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Large Scale Operational Soil Moisture Mapping from Passive MW Radiometry: SMOS product evaluation in Europe & USA

Khidir Abdalla Kwal Deng^{1,2}, Salim Lamine³, Andrew Pavlides⁴, George P. Petropoulos^{4,5,*}, Yansong Bao^{1,2}, Prashant K. Srivastava⁶, Yuanhong Guan⁷

¹ Collaborative Innovation Center on Forecast and Evaluation of Meteorological Disasters, Nanjing University of Information Science & Technology, Nanjing 210044, China, Email: khidir14332@yahoo.com

² School of Atmospheric physics, Nanjing University of Information Science and Technology, Nanjing 210044, China, Email: ysbao@nuist.edu.cn

³ Faculty of Natural Sciences, Life and Earth Sciences, University Akli Mohand Oulhadj of Bouira, 10000, Bouira, Algeria, Email: salim.lamine@gmail.com

⁴ School of Mineral & Resources Engineering, Technical University of Crete, Crete, Greece, Email: apavlides24@yahoo.com

⁵ Department of Soil & Water Resources, Institute of Industrial & Forage Crops, Hellenic Agricultural Organization "Demeter", Larisa, Greece, Email: petropoulos.george@gmail.com

⁶ Institute of Environment and Sustainable Development, Banaras Hindu University, India, Email: prashant.iesd@bhu.ac.in

⁷ School of Mathematics and Statistics, Nanjing University of Information Science and Technology, Nanjing 210044, China, Email: guanyh@nuist.edu.cn

*. Correspondence: petropoulos.george@gmail.com

ABSTRACT

Earth Observation (EO) allows deriving from a range of sensors, often globally, operational estimates of surface soil moisture (SSM) at range of spatiotemporal resolutions. Yet, an evaluation of the accuracy of those products in a variety of environmental conditions has been often limited. In this study the accuracy of the SMOS SSM global operational product across 2 continents (USA, and Europe) is investigated. SMOS predictions were compared against near concurrent *in-situ* SSM measurements from the FLUXNET observational network. In total, 7 experimental sites were used to assess the accuracy of SMOS derived soil moisture for 2 complete years of observations (2010 to 2011). The accuracy of the SMOS SSM product is investigated in different seasons for the seasonal cycle as well as different continents and land types. Results showed a generally reasonable agreement between the SMOS product and the *in-situ* soil moisture measurements in the 0-5 cm soil moisture layer. Root Mean Square Error (RMSE) in most cases was close to $0.1 \text{ m}^3 \text{ m}^{-3}$ (minimum $0.067 \text{ m}^3 \text{ m}^{-3}$). With a few exceptions, Pearson's correlation coefficient was found up to approx. 55%. Grassland, shrublands and woody savanna land cover types attained a satisfactory agreement between satellite derived and *in-situ* measurements but needleleaf forests had lower correlation. Better agreement was found for the grassland sites in both continents. Seasonally, summer and autumn underperformed spring and winter. Our study results provide supportive evidence of the potential value of this operational product for meso-scale studies in a range of practical applications, helping to address key challenges present nowadays linked to food and water security.

Keywords: surface soil moisture, earth observation, operational products, SMOS, validation

36 **1. Introduction**

37 Soil moisture corresponds to water in both the uppermost layer of the land surface - called Surface Soil
38 Moisture (SSM) - and the root zone or vadose area. This parameter is strongly affected by many factors
39 such as soil texture, organic materials, and topography as well as land use/land cover and rainfall
40 (Srivastava et al. 2016a; Raffelli et al. 2017) Soil moisture, particularly SSM plays a significant role in
41 the distribution of the mass and energy fluxes between the land and the atmosphere, and it controls the
42 different components of the water and energy balance (Seneviratne et al. 2010; Bao et al. 2018).
43 Furthermore, it is a key state variable in organizing the natural ecosystems and biodiversity (Vereecken et
44 al. 2008), also important to modeling extreme events such as flooding or landslides prediction (Bittelli et
45 al. 2012; Wanders et al. 2014), drought monitoring (Sánchez-Ruiz et al. 2014), and numerical weather
46 prediction (De Rosnay et al., 2013). Considering many aspects in life such as food security and water
47 resources management, it is essential for agriculture and irrigation management practices. Particularly, in
48 developing irrigation management practices for more crop production and optimum use of water
49 resources especially in arid and semi-arid regions (Rotzer et al. 2014; Canone et al. 2015; Brocca,
50 Ciabatta, et al. 2017; Canone et al. 2017). Thus, large scale SSM accuracy evaluation spatially and
51 temporally represents an important topic to be investigated.

52 SSM point based measurements at particular locations fail to effectively capture this variability. There are
53 different approaches used for soil moisture measurements (a good review can be found for example in
54 Petropoulos et al. 2015a) , including the establishment of relevant operational networks (Petropoulos et al.
55 2017). *In-situ* techniques, such as the gravimetric Time Domain Reflectometry (TDR) and the Frequency
56 Domain Reflectometry (FDR) techniques (Brocca et al. 2017) provide accurately SSM. However, they are
57 of too sparse spatial coverage to characterize the spatiotemporal features of soil moisture at large-scale
58 (Crow et al. 2012; Pierdicca et al. 2012). Newly developed techniques such as cosmic ray and GPS
59 moderately address this issue (Dorigo et al. 2013).

60 Earth Observation (EO) provides promising methods to survey SSM at large scale at satisfactory
61 spatiotemporal resolution (Srivastava et al. 2016b; Petropoulos et al. 2018a). In the past two decades
62 immense progress has been achieved on developing soil moisture products by using EO from microwave,
63 optical and thermal satellite sensors (for a review see e.g. (Petropoulos et al. 2018b). Several microwave
64 instruments were launched for developing SSM global products from active/passive microwave signals.
65 Currently, L-band microwave sensors are considered the most promising for SSM estimation. The Soil
66 Moisture Ocean Salinity (SMOS) mission of European Space Agency (ESA) carries the first operational
67 L-band radiometer to measure SSM at spatial resolution of ~40 km (Kerr et al. 2012; Djamai et al. 2015).
68 Currently, the satellite Scatterometers of the European Remote Sensing (ERS-1/2) and the Advanced
69 Scatterometer (ASCAT) onboard of the Meteorological Operational satellite program Metop-A and
70 Metop-B (2007–2014) provide soil moisture retrievals at global scale.

71 In order to obtain long term soil moisture estimation at global scale, passive and active microwave soil
72 moisture products have been used in combination. For example, a method to derive soil moisture from
73 SMAP/Sentinel-1 data such as SMAP L-band brightness temperatures and Copernicus Sentinel-1 C-band
74 backscatter coefficients has been developed (Entekhabi et al. 2017). Likewise, there are efforts to merge
75 the passive and active soil moisture products under the European Space Agency Climate Change Initiative
76 soil moisture product (CCI SM), in an attempt to generate a long term global scale soil moisture record
77 (Liu and Parinussa et al. 2011; Draper et al. 2012) The Water Cycle Observation Mission (WCOM)

78 satellite is being developed by the Chinese Academy of Sciences to combine the passive and active
79 microwave sensors and is expected to be launched in 2020 (Shi et al. 2014).

80 Due to its lower sensitivity to surface roughness and vegetation cover, the L-band is more appropriate for
81 assessing soil moisture conditions (Calvet *et al.*, 2011). This makes the L- the most suitable microwave
82 band for soil moisture measurement from space. In the recent years, the product has been evaluated by
83 various studies in several geographical regions around the globe like USA (Zhuo et al. 2015), Argentina
84 (Grings et al. 2015), Europe (Roßter et al. 2014; González-Zamora et al. 2015), China (Cui et al. 2017),
85 India (Chakravorty et al. 2016) and West-Africa (Louvet et al. 2015).

86 Despite the major importance of soil moisture and measuring it effectively in global scale, a systematic
87 presentation of the accuracy of the MIRAS instrument of SMOS has been examined so far by very few
88 studies (Petropoulos et al. 2014; Fascetti et al. 2014; Petropoulos et al. 2015b; Djamai et al. 2015; Liu et
89 al. 2018; Chen et al. 2018). The motivation of our study was to investigate the accuracy of soil moisture
90 measurements by SMOS in the Northern hemisphere. SMOS SSM is acquired by using remote sensing
91 through indirect measurement techniques. There are many factors influencing their retrievals (e.g. radio
92 interference, vegetation cover, soil roughness, etc. (see for example Petropoulos et al. 2014). Therefore,
93 comprehensive evaluation of those operational products through all the seasons on different vegetation
94 cover types is highly required, so that the data provider and the user can clearly understand the
95 uncertainties associated with the data and assist in further algorithm development (Srivastava et al. 2014).

96 Although a number of studies have been focused on evaluation of SMOS, there are rare studies available
97 on assessment of products over the Northern hemisphere. In this context, this study explores SMOS soil
98 moisture product accuracy in different seasons and variety of land cover types at selected sites belonging
99 to the FLUXNET global *in-situ* measurements network to investigate the different factors that might
100 influence the accuracy of the soil moisture product estimations. A better understanding of MIRAS SSM
101 data can lead to rapid developments in important areas of the economy, such as agriculture, monitoring
102 plant growth as well as food and water security.

103

104 **2. Data description**

105 **2.1 In-situ measurements**

106 FLUXNET (<http://fluxnet.ornl.gov/obtain-data>) is the largest global network of micrometeorological
107 fluxes and ancillary parameters (Baldocchi et al. 1995) in the regional and global scale. SSM is measured
108 at 30-min intervals using standardized instrumentation across sites. After data are collected standard
109 procedures for error corrections, gap-filling and quality control take place to make sure the data are
110 consistent for all sites and datasets. Erroneous data measurements with obvious instrument errors are
111 removed from the in-situ data.

112 In this study, in-situ data for the years 2010 and 2011 were acquired from seven sites. Three of those sites
113 were situated in Europe (AGU, LJU, and MAU) and four were in the United States (ME2, VAR, TON,
114 WHS). Only sites with continuous long term datasets, at surfaces top 5cm depth were selected. Another
115 factor during the selection of sites was homogeneity in the land cover type. To avoid any mixed pixel
116 effects on the overall performance, satellite pixels are chosen over the FLUXNET towers having the
117 largest homogenous land cover.

118 The 7 sites selected in this study are: ES Agu, US-WHS & ES-LJu —open shrubland, US-Me2—
119 Evergreen Needle-Leaf Forest, US-Var —grasslands, FR-MAU —croplands. For FR-MAU, only data
120 from 2011 were available. All in-situ data were obtained from the FLUXNET website and where possible,
121 verified by the site manager above.

122

123 2.2 SMOS Soil Moisture Product

124 The SMOS mission is a part of European Space Agency. It is the first L-band microwave satellite devoted
125 to provide global measurements of soil moisture over land and ocean salinity by observing natural
126 microwave emissions from the earth surface. The SMOS satellite was launched in November 2009, its
127 orbit is 763km which is approximately circular with a 6 a.m. (ascending) and 6 p.m. (descending)
128 equatorial local crossing time and still works surpassing 5 year its proposed service period.

129 The interferometric radiometer onboard of SMOS satellite operates in the L-band microwave. The SMOS
130 platform main instrument is Microwave Imaging Radiometer with Aperture Synthesis (MIRAS), a dual
131 polarized 2-D interferometer that records emitted energy from earth surface in microwave L-band (1.4
132 GHz). It is aimed to provide near-surface soil moisture estimations with global coverage, a three days
133 revisit time at the equator and approximately daily at the pole, spatial resolution of around 40 km (Kerr et
134 al. 2001). The SMOS SSM products are defined on the Icosahedral Snyder Equal Area projection (ISEA
135 4H9 grid) with aperture 4, resolution 9. The shape of cells is a hexagon (Srivastava et al. 2016a). Its
136 mission expected accuracy of 4% which expected to be achievable over relatively uniform area (Panciera
137 et al. 2011). The soil moisture retrievals evaluated in this study are the SMOS products version (v05)
138 image granules which were acquired from Eoli-SA portal covering the full years of 2011 and 2012.

139

140 3. Methods

141 *In-situ* measurements recorded in FLUXNET at the time closer of SMOS overpass were selected for the
142 comparisons performed in this study. After quality assessment, the data values were extracted (Excel
143 Macro VBA) and assigned to point shapefiles of the study site (Tabular join in ArcMap 10.2). The
144 shapefiles were imported on top of the pre-processed SMOS image pixels in the BEAM VISAT and
145 SMOS toolbox. These pixels were further analyzed using Microsoft Excel and Matlab 2016a.
146 Comparisons of the in-situ soil moisture (0–5 cm) and the satellite soil moisture retrievals were performed
147 and are presented in the results below. Evaluation was performed on point by point comparison of the *in-*
148 *situ* and satellite products. The statistical performance measures used were: The Root Mean Square Error
149 (RMSE), Pearson’s Correlation coefficient (R) including slope and intercept, Spearman’s rank correlation
150 coefficient (Rs), the Mean Error (Bias), and the standard deviation (Scatter). Those statistical measures
151 have been used in other previous studies (e.g. (G. Petropoulos et al. 2013; Deng et al. 2019). The analysis
152 was carried out on different land cover types and agreement was evaluated for 7 sites. Similarly,
153 agreement was also evaluated for the 4 seasons, spring (March- May), summer (June–August), autumn
154 (September–November) and winter (December–February), direct point-by-point comparisons were
155 performed at every in-situ station to evaluate the statistical agreement for each threshold. Analysis was
156 performed for each scenario independently for both 2010 and 2011.

157

158

159 **4. Results**

160 **4.1 Europe**

161 **4.1.1 Different land covers Performance comparisons**

162 The first study area is Europe, with three stations. The land cover type mainly covered by AGU represents
163 Shrublands, LJU represents olive orchards and MAU represents croplands. AGU and LJU are used for the
164 evaluations in 2010, and AGU, LJU and MAU are used for the year 2011.

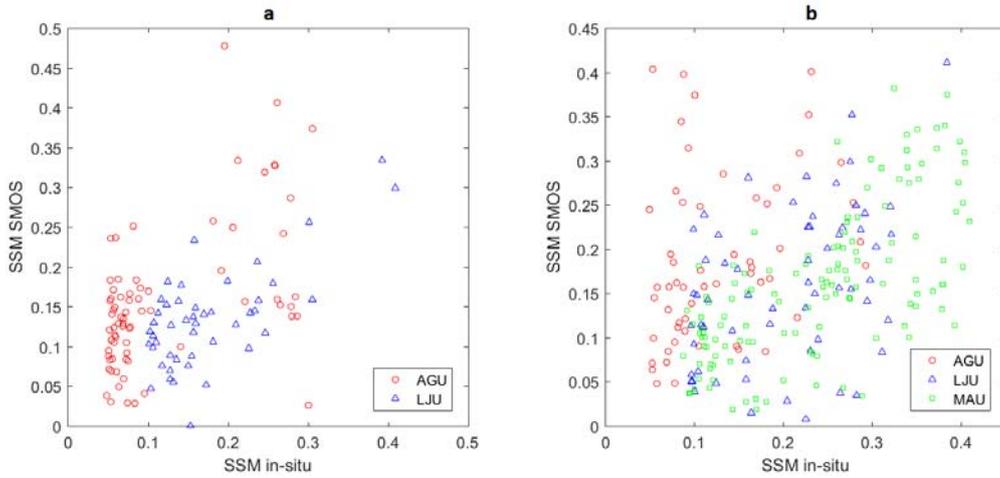
165 Table 1, Figures 1-5 show the evaluation results of SMOS SSM product in Europe for the years 2010 and
166 2011. In addition, considering that the scatter plot at a given significance level can effectively show the
167 general trends of the correlation R between the SMOS predicted SSM with the *in-situ* measurements and
168 outliers of an array, 95% confidence levels was used to intuitively reflect and compare the parameter
169 values shown in Figure 2. Generally, as indicated from the statistical metrics calculated for the case of the
170 comparisons for all sites, a relatively satisfactory agreement between the two compared datasets was
171 reported (RMSE = 0.101 m³/m³, bias = -0.024 m³/ m³, scatter = 0.099 m³/m³ and R= 0.446).

172 Further analysis was conducted to evaluate the product performance over the different land cover types.
173 As can be seen from Table 1, Figure 1 and 2, the correlation coefficient varied from 0.537-0.683 in 2010
174 over AGU& LJU to 0.303, 0.428 & 0.673 in 2011 over AGU, LJU and MAU respectively. Notably, for
175 AGU and MAU in the 2011 the RMSE is larger than 0.1 m³m⁻³ due to the presence of bias whereas the
176 correlation obtained for AGU was low (R = 0.303). On both land covers AGU (Shrubland) and LJU
177 (Olive), SSM product shows a good estimation against the *in-situ* measurements for the year 2010 (figure
178 1a). The SSM product estimation showed lower performance against the *in-situ* measurements for the
179 year 2011 (figure 1b). When the results are combined for all sites for both years, there is indication of
180 bias (-0.024) leading to an underestimation of the predicted SSM. In addition, performance for all sites
181 was better in 2010 than 2011 with overall lower RMSE for 2010 than 2011. Similar findings were
182 reported also for the Scatter. In general, SMOS product behaved similarly in the different land cover types.
183 The SMOS products for LJU (2010) had the best fitting trend with a high R (Figure 2b).

184 **Table 1:** Comparison between Satellite (SMOS) and observed SSM at the validation sites in EU based on
185 land cover type, for 2010 and 2011 as well as all sites (both years). AGU represents Shrublands, LJU
186 represents Olive Orchards and MAU represents Croplands. Units are in m³/m³

Measure	AGU 2010	AGU 2011	LJU 2010	LJU 2011	MAU 2011	All Sites
ME (bias)	0.037	0.063	-0.040	-0.046	-0.080	-0.024
MAE	0.074	0.085	0.054	0.079	0.087	0.079
RMSE	0.092	0.116	0.067	0.099	0.110	0.101
R	0.537	0.303	0.683	0.428	0.673	0.446
Rs	0.447	0.404	0.479	0.397	0.675	0.433
Scatter	0.085	0.099	0.054	0.089	0.075	0.099
Slope	0.566	0.431	0.595	0.478	0.587	0.391
Intercept	0.088	0.135	0.030	0.061	0.019	0.088
N	74	56	46	61	130	367

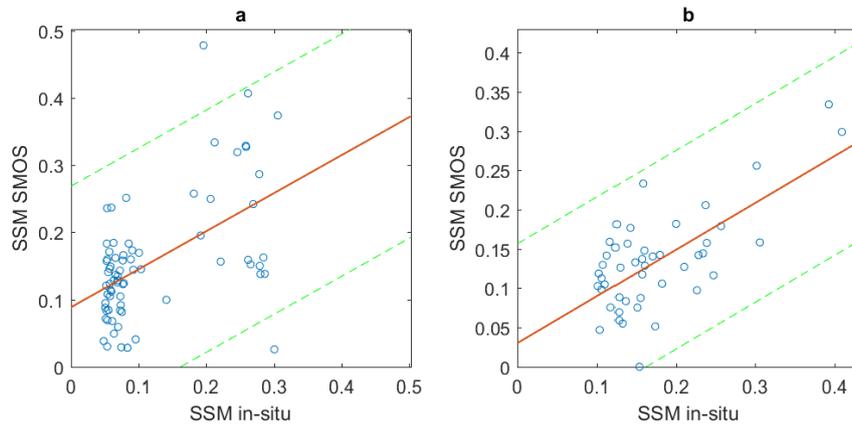
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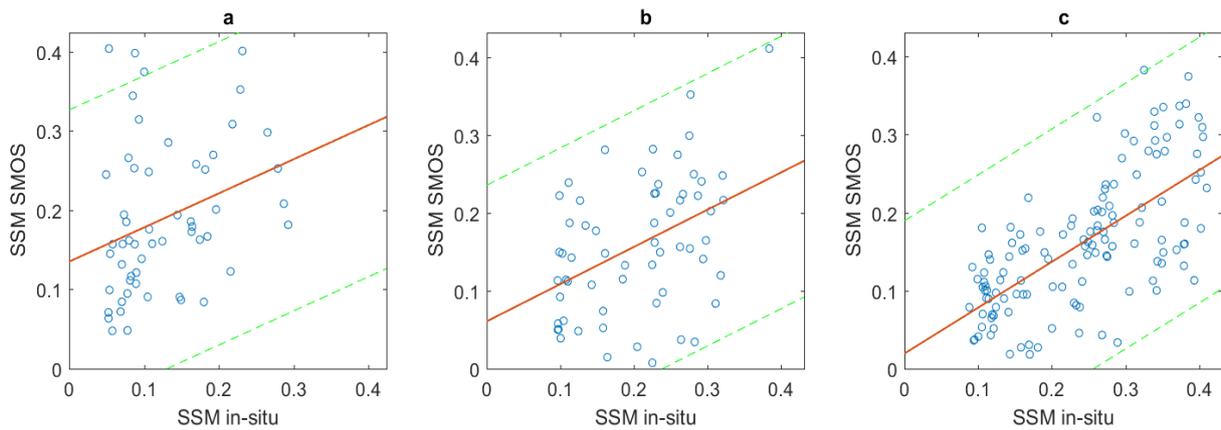
189 **Figure 1:** Agreement between in-situ and predicted SSM from SMOS for the different land cover types in
 190 **EUROPE.** Results are shown for: a) 2010: ES_AGU (red) and ES_LJU (blue). b) 2011: ES_AGU (red),
 191 ES_LJU (blue), FR_MAU (green)

192 i-



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194 ii-



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196 **Figure 2:** Agreement between in-situ and predicted SSM from SMOS for all the different land cover
 197 types in Europe. Results are shown for: i- 2010: a) ES_AGU and b) ES_LJU ii- 2011: a) ES_AGU b)
 198 ES_LJU c) FR_MAU (green)

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200 4.1.2 Temporal Variability

201 To explore the temporal trends between *in-situ* and SMOS product for different seasons during 2010 and
 202 2011, the in-situ measurements (red) and the predicted SSM (blue) over AGU, LJU in 2010, AGU, LJU
 203 and MAU in 2011 are investigated by month, when possible by the data, as shown in Figure 3. The 95%
 204 confidence intervals are shown as green dashed lines in figures 3 and 4. Due to discontinuous data and
 205 small number of data per month, the confidence margins are wide and there are gaps in the data. Thus, it
 206 is not always possible to have results for overestimation or underestimation in a given month with
 207 statistical certainty.

208 In 2010, as shown in Figure 3, the SMOS product overestimated the in-situ observations from September
 209 to November over AGU (Figure 3a) with statistical significance. In 2011, SMOS product aligns with
 210 AGU within the 95% confidence level, except for October and November although data for the previous
 211 months are scarce. Looking at the entirety of AGU though, SMOS tends to overestimate the SSM. The
 212 time series for LJU in 2011 show a greater lack of data and cannot lead to conclusions about
 213 overestimation or underestimation. For Croplands (MAU site, Figure 4c) the data is continuous. SMOS
 214 underestimates the SSM for this site especially from January to April.

215 Table 2 summarizes the comparisons between autumn, winter, spring and summer in 2010 and 2011.
 216 Figures 3- 5 show the agreement between predicted and observed soil moisture for the different seasons
 217 separately for 2010, 2011. Generally, all seasons displayed adequate RMSE (between 0.071 and 0.139)
 218 but a low correlation coefficient. No clear patterns that can be seen between the seasons in 2010 and 2011.
 219 The correlation (R 0.22) in spring 2011, could be associated with the negative bias and the smaller size in
 220 comparison to the other seasons in 2010 and 2011.

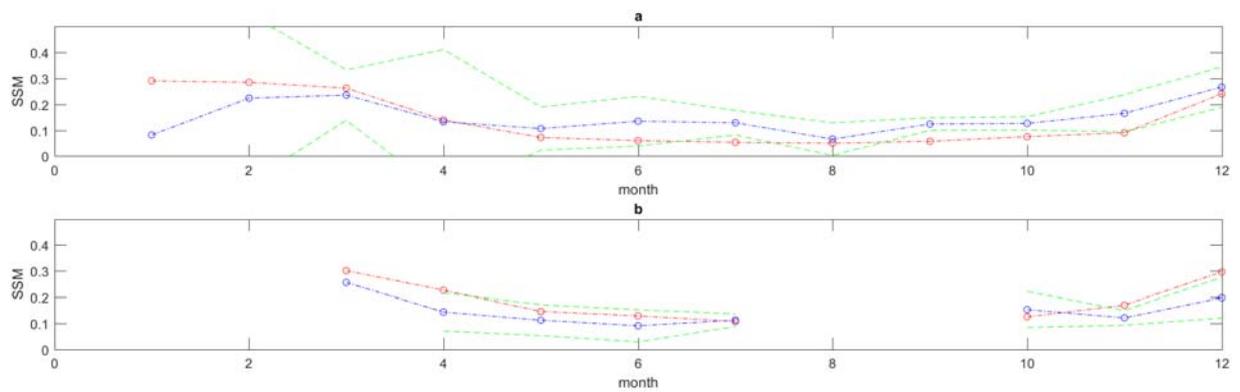
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222 **Table 2:** Comparison per season between Satellite (SMOS) and observed SSM at all validation sites in
 223 EU for 2010 and 2011. Units are in m³/m³

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Measure	Autumn 2010	Winter 2010	Spring 2010	Summer 2010	Autumn 2011	Winter 2011	Spring 2011	Summer 2011
ME (bias)	0.039	-0.051	-0.030	0.031	-0.014	-0.087	-0.063	-0.010
MAE	0.061	0.101	0.067	0.053	0.075	0.099	0.110	0.064
RMSE	0.078	0.116	0.077	0.071	0.093	0.119	0.139	0.089
R	0.305	0.234	0.559	-0.110	0.338	0.432	0.218	0.365
Rs	0.282	0.144	0.649	0.021	0.293	0.404	0.277	0.298
Scatter	0.068	0.107	0.073	0.065	0.093	0.081	0.126	0.089
Slope	0.452	0.443	0.626	-0.185	0.334	0.531	0.236	0.372
Intercept	0.093	0.100	0.037	0.128	0.095	0.053	0.115	0.089
N	44	19	26	31	79	55	50	63

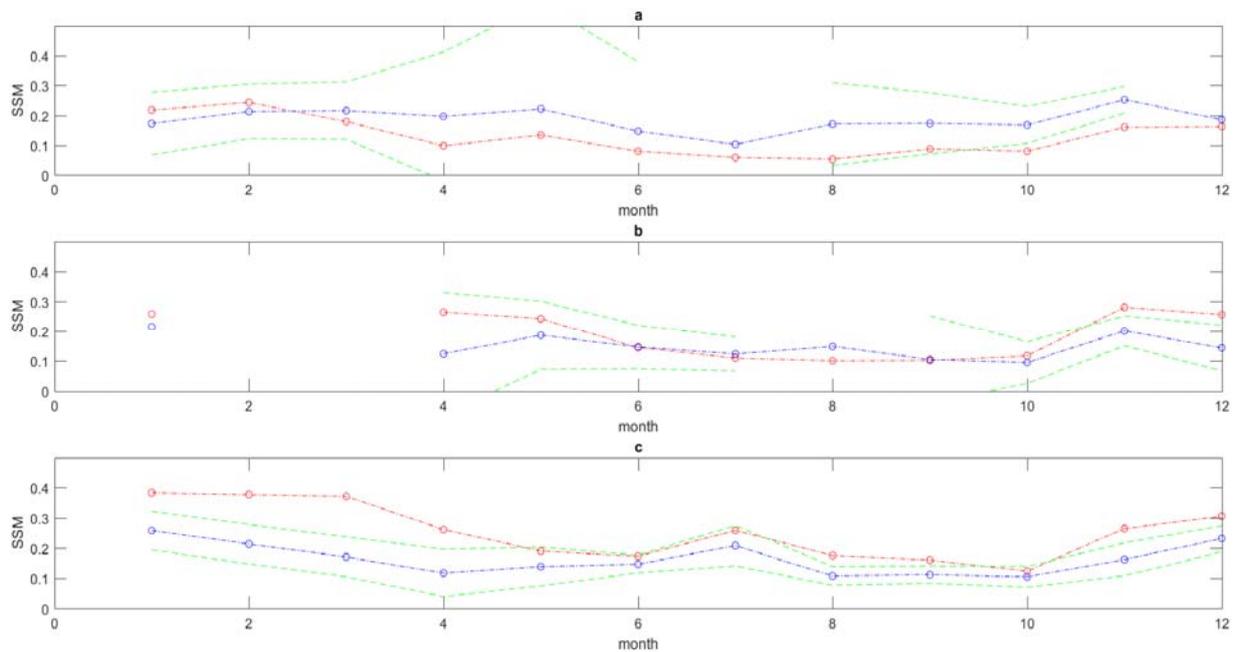
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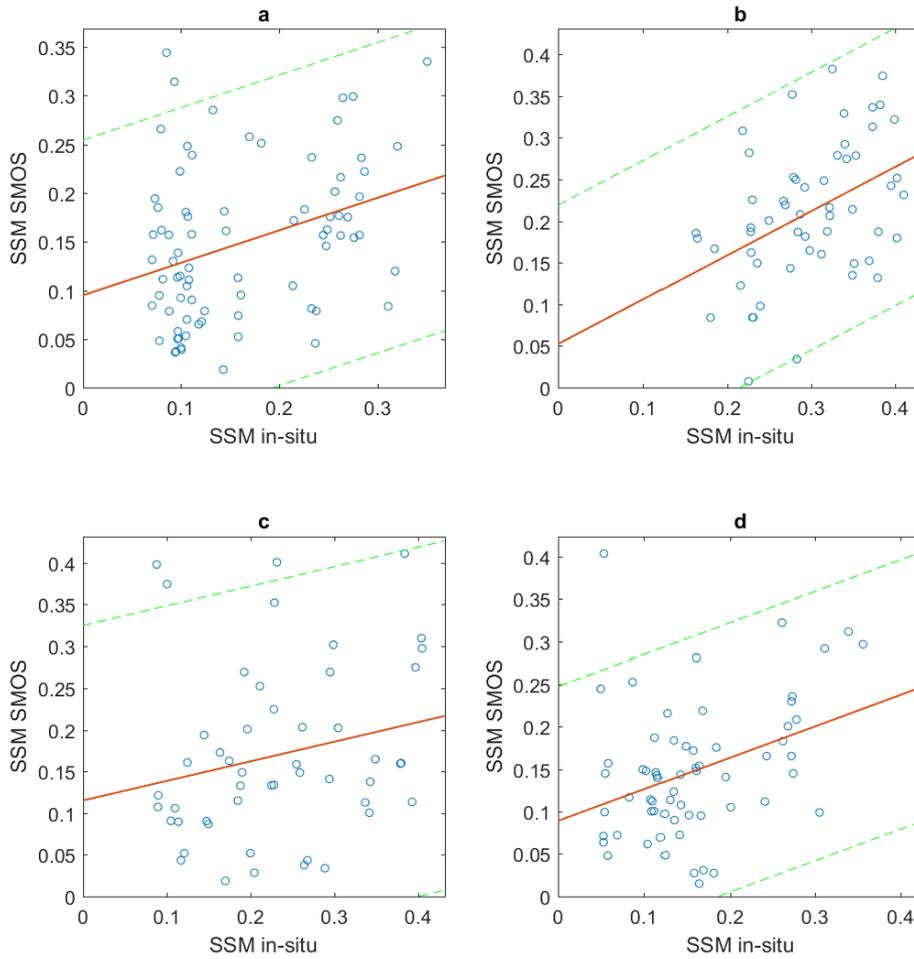
227 **Figure 3:** Agreement between in-situ (red) and predicted SSM (blue) from SMOS for the different land
 228 cover types throughout **2010** in **EUROPE**. Results are shown for: (a) ES_AGU and (b) ES_LJU.

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230 **Figure 4:** Agreement between in-situ (red) and predicted SSM (blue) from SMOS for the different land
 231 cover types throughout **2011** in **EUROPE**. Results are shown for: (a) ES_AGU and (b) ES_LJU and (c)
 232 FR_MAU

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251 **Figure 5:** Agreement between in-situ and predicted SSM from SMOS for the different seasons for all
252 sites together shown here for year EUROPE. In particular, for (a) Autumn, (b) Winter, (c) Spring, (d) for
253 2011

254

255 4.2 USA:

256 4.2.1 Comparisons for different land use/cover types

257 A total of four stations were included in the USA. The characteristic land surface cover types in this area
258 are as follows: ME2 stands for Evergreen Needle-leaf Forest (ENF), TON represents Woody Savannahs
259 (WSA), VAR stands for Grasslands (GRA) and WHS represents Open Shrublands (OSH). Table 3 and
260 Figures 6-9 show the comparison statistics between SMOS product and the in-situ measurements over
261 different land cover types. The correlation coefficient of the predicted and the observed measurements is
262 included in the scatterplots (Figure 6).

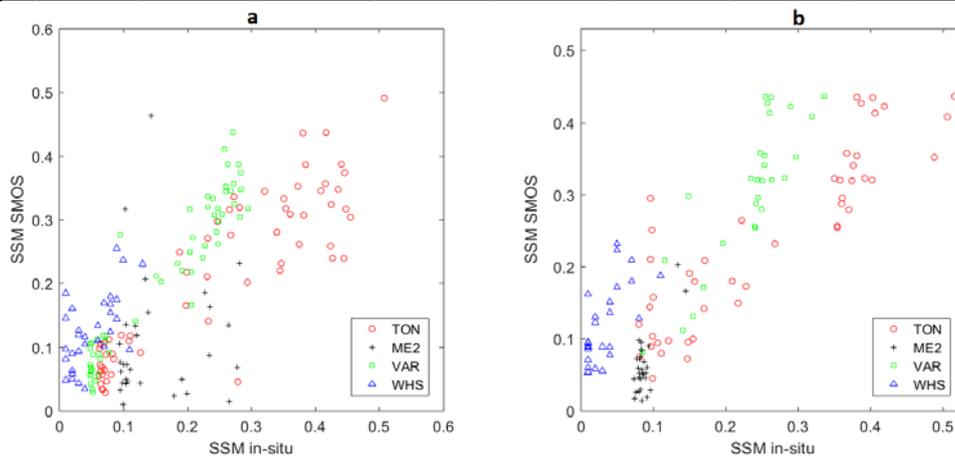
263 Overall, as indicated from the statistical metrics computed for the analysis of the combination of all sites,
264 a very good agreement between the two datasets was indicated (RMSE = 0.080 m³/m³, bias = 0.014 m³/
265 m³, scatter = 0.079 m³/m³ and R= 0.80). In the case of the different Land cover comparison, SMOS

266 product had negative bias over ME2 (ENF) and TON (WSA), while the positive bias over VAR (GRA)
 267 and WHS (OSH) for both 2010 and 2011. In term of correlation coefficient, SMOS product performance
 268 was good for all sites except ME2 (ENF) which has minimum correlation coefficient for $R_{ME2_2010} =$
 269 0.116 reflected by the $RMSE=0.117$, in contrast to the high performance in 2011 for the same site
 270 ($R_{ME2_2011} = 0.759$, $RMSE=0.040$). In addition, SMOS product shows maximum scatter (or standard
 271 deviation) on ME2₂₀₁₀ (Scatter = 0.110), while the scattering is minimum in ME2₂₀₁₁ (Scatter = 0.031).
 272 Moreover, the product has shown small RMSE values (all between 0.040 and 0.117), illustrating
 273 preferable correlation to the in-situ measurements. In addition, the SMOS products for TON, VAR and
 274 WHS showed good data quality in terms of accuracy, stability, and correlation coefficient over the
 275 different land cover types in USA, as seen in the scatter plots of Figure 7. In contrast SMOS product over
 276 Evergreen Needle-forest ME2 could not effectively coincide with the in-situ measurements (Figure 7: (i)-
 277 a and (ii)-a) displaying either very high or very low slope and high intercept.

278

279 **Table 3:** Comparison between Satellite (SMOS) and observed SSM at the validation sites in USA based
 280 on land cover type, for 2010 and 2011 as well as all sites (both years). ME2 stands for Evergreen Needle-
 281 leaf Forest, TON represents Woody Savannahs, VAR stands for Grasslands and WHS represents Open
 282 Shrublands. Units are in m^3/m^3

Measure	ME2 2010	ME2 2011	TON 2010	TON 2011	VAR 2010	VAR 2011	WHS 2010	WHS 2011	All Sites
ME (bias)	-0.045	-0.025	-0.033	-0.016	0.047	0.073	0.074	0.082	0.014
MAE	0.089	0.033	0.054	0.055	0.051	0.077	0.075	0.082	0.062
RMSE	0.117	0.040	0.075	0.068	0.065	0.093	0.086	0.092	0.080
R	0.116	0.759	0.890	0.875	0.947	0.842	0.616	0.682	0.803
Rs	0.232	0.276	0.840	0.849	0.913	0.845	0.528	0.629	0.736
Scatter	0.110	0.031	0.068	0.067	0.046	0.059	0.045	0.042	0.079
Slope	0.176	2.036	0.768	0.758	1.238	1.406	1.014	1.436	0.789
Intercept	0.079	-0.116	0.024	0.046	0.008	-0.022	0.073	0.068	0.050
N	31	29	61	46	58	28	31	29	313



283
 284 **Figure 6:** Agreement between in-situ and predicted SSM from SMOS for all the different land cover
 285 types in USA. Results are shown for US_ME2 (black), US_TON (red), US_VAR (green) and US_WHS
 286 (blue): a) 2010 b) 2011

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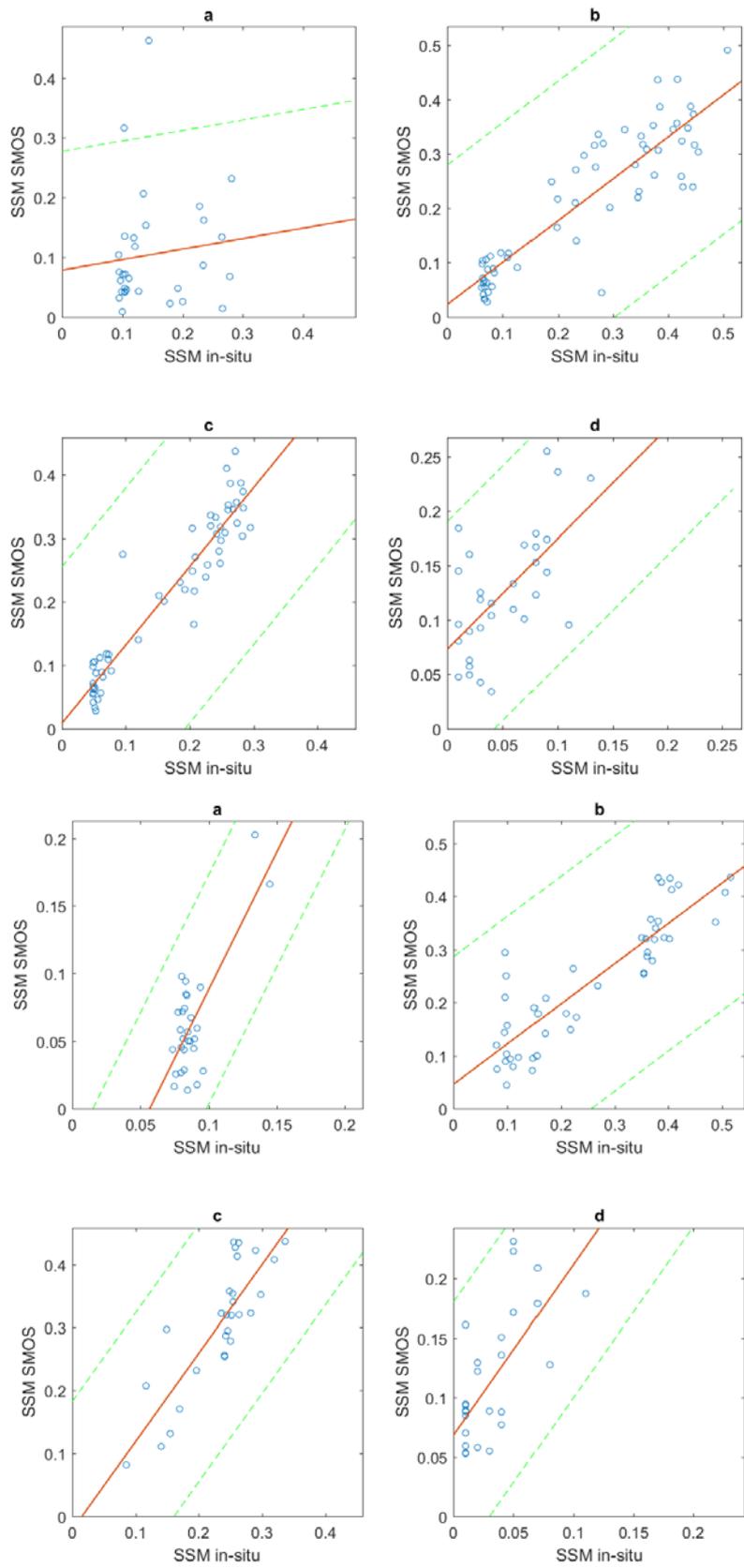
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304 **Figure 7:** Agreement between in-situ and predicted SSM from SMOS for the different land cover types in
305 USA. Results are shown for (a) US_ME2, (b) US_TON, (c) US_VAR and (d) US_WHS. i- 2010 ii- 2011

306

307 4.2.2 Temporal Variability

308 Figure 8 shows the temporal fitting trend between the in-situ measurements (red) and the predicted SSM
309 (blue) over TON (Woody Savannahs), VAR (Grasslands) and WHS (Open Shrublands). For ME2 in 2010
310 and for all sites in 2011, the data were discontinuous with large gaps so the temporal fitting was not
311 included. Even for the remaining sites, in some cases there was a small number of data per month leading
312 to wide confidence margins. Thus, it is not always possible to have results for overestimation or
313 underestimation in a given month with statistical certainty.

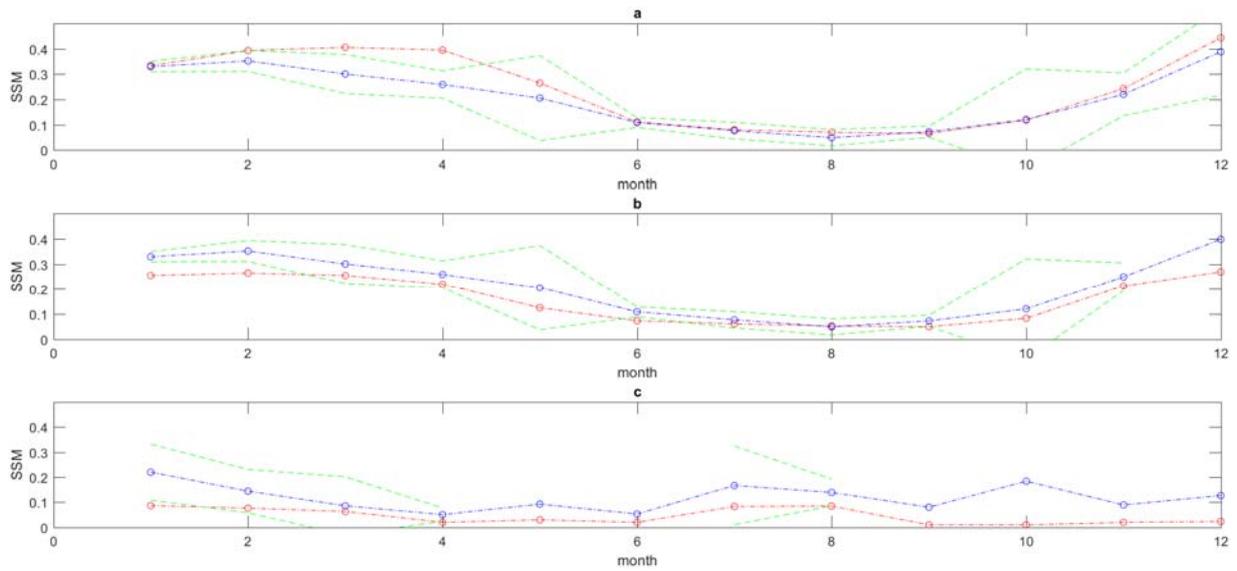
314 SMOS SSM estimations for TON (Woody Savannahs) is in good agreement with *in-situ* SSM from May
315 to December with statistically significant underestimation for March and April. For VAR (Grasslands,
316 figure 8b) and WHS (Shrublands, figure 8c), a slight (and not always statistically significant)
317 overestimation of the SSM by SMOS can be witnessed throughout the year.

318 Table 4 summarizes the comparisons between the seasons and Figure 9 shows the agreement between
319 SSM SMOS and the in-situ measurements for the different seasons separately for 2010 and 2011. In
320 general, the SSM SMOS product has shown low RMSE in all the seasons as shown in Table 4 and Figure
321 9. However, the correlation coefficient R was inferior in autumn 2011 and for both summers and
322 generally good in the other seasons in both years. RMSE has the highest values in winter in both years.
323 Spring of 2011 has the highest Pearson’s coefficient from all sites investigated in all continents and the
324 lowest bias 0.012 m³/m³.

325 **Table 4:** Comparison per season between Satellite (SMOS) and observed SSM at all validation sites in
326 USA for 2010 and 2011. Units are in m³/m³

Measure	Autumn 2010	Winter 2010	Spring 2010	Summer 2010	Autumn 2011	Winter 2011	Spring 2011	Summer 2011
ME (bias)	0.009	0.045	-0.013	-0.011	0.006	0.037	0.012	0.055
MAE	0.050	0.079	0.069	0.057	0.049	0.071	0.067	0.064
RMSE	0.074	0.091	0.086	0.082	0.063	0.084	0.078	0.083
R	0.662	0.827	0.792	0.109	0.318	0.837	0.868	0.403
Rs	0.560	0.679	0.735	0.161	0.208	0.805	0.856	0.324
Scatter	0.075	0.080	0.086	0.082	0.063	0.077	0.079	0.065
Slope	0.804	0.590	0.562	0.098	0.461	0.738	0.752	0.569
Intercept	0.033	0.149	0.089	0.083	0.054	0.096	0.078	0.103
N	53	46	31	51	54	45	20	13

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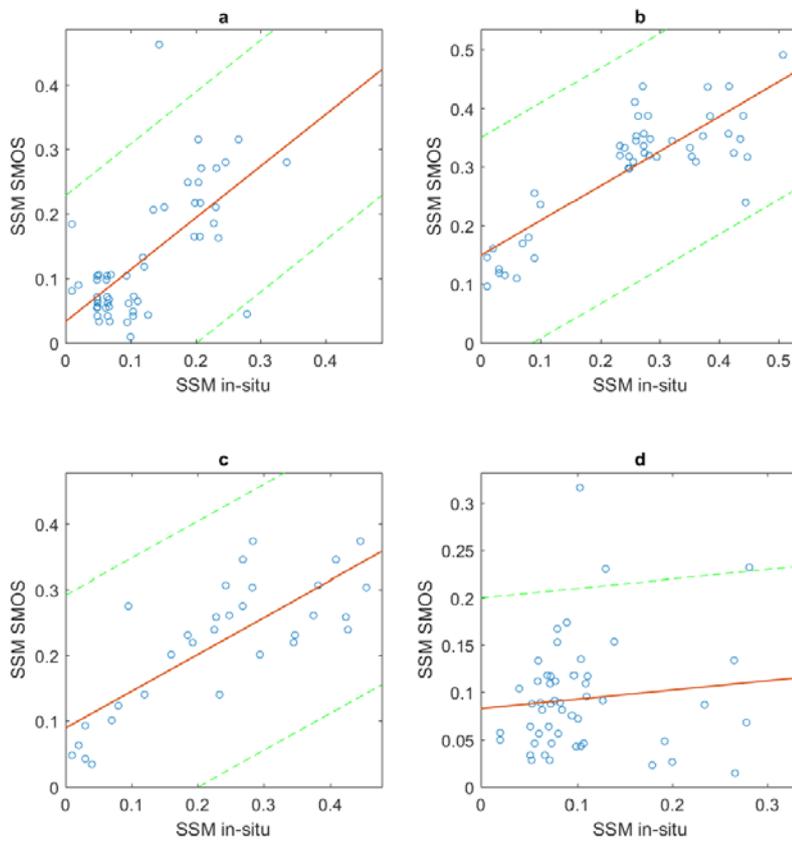
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330 **Figure 8:** Agreement between in-situ (red) and predicted SSM (blue) from SMOS for the different land
 331 cover types throughout 2010 in the USA. Results are shown for: (a) US_TON, (b) US_VAR and (c)
 332 US_WHS

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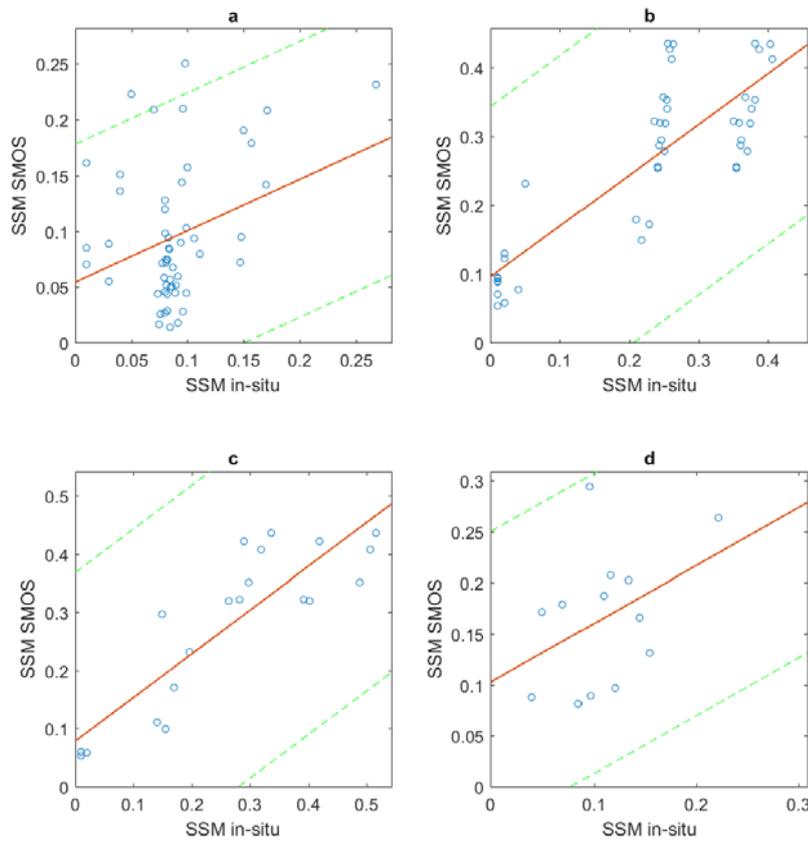
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366 **Figure 9:** Agreement between in-situ and predicted SSM from SMOS for the different seasons for all
367 sites together shown here for year USA. In particular, for (a) Autumn, (b) Winter, (c) Spring, (d) Summer
368 [ii] 2010 and [iii] 2011

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370 5. Discussion

371 In this study, SSM SMOS operational product is evaluated using in-situ measurements in two continental
372 regions based on different land cover types. Three in-situ networks in Europe that included: AGU
373 (Shrublands), LJU (olive orchards) and MAU (croplands). In USA four in-situ networks were used
374 namely: ME2 (Evergreen Needle-leaf Forest), TON (Woody Savannas), VAR (Grasslands) and WHS
375 (Open Shrublands). The performance was evaluated using metrics defined in previous work (Petropoulos
376 et al. 2013). Results are shown in section 4. In this section, an extended discussion is conducted on SMOS
377 SSM product overall performance on different land cover types in order to further improve the algorithm.

378 To summarize, SMOS SSM product is generally applicable in all the selected areas. As shown in Tables
379 1-4, several errors metrics e.g. RMSE, Bias, Scatter, and R showed satisfactory accuracy over selected
380 sites in Europe and USA. In Europe, SMOS SSM has shown reasonable R values, except over the Open
381 Shrubland of AGU₂₀₁₁. This could be due to number of factors such as the retrieved SMOS SSM product
382 observed at a depth of 0-5 cm, whereas in-situ measurement sensors observed at 5 cm. Thus, the strong

383 response to wet and dry period at shallow depth could be a reasonable explanation for discrepancies in
384 agreement (Petropoulos et al. 2013). The SMOS values usually range between (0.001–0.7 m³ m⁻³),
385 although the values generally presented a dry bias, which causes an underestimation. There is not strong
386 evidence suggesting systematic overestimation or underestimation of SSM by SMOS. This result is
387 coincident with some previous studies that have validated SMOS (Petropoulos et al. 2014). Recent
388 studies (Gumuzzio et al. 2016; Cui et al. 2017) have suggested that the error in SMOS could be due to
389 lack of scale representation between SMOS and the in-situ observations of surface temperature, land
390 cover information, soil condition in particular and the RFI.

391 Previous studies focusing on the product comparisons at the annual scale shows that soil moisture
392 estimates are driven to a certain extent by the seasonal cycle (Qin et al. 2013; Petropoulos et al. 2018b). In
393 our study, in several cases of the seasonal cycle investigation there was a slight underestimation of SSM
394 by SMOS. The negative correlation in summer can be explained mainly by lack of spatial sampling
395 between predicted and observed comparison (Al Bitar et al. 2012). Also, it can be partially attributed to
396 lesser fractional vegetation cover than other seasons and/or could be associated with the smaller sample
397 size than other seasons. On the temporal series comparisons, the predicted SMOS SSM product often
398 overestimates slightly the in-situ observations from May to June or August. This could be explained by
399 the presence of dew which is most prominent during summer, spring and autumn, respectively (De Jeu et
400 al. 2005; Du et al. 2012). In winter, SMOS predictions have low accuracy for European sites and perform
401 very well in USA sites. As such, no conclusions about the performance of SMOS in winter can be made
402 from our study. Summer correlation coefficient is generally suboptimal compared to other seasons for
403 both 2010 and 2011 for USA and European sites.

404 In the USA sites SMOS showed good agreement between the two datasets. In term of correlation
405 coefficient, SMOS product performance was good for TON (WSA), VAR (GRA), and WHS (OSH) in
406 2010 and 2011, but underperformed for ME2₂₀₁₀ (ENF). In the case of the different vegetation cover
407 comparison, SMOS product had negative bias over ME2 (ENF) and TON (WSA). The product had
408 positive bias over VAR (GRA) and WHS (OSH) in both years 2010 & 2011. Studies linked SMOS errors
409 to global parameters such as soil texture, RFI, and land cover suggested that globally the forest presence
410 in the radiometer field of view appears to have the great influence on SMOS error up to (56.8%) whereas
411 1.7% of the RFI. The extent of the impact varies among different continents; however, soil texture was
412 highlighted as the main influence over Europe whereas RFI had the greatest influence over Asia.
413 Additionally, a land cover difference as a result of spatial heterogeneity could increase the error in SM
414 within a 0.25°-resolution pixel. Whereas, forest as well dense vegetation could increase the SSM error by
415 negatively affecting microwave penetration (Rotzer et al. 2012; Leroux et al. 2013; Liu et al. 2018)

416 The Correlation coefficient (R) was inferior in summer and good for winter and spring for both 2010 and
417 2011. In both years RMSE has high values in winter. This could be due to the frozen soil. During summer
418 and spring the error could partially be explained by the presence of dew which has a significant effect on
419 passive microwave observations by increasing horizontal brightness, and is most prominent during
420 summer, spring, and autumn, respectively (De Jeu et al. 2005; Du et al. 2012). RFI can be defined as the
421 disturbance that affects an electromagnetic radiation emitted from an external source (Murray 2013). It is
422 a major problem in SMOS SSM retrieval, which decreases the efficiency of retrieved soil moisture
423 (Jackson et al. 1999). These disturbances largely reduce or limit the quality of the data. Hence, signal
424 contamination removal in L-band is an ongoing research challenge in Europe and many other parts of the
425 world (Kerr et al. 2012; Oliva et al. 2012; Daganzo-Eusebio et al. 2013).

426 Overall, the lack of agreement between the predicted and observed SSM for all scenarios examined here
427 can be attributed to a number of factors, such as: (1) the topographic and vegetation properties complexity.
428 It is known that high vegetation density (e.g., taller and/or denser), frozen soils, snow cover, and volume
429 scattering in dry soils are very critical for SSM operational products retrieval accuracy (Brocca et al.
430 2017). With regards to vegetation effects in particular the quality of SSM retrievals could be strongly
431 affected by the vegetation structure and water content. (2) The differences in terms of the SSM sensing
432 depth between the compared datasets. In our study the surface ground measurement used for the
433 evaluation is at 5cm. while the sampling depth of the effective soil moisture of SMOS varies as a function
434 of topography and land cover characteristics (Deng et al. 2019). (3) Differences in spatial observation
435 scale. Since the exact scale of the satellite observation could not be represented by ground observation,
436 the average point-based measurement is used as a “reference”. However, as is also argued in many studies,
437 it is difficult to characterize the spatial soil moisture patterns by using in-situ measurement. It is able only
438 to reproduce the temporal dynamic of soil moisture but not the absolute value (Petropoulos et al. 2015).
439 Sometimes, if in-situ sensors are not dense enough, it causes mismatch in scales and hence poor accuracy
440 in comparison. (4) Errors caused by measurements accuracy of the sensors Land surface factors, such as
441 topography, seasons and land cover types (particularly at the presence of forests) have been pointed out as
442 elements to affect the product accuracy and consistency. In addition to that they affect the quality of the
443 product that can be expected by the final user (Dorigo et al. 2013; Petropoulos et al. 2014).

444 **6. Conclusions – Future work**

445 The quantification of the SMOS SSM product accuracy is crucial for hydrological applications of the
446 product and for the retrieval algorithm refinement. This study explores the performance of the SMOS
447 product. SMOS data were compared to in-situ measurements from FLUXNET validated observational
448 networks over different land cover, seasonal and varied climatic zones. This comparison increases our
449 understanding to the product application at continental expanse. SMOS product is available for a long
450 term period that can be used in modeling of scale-related researches such as land surface and hydrological
451 studies. On the other hand the present evaluation can provide help and feedback for the current retrieval
452 algorithm improvement.

453 The study results showed that direct comparison of SMOS operational product with in-situ observations
454 indicate good performance of the product within these sites in respect of the RMSE. The main findings of
455 the study can be summarized as follows:

- 456 (1) The overall comparison at variety of sites showed generally reasonable agreement between the
457 SMOS product and the in-situ measurement of soil moisture, but at different vegetation cover,
458 some SMOS observations show negative bias. The results were largely comparable to pervious
459 related validation studies.
- 460 (2) The agreement between the in-situ measurements and the product SSM estimations is
461 observed in regard to different vegetation covers, where SMOS product displayed negative
462 bias over ME2 (ENF) and TON (WSA), while the positive bias over VAR (GRA) and WHS
463 (OSH). This conclusion suggests that the vegetation effects must be carefully accounted for
464 consistent SSM estimations. SMOS uses nadir optical depth and different polarization
465 incidence angle to estimate vegetation optical depth. Land cover impacts the variation of soil
466 moisture content because of increasing transpiration losses and rainfall interception.
467 Furthermore, the type of land cover also influences the vegetation attenuation and scattering

468 albedo which can affect the overall soil moisture retrieval. Therefore, analyzing the effect of
469 vegetation on these algorithms would be important.

470 (3) The seasonal periods where the predicted and observed SSM exhibited low correlation
471 coefficient are summer and autumn. This could partially be explained by the presence of dew
472 which is most prominent during summer, spring and autumn. Seasonality is one of the major
473 controls on soil moisture dynamic and its variability can have very important impacts which
474 can influence overall performance of the soil moisture retrieval.

475 The work presented was focused on temperate areas. As the results were promising, work is ongoing in
476 expanding the SMOS SSM, SMAP, ASCAT operational SSM product evaluation over cold and arid
477 regions. The effect of vegetation cover factor that affects the data quality mentioned in the previous
478 paragraphs will be comprehensively considered. In the future, the integration of numerical weather
479 models, meteorological variables, local hydrological models and information on land cover could also be
480 utilized to more accurately analyze the effect of seasonality on soil moisture estimation as already
481 demonstrated in other studies (Srivastava et al. 2013).

482

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494 **Author Contributions**

495 Methodology Development, Data Acquisition, Analysis: all co-authors; Writing: lead by KAKD with
496 contributions also from all other co-authors.

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