting of hydrological time series, *Physics and Chemistry of the Earth*, 31,1172-1179.
- 6 Martinerie, J.M., Albano, A.M., Mees, A.I., and Rapp, P.E., (1992), Mutual information, strange attractors, and  
7 the optimal estimation of dimension, *Phys. Rev. A*, 45, 7058.
- 8 Mees, A.I., Rapp, P.E., and Jennings, L.S., (1987), Singular-value decomposition and embedding dimension, *Phys.*  
9 *Rev. A*, 36, 340.
- 10 Muttill, N. and Chau, K.W., (2006), Neural network and genetic programming for modelling coastal algal blooms.  
11 *Int. J. Environment and Pollution*, Vol. 28, Nos. 3/4, pp.223–238.
- 12 Nash, J. E. and Sutcliffe, J. V., (1970), River flow forecasting through conceptual models; part I – a discussion of  
13 principles, *J. Hydrol.*, 10, 282–290, 1970.
- 14 Pasternack, G.B., (1999), Does the river run wild? Assessing chaos in hydrological systems. *Adv. Water Resour.*  
15 23, 253–260.
- 16 Procaccia, I., (1988), Complex or just complicated? *Nature*, 333, 498–499.
- 17 Ruelle, D., (1990), *Proc. R. Soc. London, Ser. A* 427,241.
- 18 Salas, J.D., Tabios III, G.Q., and Bartolini, P., (1985), Approaches to multivariate modeling of water resources  
19 time series, *Water Resources Bulletin*, 21(4), 683-708.
- 20 Sauer, T., Yorke, J. A., and Casdagli, M., (1991), Embedology, *J. Stat. Phys.*, 65, 579– 616.
- 21 Schertzer, D., Tchiguirinskaia, I., Lovejoy, S., Hubert, P., Bendjoudi, H., Larchevêque, M., (2002), Which chaos  
22 in the rainfall runoff process? A discussion on Evidence of chaos in the rainfall-runoff process by Sivakumar.  
23 *Hydrol. Sci. J.* 47 (1), 139–147.
- 24 Sivakumar, B., Jayawardena, A. W. and Fernando, T. M. K., (2002), River flow forecasting: use of phase-space  
25 reconstruction and artificial neural networks approaches, *Journal of Hydrology*, 265(1), 225-245.
- 26 Sivapragasam, C., Liong, S.Y. and Pasha, M.F.K., (2001), Rainfall and discharge forecasting with SSA-SVM  
27 approach, *J. Hydroinf.*, 3(7), 141–152.
- 28 Sivakumar, B., Liong, S.Y., and Liaw, C.Y., (1998), Evidence of chaotic behavior in Singapore rainfall, *Journal*  
29 *of the American water resources association*, 34(2), 301-310.
- 30 Sivakumar, B., Berndtsson, R., Olsson, J., Jinno, K., (2001), Evidence of chaos in rainfall-runoff process.  
31 *Hydrological Sciences Journal*, 46(1), 131-146.
- 32 Solomatine, D.P., Maskey, M., and Shrestha, D.L., (2008), Instance-based learning compared to other data-driven  
33 methods in hydrological forecasting, *Hydrological Processes*, 22(2), 275-287.
- 34 Sugihara, G and May, R.M., (1990), Nonlinear forecasting as a way of distinguishing chaos from measurement  
35 error in time series. *Nature* 344, 734– 741.
- 36 Takens, F., (1981), *Dynamical systems and Turbulence*, Lecture note in Mathematics, Vol. 898, Springer, New  
37 York.

- 1 Theiler, J., (1986), Spurious dimension from correlation algorithms applied to limited time-series data. *Phys. Rev.*  
2 *A* 34 (3), 2427–2432.
- 3 Toth, E, Brath, A., and Montanari, A., (2000), Comparison of short-term rainfall prediction models for real-time  
4 flood forecasting, *J of Hydrology* 239, 132–147.
- 5 Tsonis, A.A., (1992), *Chaos: From Theory to Applications*, Plenum Press, New York.
- 6 Wang, W., Van Gelder, P.H.A.J.M., Vrijling, J.K., Ma, J., (2006a), Testing for nonlinearity of streamflow  
7 processes at different timescales. *Journal of Hydrology*, 322, 247-268.
- 8 Wang, W., Van Gelder, P.H.A.J.M., Vrijling, J.K., Ma, J., (2006b), Forecasting Daily Streamflow Using Hybrid  
9 ANN Models, *Journal of Hydrology*, 324, 383-399.
- 10 Yakowitz, S. (1987), Nearest neighbor method for time series analysis. *Journal of Time Series Analysis*, 8 (2),  
11 235-247.
- 12 Yakowitz, S., and Karlsson, M., (1987), Nearest neighbor methods for time series with application to rainfall-  
13 discharge prediction, In: *Stochastic Hydrology*, (eds.) J.B. MacNeil, and G.D. Umphrey, pp. 149-160, Dordrecht ;  
14 Boston : D. Reidel ; Norwell, Mass.
- 15 Yu, P.S., and Tseng, T.Y. (1996), A model to forecast flow with uncertainty analysis, *Hydrological Sciences-*  
16 *Journal*, 41(3), 327-344.
- 17 Yu, X.Y., Liang, S.Y. and Babovic, V., (2004), EC-SVM approach for real-time hydrologic forecasting, *Journal*  
18 *of Hydroinformatics*, 6(3), 209-233.

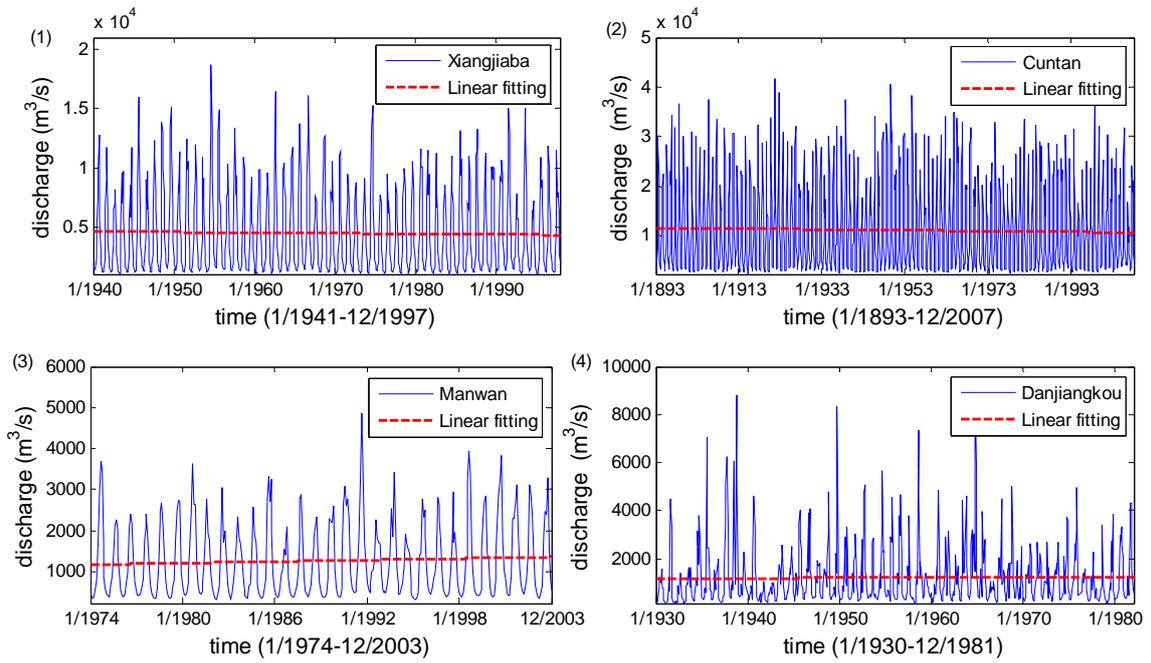


Figure 1. Monthly discharge series of (1) Xiangjiaba, (2) Cuntan, (3) Manwan, and (4) Danjiangkou

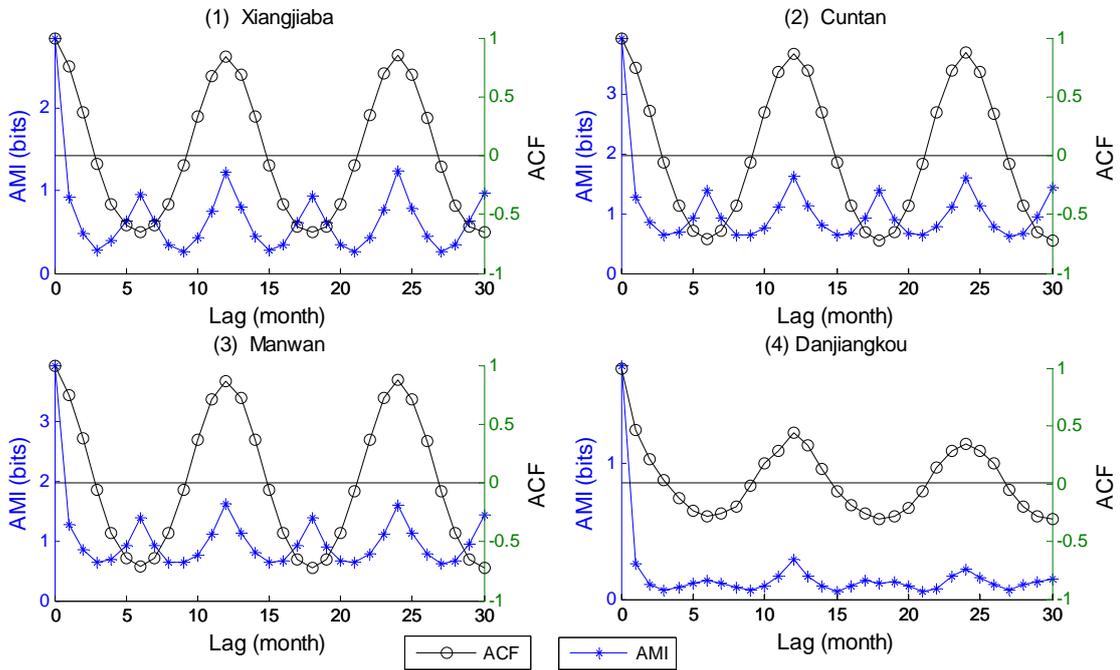


Figure 2. ACF and AMI of (1) Xiangjiaba, (2) Cuntan, (3) Manwan, and (4) Danjiangkou

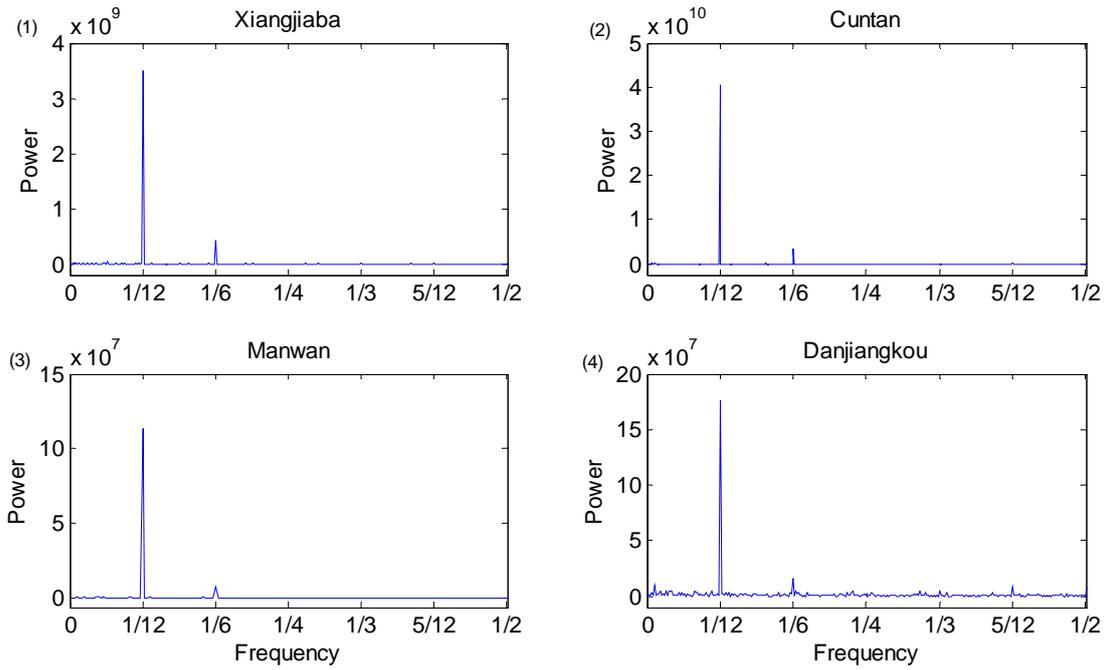


Figure 3. Periodogram of (1) Xiangjiaba, (2) Cuntan (3) Manwan, and (4) Danjiangkou

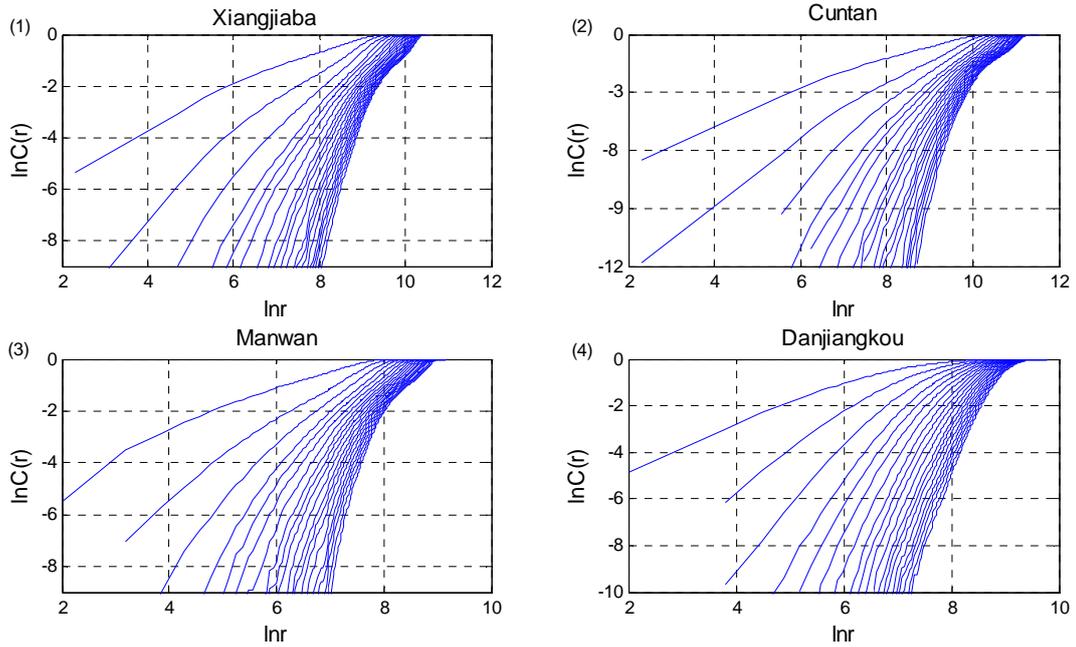
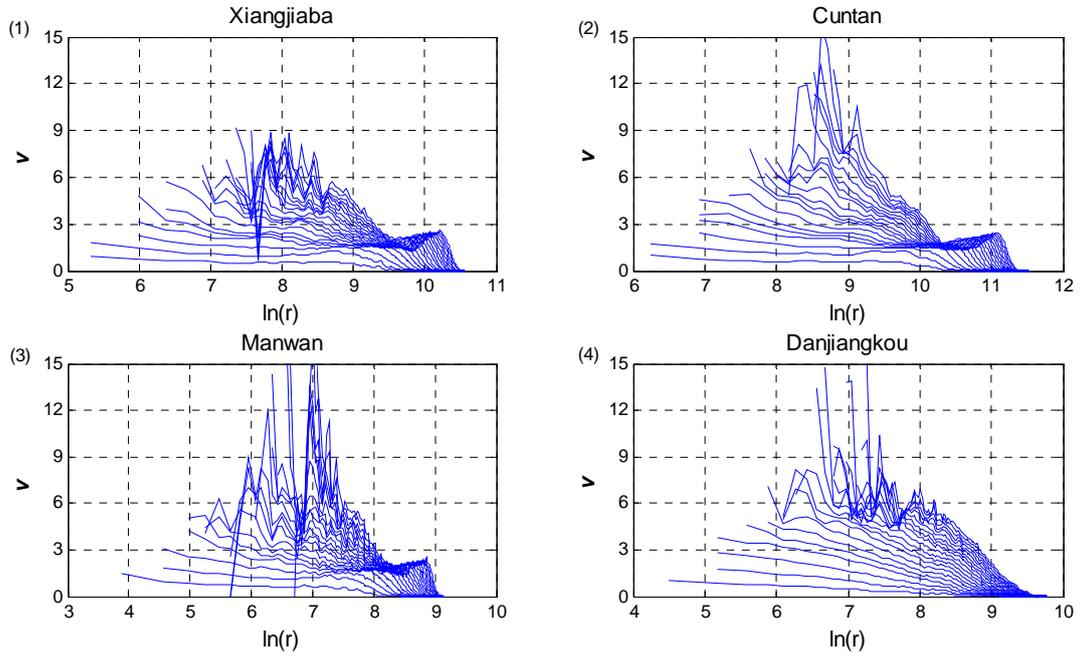


Figure 4.  $\ln C(r)$  versus  $\ln r$  plot for (1) Xiangjiaba, (2) Cuntan (3) Manwan, and (4) Danjiangkou

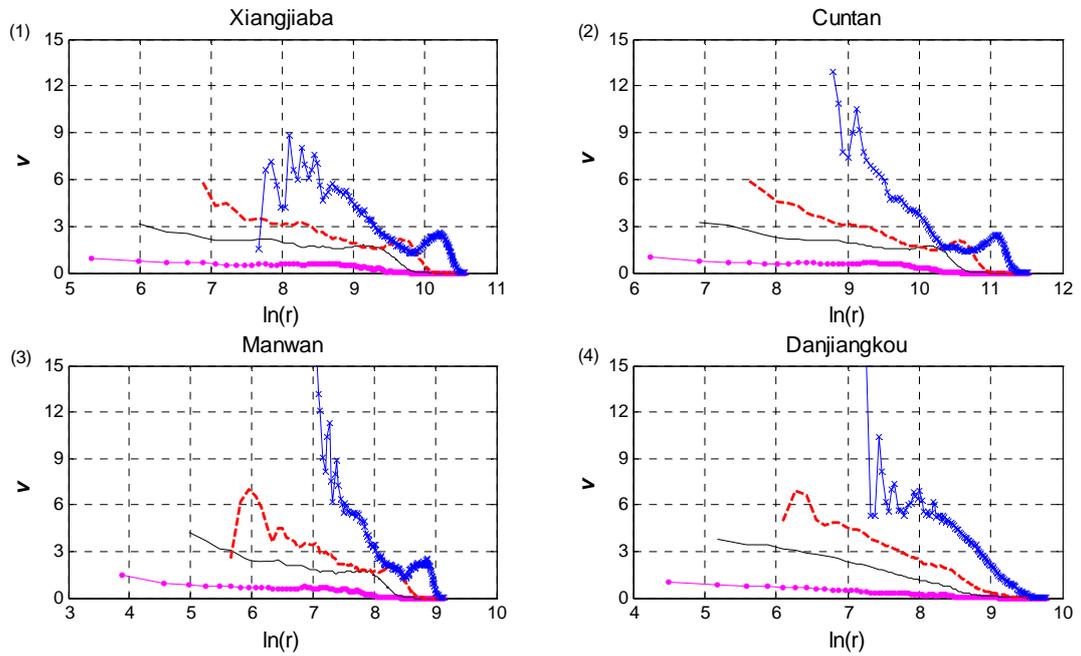
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Figure 5. The estimation of correlation dimension ( $d_2$ ) for (1) Xiangjiaba, (2) Cuntan (3) Manwan, and (4) Danjiangkou.  $\tau = 3$  and  $m$  is at the interval of  $[1,20]$  (increasing from bottom to top in each pane).



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Figure 6. The estimation of correlation dimension ( $d_2$ ) for (1) Xiangjiaba, (2) Cuntan (3) Manwan, and (4) Danjiangkou.  $\tau = 3$  and  $m$  is at 1(solid line with dots),4(solid line),8 (dashed line), and 20 (solid line with crosses).

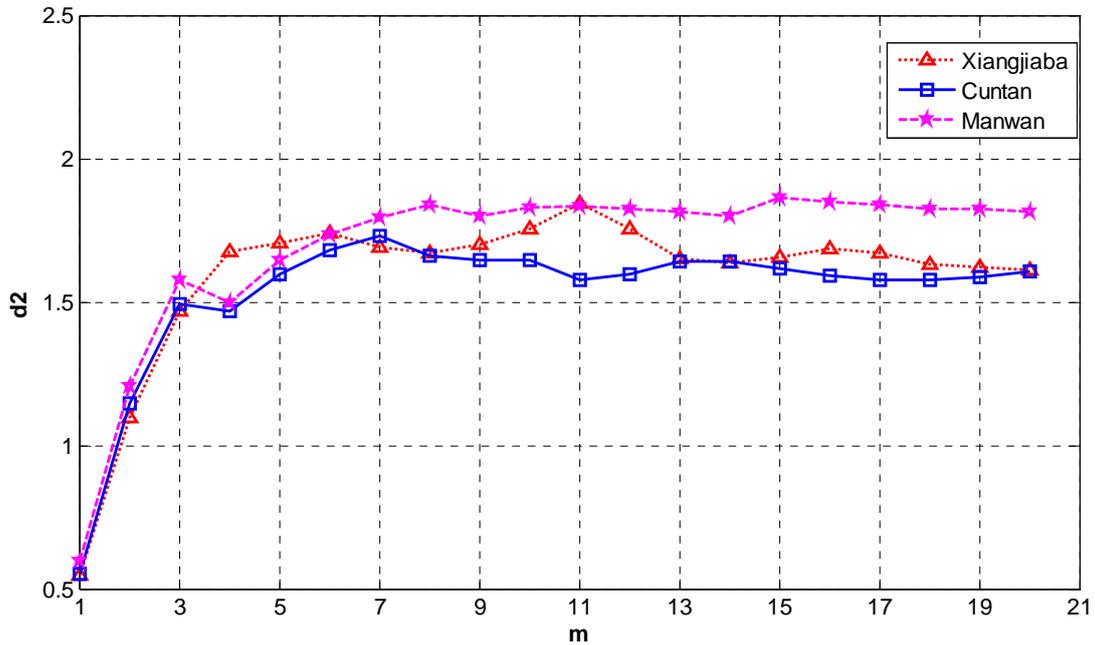


Figure 7. Relationship between  $d_2$  and  $m$  for monthly discharges of Xiangjiaba, Cuntan, and Manwan

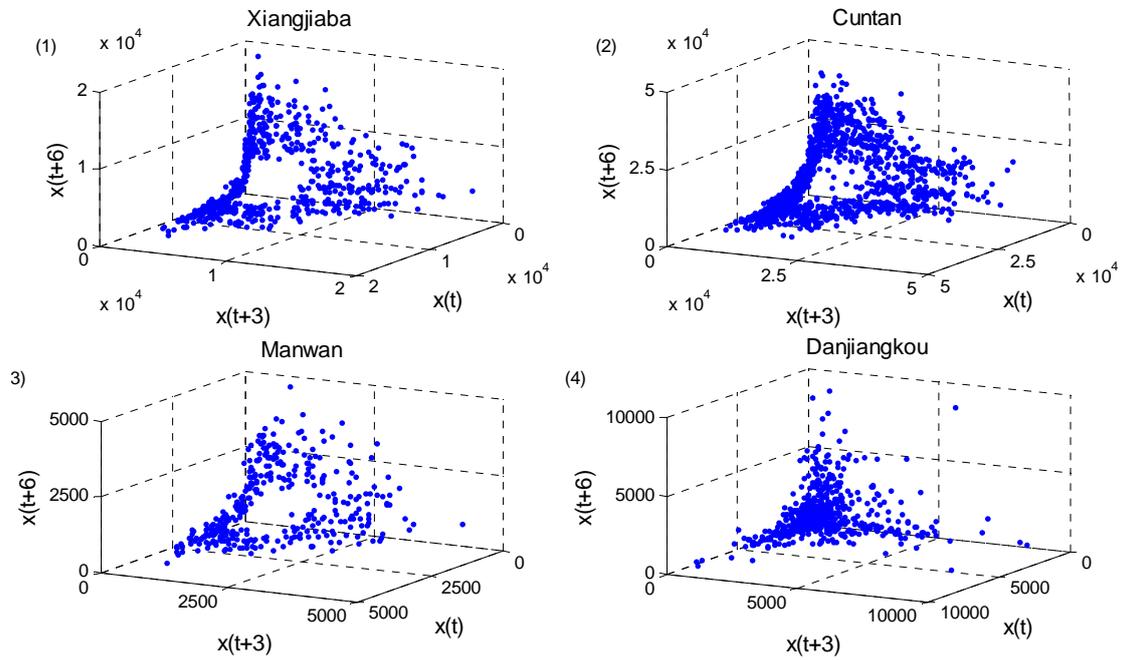
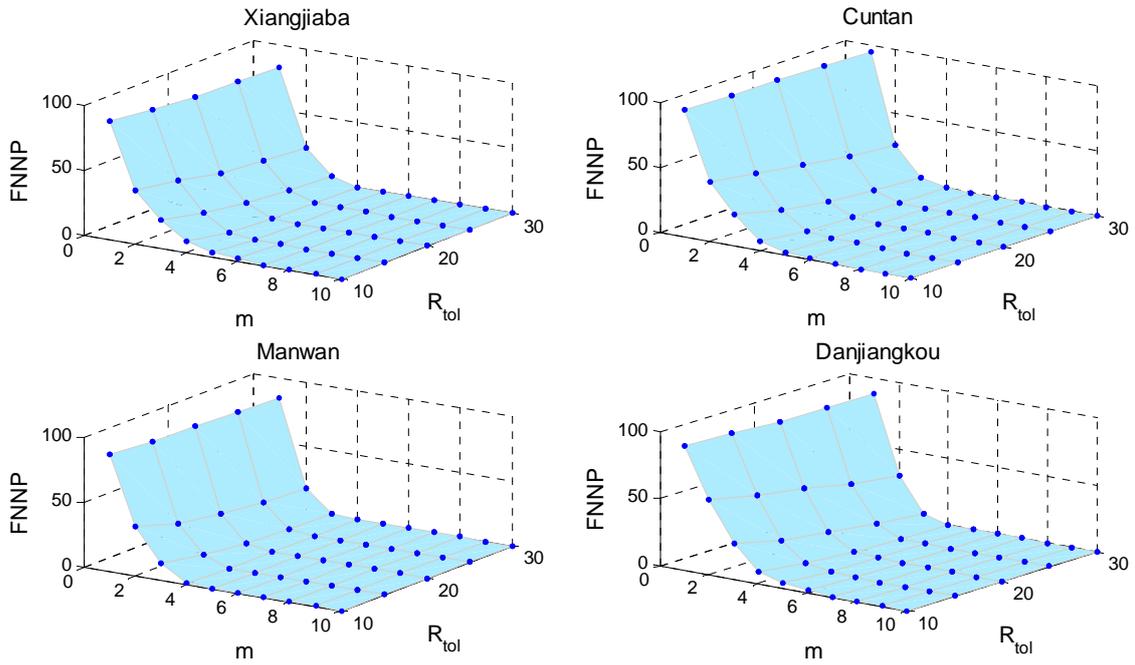
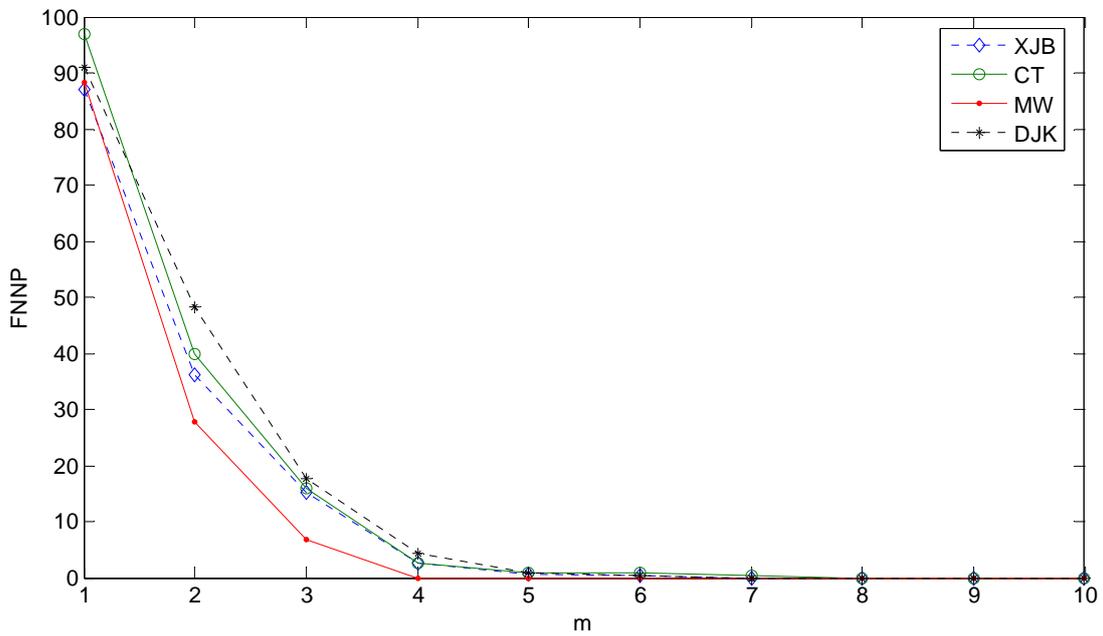


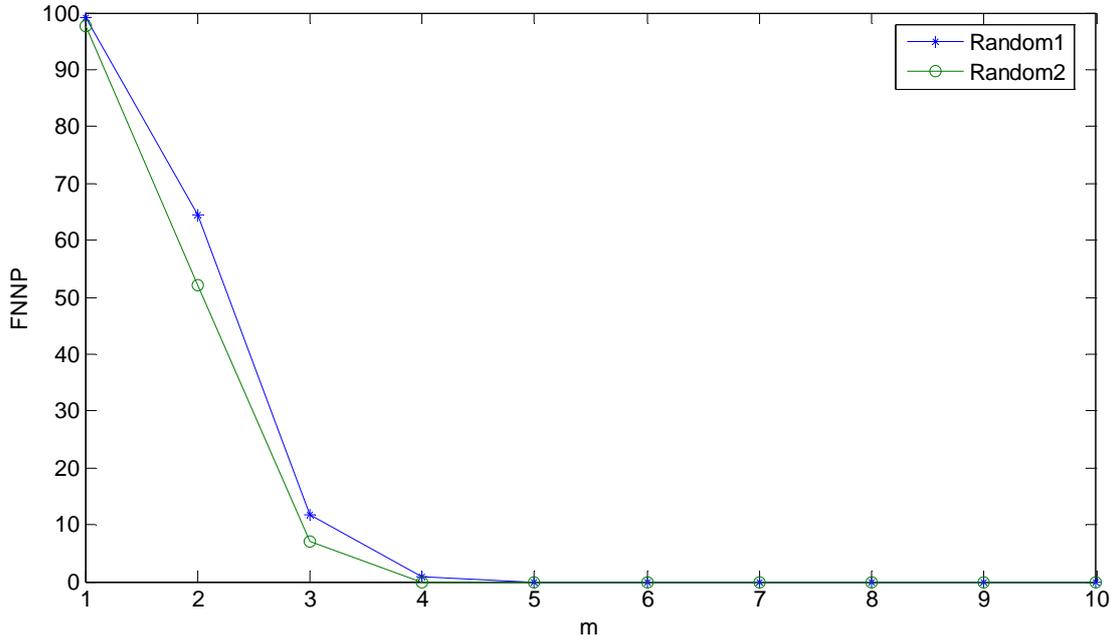
Figure 8. Phase portraits for (1) Xiangjiaba, (2) Cuntan (3) Manwan, and (4) Danjiangkou when  $m = 3$  and  $\tau = 3$ .



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2 Figure 9. FNNP plots for (1) Xiangjiaba, (2) Cuntan (3) Manwan, and (4) Danjiangkou when  $\tau = 3$  and  $R_{tol}$  is 10  
3 to 30.

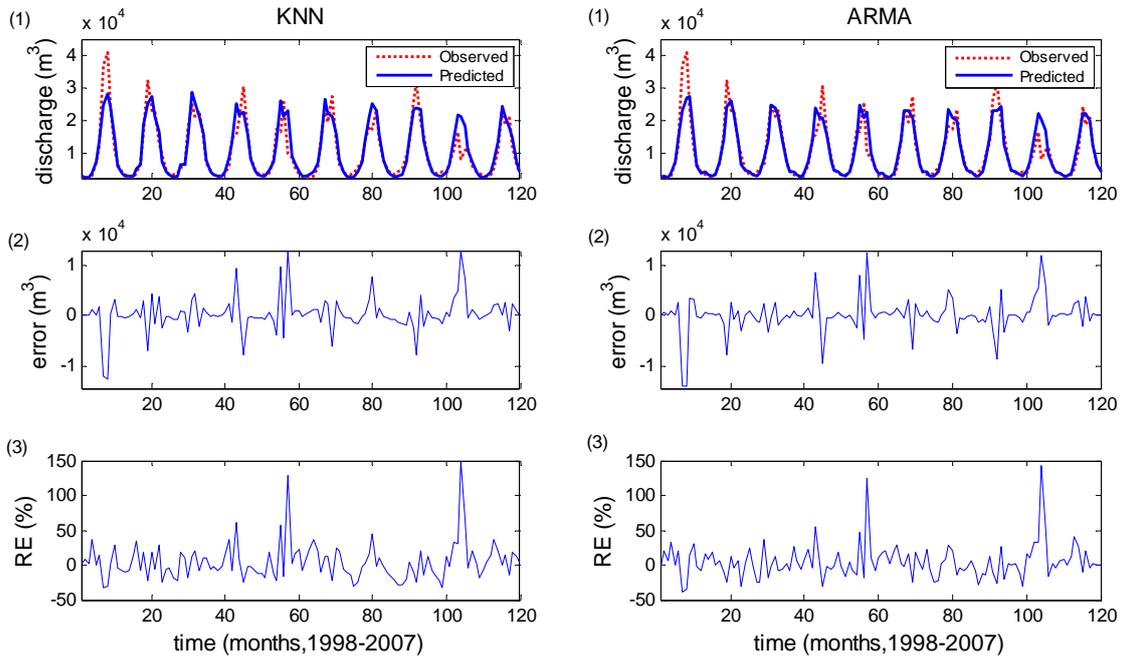


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5 Figure 10. FNNP for Xiangjiaba, Cuntan, Manwan, and Danjiangkou when  $\tau = 3$  and  $R_{tol} = 15$ .  
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Figure 11. FNNP for Random process 1 with a data size of 1380 and Random process 2 with a data size of 696



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Figure 12. KNN (left column) and ARMA (right column) model results for one-day-ahead forecast on Cuntan: (1) observed vs. predicted (2) error (observed – predicted) and (3) RE (representing Relative Error)

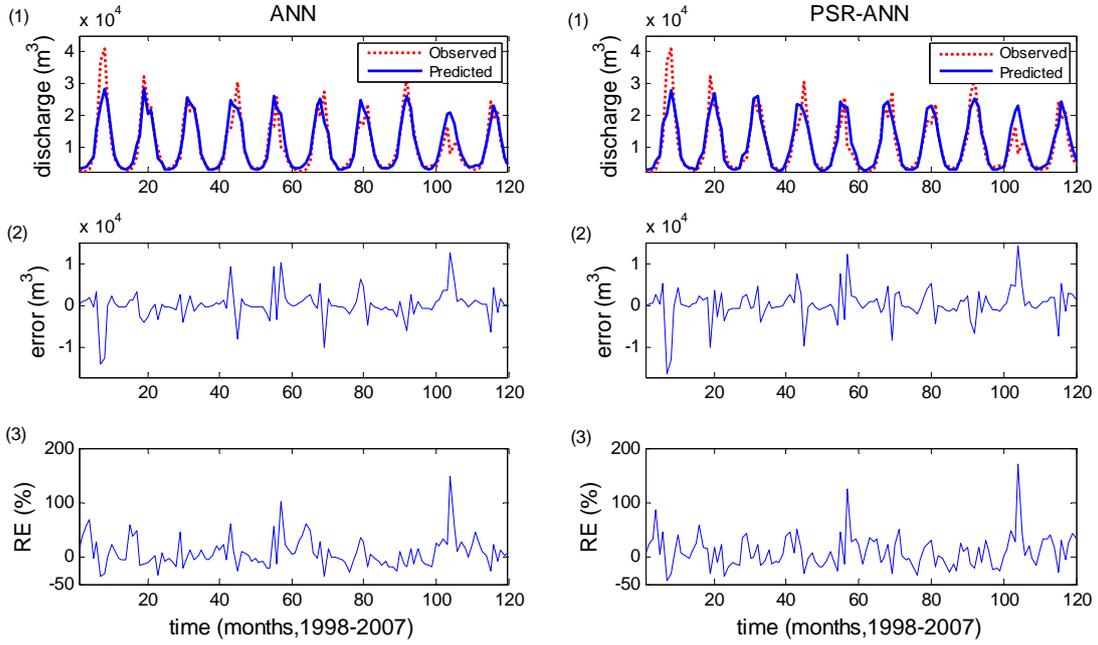


Figure 13. ANN (left column) and ANN-PSR (right column) model results for one-day-ahead forecast on Cuntan: (1) observed vs. predicted (2) error (observed – predicted) and (3) RE (representing Relative Error)

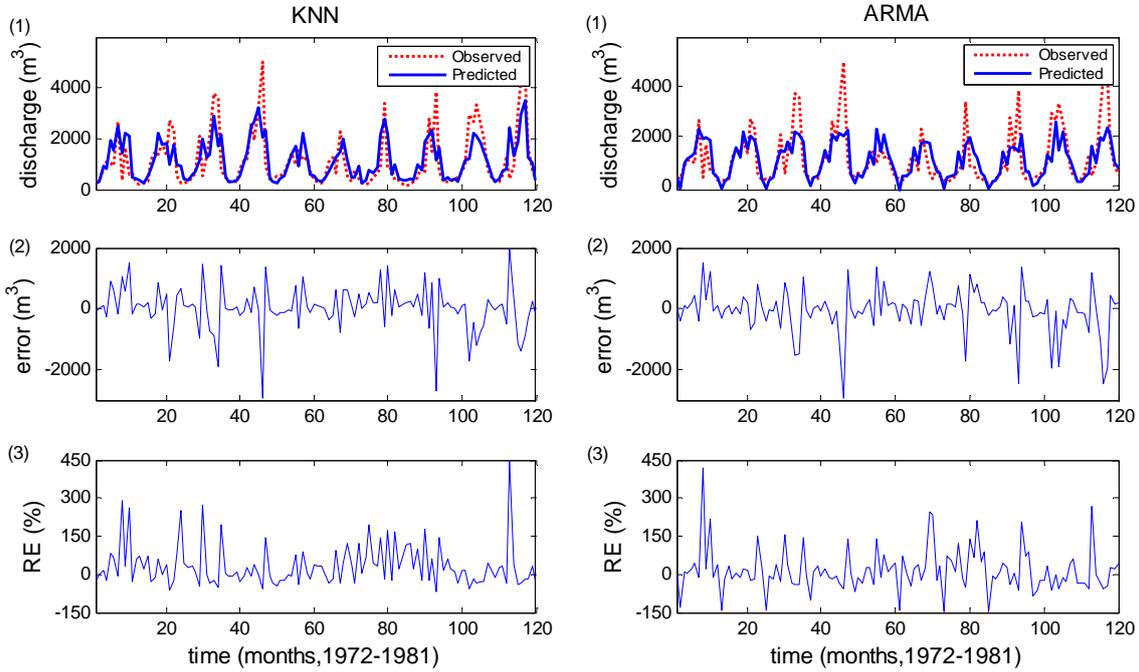


Figure 14. KNN (left lists) and ARMA (right lists) model results for one-day-ahead forecast on Danjiangkou: (1) observed vs. predicted (2) error (observed – predicted) and (3) RE (representing Relative Error)

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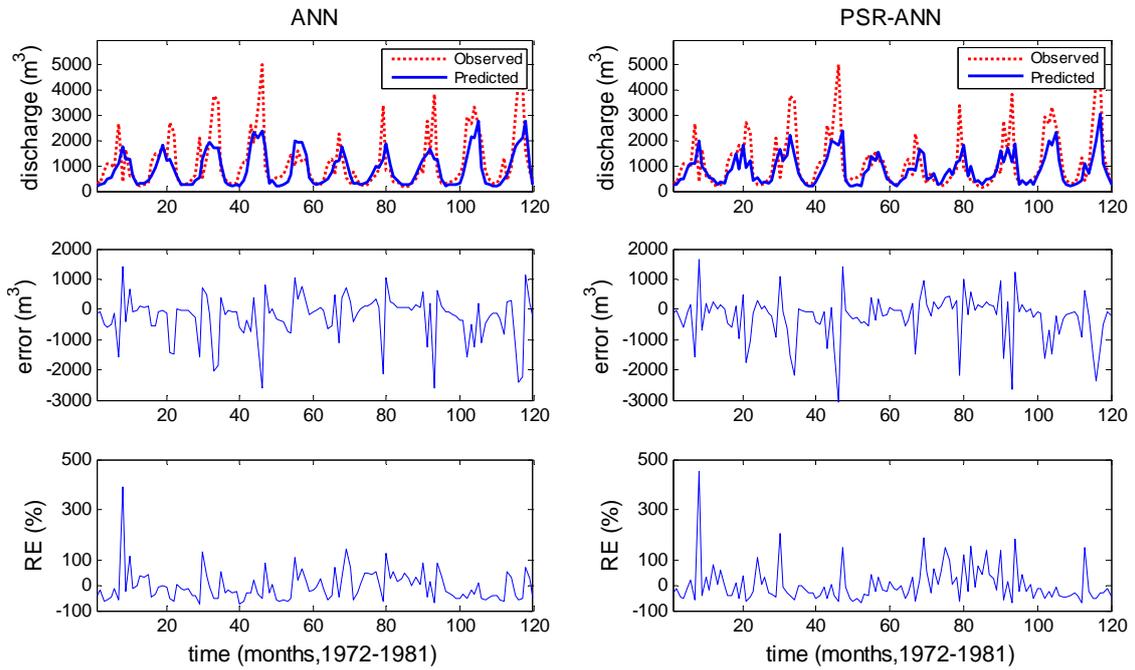


Figure 15. KNN (left lists) and ARMA (right lists) model results for one-day-ahead forecast on Danjiangkou: (1) observed vs. predicted (2) error (observed – predicted) and (3) RE (representing Relative Error)

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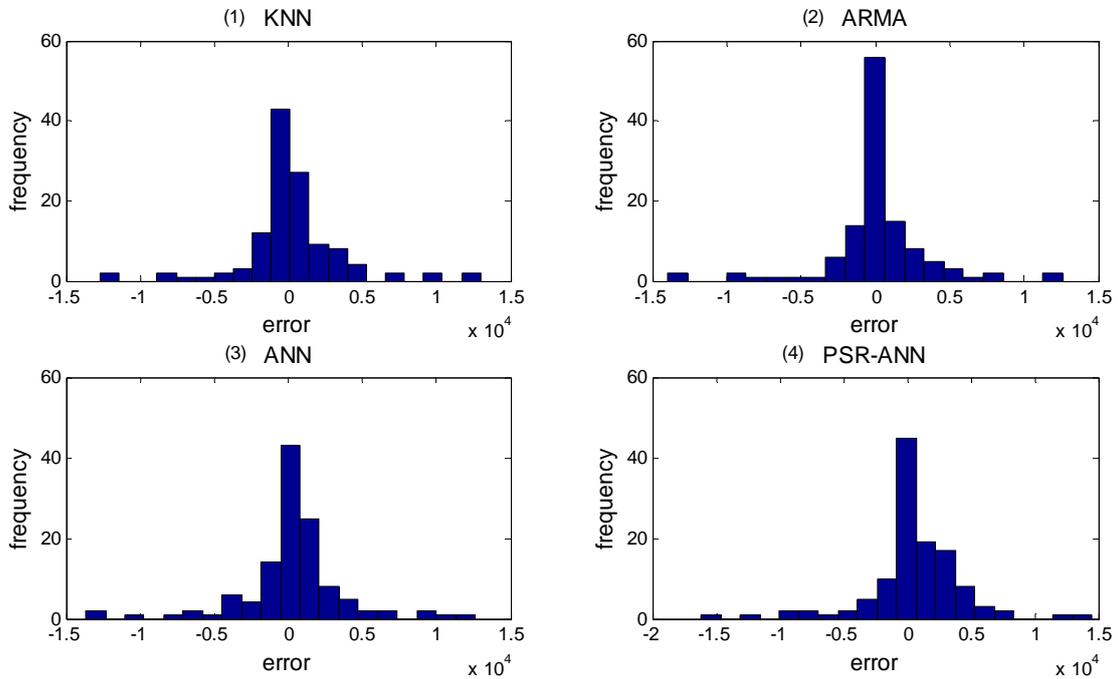


Figure 16. Histograms of errors of (1) K-NN, (2) ARMA, (3) ANN, and (4) ANN-PSR for Cuntan

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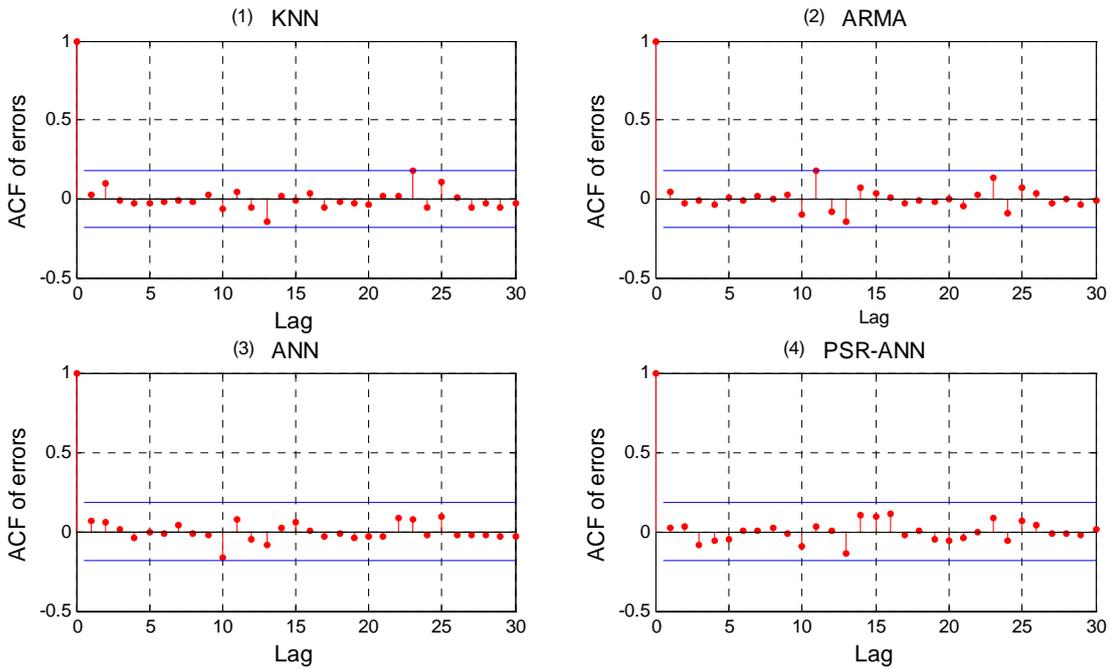


Figure 17. ACF of errors of (1) K-NN, (2) ARMA, (3) ANN, and (4) ANN-PSR for Cuntan

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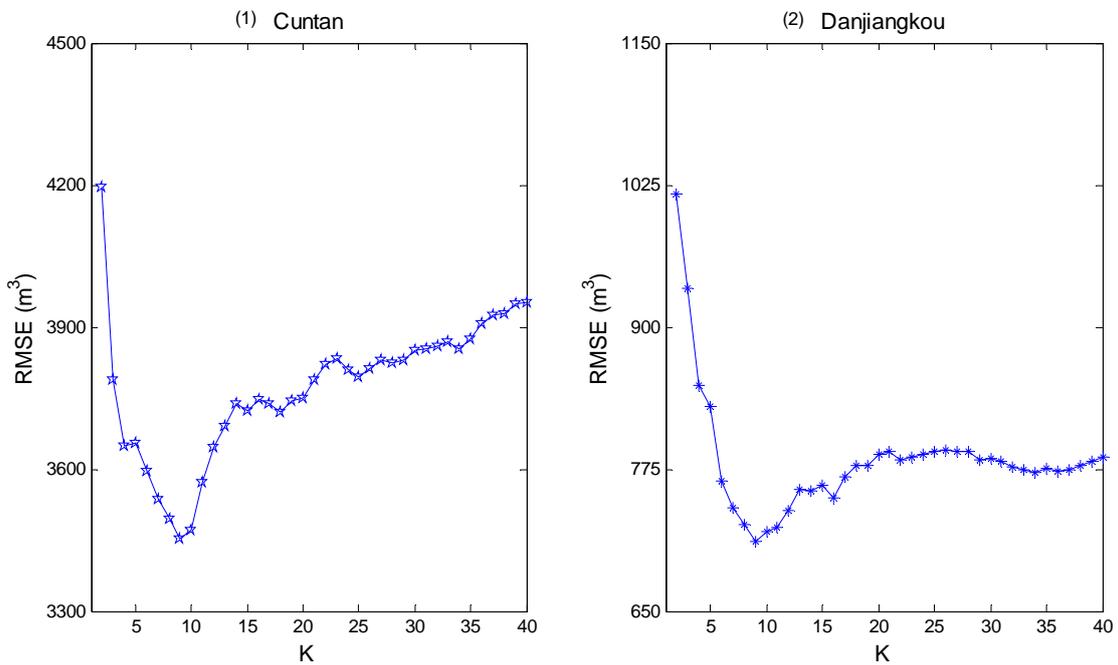
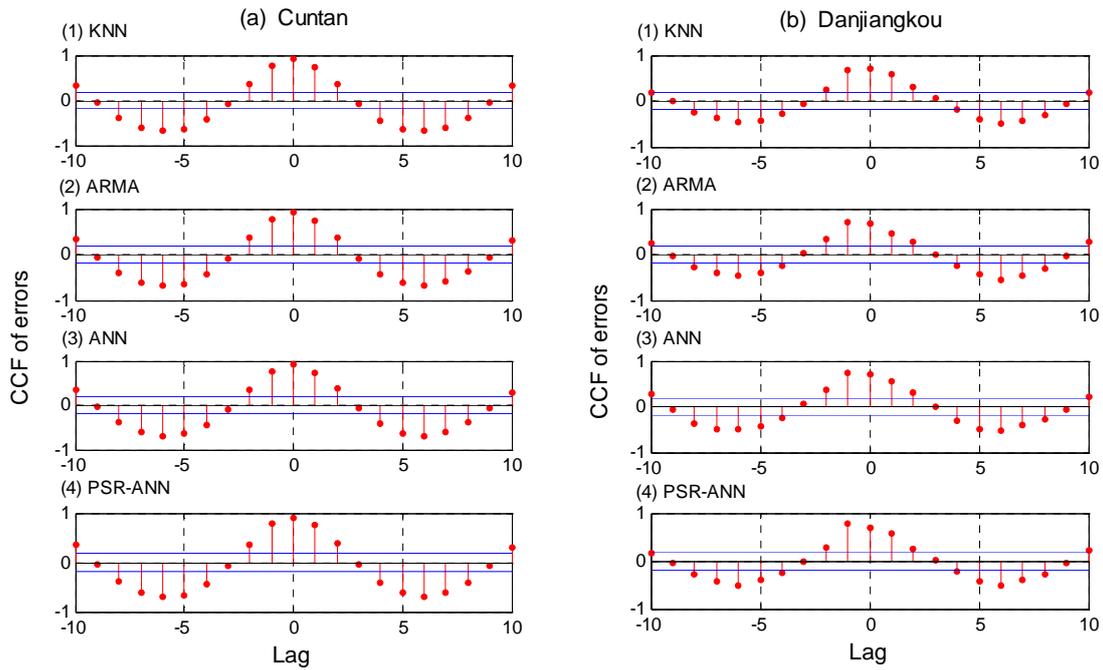
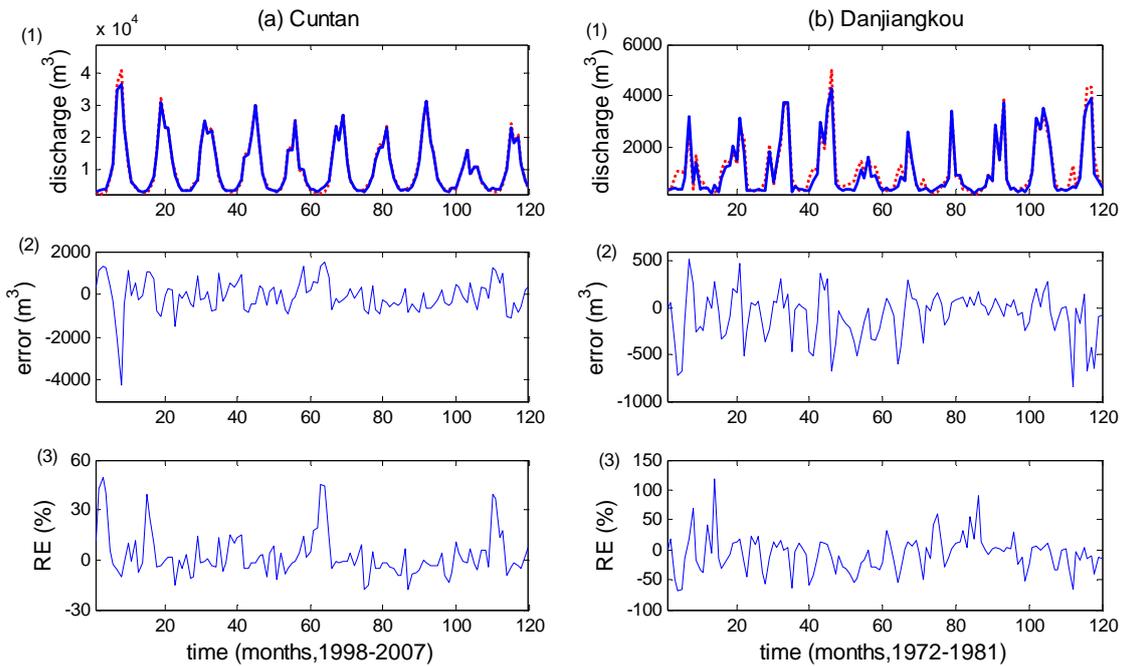


Figure 18. The check of stability of  $k$  in KNN method for (1) Cunan and (2) Danjiangkou

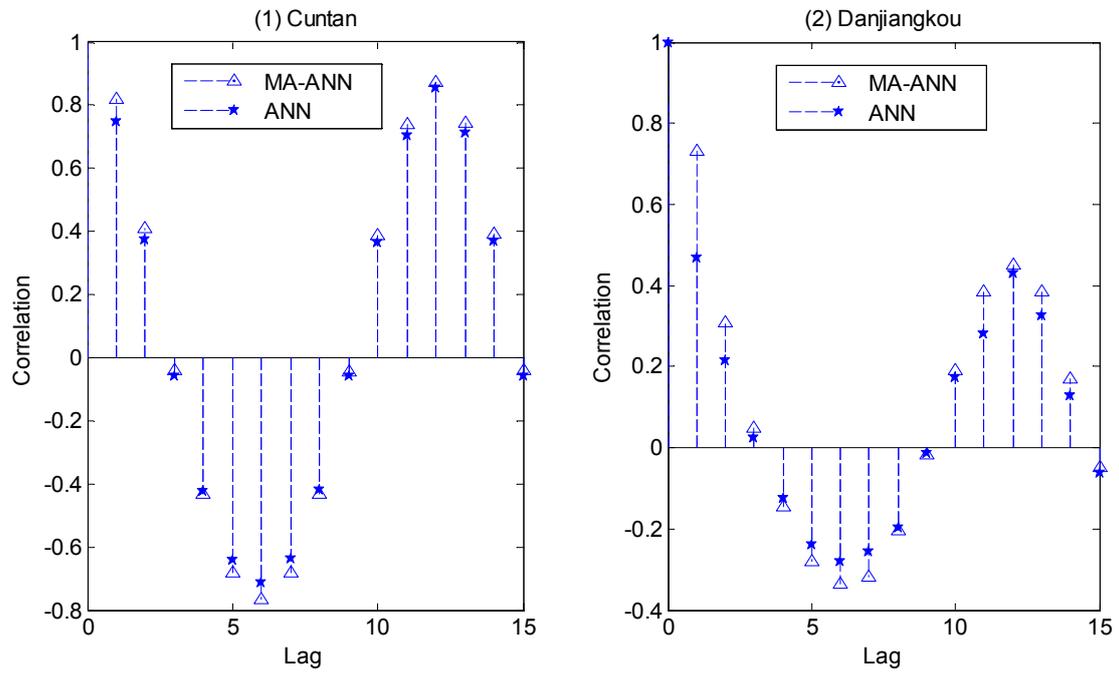
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Figure 19. The check of the prediction lag effect via CCF between observed and predicted values. (a) for Cuntan and (b) for Danjiangkou ((1) K-NN, (2) ARMA, (3) ANN, and (4) ANN-PSR)



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Figure 20. MA-ANN model results of one-day-ahead forecast for (a) Cuntan and (b) Danjiangkou. ((1) observed vs. predicted (2) error (observed – predicted) and (3) RE (representing Relative Error) )



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Figure 21. CCFs between inputs and outputs from the MA-ANN model and the ANN model for (1) Cuntan and (2) Danjiangkou.

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Table 1. Statistical parameters for training, testing, and validation sets

Watershed and datasets	Statistical parameters					Watershed area and data period
	Mean	Standard deviation	Minimum	Maximum	Range	
<b>Xiangjiaba</b>						
Original data	4538.2	3671.5	1100	18700	17600	Area: 45.88×10 <sup>4</sup> km <sup>2</sup> Data period: January 1940 to December 1997
Training set	4678.2	3832.4	1130	18700	17570	
Cross-validation set	4297.6	3411.2	1110	15300	14190	
Testing set	4475.2	3552.3	1100	15000	13900	
<b>Cuntan</b>						
Original data	11118	8649.7	2250	41700	39450	Area:  Data period: January 1893 to December 2007
Training set	11342	8892.6	2250	41700	39450	
Cross-validation set	10766	8195.7	2420	34800	32380	
Testing set	10774	8488.7	2610	40800	38190	
<b>Manwan</b>						
Original data	1268.2	908.66	316	4860	4544	Area: 11.45×10 <sup>4</sup> km <sup>2</sup> Data period: January 1974 to December 2003
Training set	1192.7	862.29	329	3710	3381	
Cross-validation set	1371.3	944.26	373	4860	4487	
Testing set	1300.1	943.19	316	3944	3628	
<b>Danjiangkou</b>						
Original data	1203.3	1285.1	139	9010	8871	Area: 9.5×10 <sup>4</sup> km <sup>2</sup> Data period: January 1930 to December 1981
Training set	1203	1348.1	140	8800	8660	
Cross-validation set	1256.1	1310.7	139	9010	8871	
Testing set	1130	1053.5	159	4980	4821	

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Table 2 Statistics of ANN model performance using various training algorithms with 30 time runs

Training algorithm and performance	Statistical parameters				
	Mean	Standard deviation	Minimum	Maximum	Range
<b>LM</b>					
RMSE	844.9	114.1	774.6	1429.5	654.9
CE	0.3	0.2	-0.9	0.5	1.3
PI	0.3	0.2	-1.0	0.4	1.4
Training time(sec.)	0.7	0.2	0.5	1.5	0.9
<b>PSO</b>					
RMSE	795.9	29.1	763.1	912.5	149.4
CE	0.4	0.0	0.2	0.5	0.2
PI	0.4	0.0	0.2	0.4	0.2
Training time(sec)	74.6	1.0	70.9	76.0	5.0
<b>LM-GA</b>					
RMSE	808.7	24.1	760.1	861.0	100.9

CE	0.4	0.0	0.3	0.5	0.1
PI	0.4	0.0	0.3	0.4	0.1
Training time(sec.)	19.3	0.7	18.3	21.0	2.7

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2 Table 3. Comparison of model performance for one-step-ahead forecast on validation data and  
3 model parameters

Watershed	Model	RMSE	CE	PI	Model parameters		
					Architecture	( $\tau, m, k$ )	(p, q)
Cuntan							
	ARMA	3498.60	0.83	0.69			(11,11)
	ANN*	3557.10	0.82	0.68	(12,13,1)		
	K-NN	3495.50	0.83	0.69		(5,7,8)	
	ANN-PSR*	3878.60	0.79	0.61	(7,9,1)		
Danjiangkou							
	ARMA	763.81	0.47	0.44			(11,10)
	ANN*	794.98	0.43	0.39	(12,5,1)		
	K-NN	741.58	0.50	0.47		(5,6,7)	
	ANN-PSR*	814.74	0.40	0.36	(6,9,1)		

4 Note:\* the index of performance is an average over the best 10 runs of total 30 runs.

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7 Table 4. Improvement of model performance for one-step-ahead forecast from K-NN and MA-  
8 ANN models and model parameters

Watershed	Model	RMSE	CE	PI	Model parameters		
					Architecture	( $\tau, m, k$ )	Memory length for the moving average
Cuntan							
	MA-ANN	687.37	0.99	0.99	(12,13,1)		3
	K-NN	3453.90	0.83	0.69		(5,7,9)	
Danjiangkou							
	MA-ANN	268.71	0.93	0.93	(12,5,1)		6
	K-NN	710.73	0.54	0.51		(5,6,9)	

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