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Altimetric Analysis of the Sea-Surface GPS-Reflected Signals

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4 *Abstract*—We describe the development and implementation of 5 a method for extracting altimetric information using the Passive 6 Reflectometry and Interferometry System (PARIS), i.e., from GPS 7 signals after their reflection off the sea surface. We have for-8 malized one idea laid out in the description of a bistatic system 9 for ocean altimetry using the GPS signal, by Hajj and Zuffada 10 (Jet Propulsion Laboratory), and have extended it to real sit-11 uations encountered in PARIS aircraft experiments. Second, we 12 have developed the corresponding algorithms to produce real-time 13 altimetric observables to be used in dedicated digital signal proces-14 sors. Finally, we have applied this method to estimate sea-surface 15 height from one flight experiment in the North Sea off the coast of 16 Norway.

17 *Index Terms*—GPS, microwave ocean altimetry, microwave 18 ocean scatterometry, sea-surface reflections.

19 I. INTRODUCTION

7 ERY LONG Baseline Interferometry (VLBI), Global 20 Navigation Satellite Systems (GNSSs), and radar altime-21 22 try (RA) are based on the measurement of ranges between 23 phase centers of microwave antennas. These measurements are 24 extracted from the cross-correlation functions, or waveforms, 25 between recorded signals with suitable replicas. These mea-26 surements are processed to provide time-space coordinates and 27 other parameters of geophysical interest, together with their 28 associated uncertainties, to different points on the Earth surface. 29 The International VLBI Service [1] and the International GNSS 30 Service [2] have the mission to provide data and products based 31 on VLBI and GNSS observations which are used in many 32 Earth Science applications. Dedicated RA instruments have 33 been placed on space platforms and provide data and products 34 used to measure sea-surface properties like sea level or its 35 roughness. Reference [3] provides a good background on RA 36 and its applications.

A different technique based on VLBI, GPS, and RA con-38 cepts, termed as Passive Reflectometry and Interferometry 39 System (PARIS), or GNSS Reflectometry (GNSS-R), was pro-40 posed in 1993 [4] to provide additional measurements of the

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M. Martín-Neira is with the European Space Research and Technology Centre, ESA, 2200 AG Noordwijk, The Netherlands (e-mail: rius@ieec.uab.es). Digital Object Identifier 10.1109/TGRS.2009.2036721 sea surface to increase the temporal and spatial coverages of RA 41 measurements. Conceptually, PARIS is a bistatic radar altimeter 42 where the transmitters are GNSS satellites. 43

In all the cases, the measured ranges are built with different 44 components: geometric delays due to the relative position of 45 the transmitter and the receiver, offsets between the transmitter 46 and receiver clock readings, ionospheric delays induced by the 47 electronic content of the ionosphere, tropospheric delay pro-48 duced by the refractivity of the troposphere, multipath caused 49 by scattering objects in the vicinity of the receiving antenna, 50 delays inside the transmitter and the receiver, and noise. Any 51 of these components could be considered as signals or as noise, 52 depending on the application and its variability. The effective 53 separation between the different components is based on the 54 assumption that the collected data during an observing period 55 show different signatures for the different effects. Different 56 strategies could be devised to obtain the delay components 57 with sufficient accuracy and precision in different scenarios. 58 Accurate instrumentation and calibration techniques will con- 59 tribute to this separation, at expenses of the cost. Removing 60 the noise effects by the differentiation of observables is another 61 possibility. Reference [5] provides general discussions on these 62 issues. 63

The accepted practices in metrology [6] classify the uncer- 64 tainties into two categories: 65

- type-A uncertainties: those which are evaluated by sta- 66 tistical methods applied to the observables (i.e., least 67 squares fits, root mean square (rms), standard deviations, 68 covariances, etc.);
- type-B uncertainties: those which are evaluated by using 70 additional relevant information (i.e., data provided in 71 calibration reports, model specifications, etc.).

Those practices are relevant in the emerging GNSS-R field, 73 where different scientific and technological disciplines con- 74 verge, and a reliable uncertainty budget should be established. 75 See, for instance, the different usage of the term *accuracy* in 76 [7] and [8]. 77

A particular uncertainty could be considered as type A or B, 78 depending on the technique used. For example, the tropospheric 79 delay could be considered type A in VLBI or GNSS for a fixed 80 station during a long observing session but is type B in RA. Be- 81 cause RA is a nadir-looking instrument, it presents more uncer- 82 tainties of type B than VLBI or GPS instruments, which allow 83 a wider angular diversity. Potentially, GNSS-R could observe 84 reflections far from nadir, but then, the interpretation of the 85 observations in terms of geophysical parameter is less reliable 86 because, in such a case, the models have greater uncertainties. 87

AQ1

- AQ4 AQ5
- AQ6

AQ7

AO8

88 The conservative approach is to reject such observations, as 89 Cox and Munk [9] made in their classical measurements of 90 the Sun glistening on the sea surface, when they limited their 91 measurements of the roughness of the sea surface to those 92 gathered when the sun elevation was above 55°.

93 Schematically, an RA [3] extracts two types of measurements 94 from the waveforms: altimeter ranges (to measure the mean 95 sea level) and parameters describing the sea-surface state. It 96 is assumed in this case that the type-A uncertainties are those 97 produced by the random nature of the recorded signals and the 98 variations of the mean sea level and the sea roughness. The 99 rest are considered as type B. This consideration is forced in 100 this case because the altimeter ranges are completely corre-101 lated with the mean sea level, tropospheric delay, instrumental 102 delays, and the radial component of the RA orbit. A similar 103 approach should be taken when considering the use of GNSS-R 104 as an altimeter.

All altimeter data must be postprocessed to produce accu-105 106 rate surface elevation measurements. This postprocessing is 107 called "retracking" and is required because the leading edge 108 of the return waveform deviates from the one produced by 109 reflection off a perfectly smooth surface, causing an error in 110 the telemetered range measurement. There are many different 111 methods described in the literature that are available for retrack-112 ing altimeter data, mostly based on fitting of a theoretical or 113 empirical model to the data (e.g., [10] and [11]). The Zavorotny 114 and Voronovich [12] model, which might consider geometric, 115 instrumental, and sea-state effects, has been widely used for 116 retracking the GNSS-R waveforms (e.g., [13] and [14]). It 117 either requires a posteriori knowledge of the sea-surface state 118 to correct the altimetric range or the model itself is used to 119 invert both altimetric and scatterometric components from the 120 data. The disadvantages of the latter approach are twofold: 121 computing time and dependence on the model. In the following 122 sections, we will provide a procedure to "retrack" the GNSS-R 123 waveforms formally independently of the model and sea-state 124 information, as well as computationally efficient. This tech-125 nique will separate the altimetric components from the scat-126 terometric components of the range with no need of fitting a 127 model. We will finally apply such a procedure to analyze the 128 data gathered during an aircraft experiment.

129II. GNSS-R: Altimetric and130Scatterometric Observables

A GNSS-R receiver collects, in addition to the direct signals, 132 the GNSS signals after their reflections from the Earth surface. 133 Its main product is the relative delay between the reflected and 134 direct signals. To that end, the receiver produces the *waveforms*, 135 i.e., the cross-correlation of the incoming signals with a well-136 known replica of them. The shape of the reflected waveform 137 is the result of adding incoherently the contribution of the 138 different points of the surface, and its shape can be remarkably 139 different by comparison with the one corresponding to the 140 direct signal, as shown in Fig. 1. The shape of the direct 141 waveform is a triangle in amplitude (squared triangle in power). 142 The half width of the triangle corresponds to the time length 143 of the modulation *chip*. The time corresponding to the peak



Fig. 1. Direct and reflected waveforms. The cross-correlation is computed in two time windows. The direct window follows the peak of the direct signal using a delay lock loop, and the reflected window position lags the direct signal with a delay computed with *a priori* information based on the real-time navigation information.

of the direct waveform is taken as the apparent arrival time of 144 the direct signal t_D . In the reflected signal, we will distinguish 145 the time t_{DER} corresponding to the maximum of the derivative 146 of the waveform and the time of the maximum waveform 147 amplitude t_{MAX} . As it will be justified in the next section, 148 the delay corresponding to the maximum of the waveform 149 derivative is a biased estimator of the *specular* delay t_S . In the 150 case of a reflection in a smooth surface, the reflected waveform 151 is also a triangle, and $t_S = t_{\text{DER}} = t_{\text{MAX}}$. The following dis-152 cussion assumes that $t_S - t_{\text{MAX}}$ is significantly larger than the 153 waveform sampling interval, which is a condition fulfilled when 154 the surface is sufficiently rough.

From these quantities, we define the observed specular delay 156

$$\Delta \tau_{\rm spec}^{\rm obs} \equiv t_{\rm DER} - t_D \tag{1}$$

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and the observed scatterometric delay

$$\Delta \tau_{\rm scatt}^{\rm obs} \equiv t_{\rm MAX} - t_{\rm DER} \tag{2}$$

which contain information related with the position and the sea 158 state, respectively.

The observed specular delay $\Delta \tau_{\rm spec}^{\rm obs}$ is obtained using 160

$$\Delta \tau_{\rm spec}^{\rm obs} \equiv (t_{\rm DER} - t_{\rm RW}) + (t_{\rm RW} - t_{\rm DW}) - (t_D - t_{\rm DW})$$
(3)

where $t_{\rm DW}$ is the delay corresponding to the origin of the direct 161 window, as measured by the navigation receiver, and $t_{\rm RW}$ is the 162 corresponding value computed by the receiver for the origin of 163 the reflection window.

Reference [15] presents a theoretical description of a bistatic 167 system for ocean altimetry using a GPS signal. It provides an 168 intuitive framework for the discussion of the capabilities of the 169 PARIS concept from the space. We have taken that paper as 170 a conceptual departure point. Those authors indicate that the 171 delay corresponding to the specular reflection could be obtained 172 by looking for the inflection point of the waveform with positive 173 slope. They support this claim after considering the case of a 174 175 waveform built as the limit of an infinite incoherent sum of 176 equal triangles shifted in delay. A search in the open literature 177 has shown that these ideas have not been developed further by 178 the involved community.

Now, we explore some properties of the derivative of the waveforms modeled using the radar equation, as formulated 181 in [12]

$$w(\tau) = T_i^2 \cdot \iint \frac{G \cdot S \cdot \Lambda^2 \left[\tau - (R_0 + R)/c\right]}{4\pi R_0^2 R^2} \cdot \sigma_0 \cdot d^2 \rho$$
(4)

182 where T_i is the coherent integration time, ρ is the integration 183 variable, $G = G(\rho)$ is the power gain of the receiving antenna, 184 $S(f(\rho))$ is the sinc-exponential function to account for different 185 frequency f Doppler effects onto different ρ positions, and the 186 triangle function Λ is defined as

$$\Lambda(\tau) = \begin{cases} 1 - |\tau|/\tau_c, & \text{if } |\tau| < \tau_c \\ 0, & \text{otherwise} \end{cases}$$
(5)

187 where τ_c is the chip length, the quantity $\sigma_0(\rho)$ is the normalized 188 bistatic radar cross section of the sea surface, and R_0 and R are 189 the distances from the transmitter and the receiver to point ρ . 190 For the rest of this derivation (Section III), τ is expressed in 191 units of τ_c , whereas the rest of this paper will give the delay in 192 units of length (in meters).

193 Equation (4) can be written [16] as the convolution product

$$w(\tau) = p(\tau) * \Lambda^2(\tau) \tag{6}$$

194 where the power per unit delay $p(\tau)$ represents the contribution 195 to the waveform of the points with a delay τ .

196 If the sea surface was a smooth surface, the shape of the 197 reflected waveform $w(\tau)$ would be the same as the direct 198 waveform but shifted by a delay corresponding to the difference 199 between the delays of the reflected and direct signals. When 200 the surface is not smooth, off-specular (or scattered) reflections 201 are generated, and the reflected power $p(\tau)$ spreads for longer 202 times, starting at the specular delay. Setting the specular delay 203 as $\tau_{\rm spec} = 0$, it reads

$$p(\tau) = 0, \qquad \tau < 0$$

> 0, \quad \tau \ge 0 (7)

204 and its derivative could be expressed in these two alternative 205 forms [17]

$$w'(\tau) = p'(\tau) * \Lambda^{2}(\tau)$$
$$= p(\tau) * \Lambda^{2'}(\tau).$$
(8)

206 The derivative of the squared triangle function is (see 207 also Fig. 2)

$$\begin{aligned} \Lambda^{2'}(\tau) &= 2\Lambda(\tau)\Lambda'(\tau) = 2(\tau+1), & -1 \le \tau \le 0 \\ &= 2(\tau-1), & 0 < \tau \le 1 \\ &= 0, & |\tau| > 1. \end{aligned}$$



Fig. 2. Panel 1 represents the reflected power as a function of the delay relative to the specular delay, as defined in (7). Panel 2 represents the function $\Lambda^2(\tau)$ defined by (5) as a thin line and its filtered version with a thick line. The following panels use the same convention to represent the unfiltered and filtered versions. The waveform $w(\tau)$ is obtained using (6), and it is represented in Panel 3. Note that both versions are very similar in shape and in the delay corresponding to the maximum power. Panel 4 represents the derivative of the waveform $w'(\tau)$. Now, the position of the derivative depends on the filtered signal. The same results are obtained using the derivative of $\Lambda^2(\tau)$, as indicated in the path which includes Panel 5.

The derivative of the waveform $w(\tau)(8)$ will present a peak 208 in the specular point if its second derivative switches sign at 209 that moment, from positive to negative. The second derivative 210 is expressed as 211

$$w''(\tau) = p(\tau) * \Lambda^{2''}(\tau) \tag{9}$$

$$= \int_{-\infty}^{\infty} p(\tilde{\tau}) \Lambda^{2''}(\tau - \tilde{\tau}) \,\mathrm{d}\tilde{\tau}.$$
 (10)

However, $p(\tilde{\tau})$ is defined in $\tilde{\tau} > 0$ (no power earlier than 212 specular ray path) 213

$$w''(\tau) = \int_{0}^{\infty} p(\tilde{\tau}) \Lambda^{2''}(\tau - \tilde{\tau}) \,\mathrm{d}\tilde{\tau}.$$
 (11)

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 $\Lambda^{2^{\prime\prime}}(x)$ reads

$$\begin{aligned} & e^{2n}(x) = 2, & -1 \le x < 0 \\ & = 2, & 0 < x \le 1 \\ & = \lim_{\epsilon \to 0} \frac{-4}{2\epsilon}, & -\epsilon < x < \epsilon \\ & = 0, & |x| > 1. \end{aligned}$$

 $w''(\tau)$ in (11) is zero when $\tau < -1$. For $\tau \ge -1$, it reads 215

$$w''(\tau) = \lim_{\epsilon \to 0} \left[\int_{\tau-1}^{\tau-\epsilon} 2p(\tilde{\tau}) \,\mathrm{d}\tilde{\tau} + \int_{\tau+\epsilon}^{\tau+1} 2p(\tilde{\tau}) \,\mathrm{d}\tilde{\tau} + \int_{\tau-\epsilon}^{\tau+\epsilon} \frac{-4}{2\epsilon} p(\tilde{\tau}) \,\mathrm{d}\tilde{\tau} \right].$$
(12)

216 The last term is

$$\lim_{\epsilon \to 0} \frac{-4}{2\epsilon} \int_{\tau-\epsilon}^{\tau+\epsilon} p(\tilde{\tau}) \,\mathrm{d}\tilde{\tau} = \lim_{\epsilon \to 0} \frac{-4}{2\epsilon} 2\epsilon p(\tau) = 4p(\tau) \tag{13}$$

217 and the first and second terms can be expressed as

$$\lim_{\epsilon \to 0} 2 \left[\int_{\tau-1}^{\tau-\epsilon} p(\tilde{\tau}) \,\mathrm{d}\tilde{\tau} + \int_{\tau+\epsilon}^{\tau+1} p(\tilde{\tau}) \,\mathrm{d}\tilde{\tau} \right] \\ = 2 \int_{\tau-1}^{\tau+1} p(\tilde{\tau}) \,\mathrm{d}\tilde{\tau} \simeq 4 \langle p(\tau) \rangle_{[\tau-1,\tau+1]}$$
(14)

218 with $\langle p(\tau) \rangle_{[\tau-1,\tau+1]}$ being the average value of $p(\tau)$ in the 219 range $[\tau - 1, \tau + 1]$. Therefore, combining (13) and (14) in 220 (12), the second derivative of the waveform becomes

$$w''(\tau) = 2 \int_{\tau-1}^{\tau+1} p(\tilde{\tau}) \,\mathrm{d}\tilde{\tau} - 4p(\tau) \simeq 4 \left(\langle p(\tau) \rangle_{[\tau-1,\tau+1]} - p(\tau) \right)$$
(15)

221 i.e., proportional to the difference between the power averaged 222 along two code chip lengths and its instantaneous value.

As stated in (7), there is a positive discontinuity at the delay 224 of the specular point: $p(\tau)$ is zero for $\tau < 0$ and positive for 225 $\tau \ge 0$. If we also assume that $p(\tau)$ is a continuous function 226 monotonically decreasing, then $w''(\tau)$ in (15) will switch from 227 positive to negative sign at the specular point (maximal of a 228 function is greater than the averaged value). Hence

$$t_S = t_{\text{DER}}.\tag{16}$$

Fig. 2 shows graphically the previous discussion. The con-230 volution product of $p(\tau)$, which contains information on the 231 sea state, with $\Lambda^2(\tau)$, which represents the impulse response 232 of the instrument, produces a waveform $w(\tau)$, from which the 233 derivative $w(\tau)'$ is extracted. The thin line corresponds to the 234 ideal case of infinite bandwidth, discussed previously. When 235 we consider a band-limited receiver, the impulse response is 236 filtered, and the results are somewhat different, as indicated by 237 the thick line. Note that, in particular, this filtering introduces a 238 bias in the position of the specular point.

The previous result corresponds to an idealized situation, 240 represented by (6), which is a simplification of (4). A more 241 complete model should include, in addition to the complete 242 radar equation, the consideration of a finite sampling interval of 243 the waveforms in the actual GNSS-R receivers and the presence 244 of different sources of noise: thermal, speckle, and processing 245 noises. With a more complete model, the result equivalent to 246 (16) is

$$t_{\rm DER} = t_S + \Delta \epsilon_{\rm spec} \tag{17}$$

247 where $\Delta \epsilon_{\text{spec}}$ is a correction term which appears when we 248 consider a band-limited version of Λ^2 and a finite sampling rate. 249 This term is zero in the limit case of very large bandwidth and very small sampling rate, in agreement with the analytical dis- 250 cussion; otherwise, it is different from zero. We have estimated 251 this term with a model based on (4) with the actual sampling 252 rate (20 MHz) and bandwidth (10 MHz), and we have found 253 that this term is nearly a bias with a weak dependence on the 254 mean square slope (mss), as indicated in the last column of 255 Table II. 256

To obtain the result $t_{\text{DER}} = t_S$ with our numerical model, 257 it is required that the separation between t_{MAX} and t_{DER} 258 be large, when measured in the waveform sampling intervals. 259 In the limit case when the surface is flat, we do not have 260 independent information: both delays coincide. If the separation 261 is only a few times the sampling interval of the waveforms, 262 there will be a correlation between both estimates, produced 263 by the processing algorithms used to extract the information. 264 Because the sampling rate is determined by the bandwidth of 265 the signal and the instrument, better results are expected by 266 using instruments and signals with wider bandwidth and higher 267 sampling rates. 268

The deterministic specular delay model could be for- 270 mulated as 271

$$\Delta \tau_{\rm spec} \equiv \Delta \epsilon_{\rm spec} + (t_S - t_D)$$

= $\Delta \epsilon_{\rm spec} + \Delta \tau_{\rm geo} + \Delta \tau_{\rm ion} + \Delta \tau_{\rm trop} + \Delta \tau_{\rm ins, spec}$ (18)

where the term $\Delta \epsilon_{\rm spec}$ has been introduced in (17), $\Delta \tau_{\rm geo}$ is the 272 difference between the geometrical distances of the paths fol- 273 lowed by the direct and reflected signals between the transmitter 274 and receiver phase centers, $\Delta \tau_{\rm ion}$ is the difference between the 275 reflected and direct signals due to the ionosphere, $\Delta \tau_{\rm trop}$ is 276 the difference between the reflected and direct signals due to 277 the troposphere, and $\Delta \tau_{\rm ins,spec}$ is the specular instrumental de- 278 lay which accounts for the biases induced by the differences be- 279 tween the instrumental delays experienced by the reflected and 280 direct signals, and the differences in the extraction processes. 281

The scatterometric delay model is formulated as 282

$$\Delta \tau_{\rm scatt} = (t_{\rm MAX} - t_{\rm DER}) = \Delta \epsilon_{\rm scatt} + \Delta \tau_{\rm ins, scatt}$$
(19)

where the term $\Delta \epsilon_{\text{scatt}}$ is the delay induced by the sea state 283 and $\Delta \tau_{\text{ins,scatt}}$ is a possible bias term produced by the different 284 process followed to extract the quantities t_{MAX} and t_{DER} . 285

Speckle noise is reduced by the incoherent integration of 286 waveforms, and the thermal noise and processing inaccuracies 287 are expected to be uncorrelated between samples. Table I lists 288 the different sources of uncertainties which we have and its 289 classification as type A or B. The uncertainties associated with 290 the waveform model have two parts: the *a priori* value of the 291 sea-surface state, parameterized as mss, which is considered 292 as type B, and its variation Δmss to be estimated, which is 293 considered as type A. This list is similar to the one correspond-294 ing to RA, with the addition of two terms, namely, multipath 295 and carrier-smoothed pseudoranges, which are relevant to the 296 instrument used in the experiment to be discussed later. The 297

Term	Source of uncertainties	Туре	e Comments	
$\Delta \epsilon_{spec}$	Waveform model	A/B	See Section VII	
$\Delta \tau_{geo}$	Trajectory	В	Quoted by the provider	
	Sea level	А	As in RA. See [3]	
$\Delta \tau_{trop}$	Wet delay	В	As in RA. See [3]	
	Dry delay	В	Accurate model	
$\Delta \tau_{ion}$	Ionosphere	В	Cancels	
$\Delta \tau_{ins}$	Instrument	В	Cables	
	Smoothed pseudorange	В	See [5]	
	Multipath	В	See [5]	
$\Delta \epsilon_{scatt}$	Waveform model	A/B	See Section VII	

TABLE I GNSS-R UNCERTAINTIES CLASSIFIED AS TYPE A OR/AND B

298 multipath is inherent to the use of low-gain antennas. For the 299 direct signal, we have used "carrier-smoothed pseudoranges," 300 which have reduced multipath, but they are affected by the 301 changes in the ionosphere, as discussed in [5].

V. LINEARIZED MODEL

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303 Using the observables extracted from the integrated wave-304 forms $\Delta\tau_{\rm spec}^{\rm obs}$ and $\Delta\tau_{\rm scatt}^{\rm obs}$ and the corresponding modeled 305 values, we form the prefit residuals $\delta \tau_{\rm spec}$ and $\delta \tau_{\rm scatt}$ as

$$\delta \tau_{\rm spec} = \Delta \tau_{\rm spec}^{\rm obs} - \Delta \tau_{\rm spec}^{\rm mod} \tag{20}$$

$$\delta \tau_{\rm scatt} = \Delta \tau_{\rm scatt}^{\rm obs} - \Delta \tau_{\rm scatt}^{\rm mod}.$$
 (21)

We assume that our *a priori* information has the following 306 307 deficiencies:

308 1) the variation of the geoidal height of the specular point with respect to the *a priori* value: Δh ; 309

- 2) a correction term to the *a priori* model of the mss: Δmss ; 310
- 3) two bias terms b_{spec} and b_{scatt} , which include all the 311 312 unmodeled effects.

These assumptions allow one to parameterize the two prefit 313 314 residuals as linear functions of four unknowns

$$\delta \tau_{\rm spec} = \frac{\partial \tau_{\rm spec}^{\rm mod}}{\partial h} \cdot \Delta h + \frac{\partial \tau_{\rm spec}^{\rm mod}}{\partial mss} \cdot \Delta mss + b_{\rm spec}$$
(22)

$$\delta \tau_{\text{scatt}} = \frac{\partial \tau_{\text{scatt}}^{\text{mod}}}{\partial h} \cdot \Delta h + \frac{\partial \tau_{\text{scatt}}^{\text{mod}}}{\partial mss} \cdot \Delta mss + b_{\text{scatt}}.$$
 (23)

Observing different GNSS satellites at different epochs, we 315 316 will have a series of delays $\delta \tau_{\text{spec}}(sat, t)$ and $\delta \tau_{\text{scatt}}(sat, t)$.

As it has been pointed out in the introduction, the selection of 317 318 observations near nadir is required for ocean altimetry because 319 the models are more accurate. However, in this case, we have $(\partial \tau_{
m spec}^{
m mod}/\partial h) pprox -2$, nearly constant, and the estimates of Δh 320 will be highly correlated with $b_{\rm spec}$. 321

In VLBI or GNSS positioning, the bias uncertainties could 322 323 be reduced when the receiver is placed in a fixed position 324 for a long period of time, because in such a case, there is an 325 opportunity of processing observations with a wide range of el-326 evations. Even there, it is recommended to have instrumentation 327 with predictable behavior to reduce the uncertainty. Reference 328 Sec. 6.1.3 [5], entitled Price of an Inexpensive Receiver Clock, 329 discusses these issues. An alternative to the bias estimation



Fig. 3. Distribution of the specular points for the events selected. There are traces for three different satellites: PRN22, PRN19, and PRN3. The aircraft moved from A to B, at roughly 3000-m altitude with an approximate horizontal velocity of 75 m/s, during 1900 s.

is to use the differences between observations from a single 330 satellite or different satellites. This approach, used successfully 331 in conventional space geodesy, will fail in our case because the 332 requirement of using high-elevation observations will "dilute 333 the precision." As we will show in Section VII, we will analyze 334 a GNSS-R experiment, and we will show the difficulties of 335 applying space geodetic techniques to real experiments. 336

VI. EXPERIMENT DESCRIPTION 337

To test the suitability of the use of the waveform derivative 338 to extract altimetric information, we have analyzed data taken 339 with the GOLD RTR system described in [18], during an 340 airborne CoSMOS Ocean Salinity Campaign of the European 341 Space Agency (ESA). It took place on April 15, 2006, in 342 front of the southwestern part of the Norwegian coast. The 343 aircraft was moving with an approximate velocity of 75 m/s at 344 roughly 3000-m altitude. General details of the campaign and 345 the scatterometric results have been described in [19].

We have analyzed the data subset obtained while the aircraft 347 followed a straight path and only from GPS satellites with ele- 348 vation higher than 40°, i.e., GPS satellites with pseudorandom 349 noises (PRNs) 3, 19, and 22. The data set starts at the second 350 AQ10 of the day 31780 and for about 1900 s. The satellite with the 351 highest elevation, PRN 19, was observed continuously during 352 this period. The other two were observed only intermittently, 353 alternating 2-min periods of data with 1-min gaps. This was so 354 because the correlators were scheduled to gather data required 355 by the primary experiment, with sea-surface height as a test 356 of opportunity. The altimetry analysis has used only the civil 357 coarse/acquisition (C/A) code of the GPS L1 signals. 358

The area of the experiment is shown in Fig. 3. The different 359 traces correspond to the specular points for the events selected. 360

The ellipsoidal altitude of the aircraft as a function of longi- 361 tude is shown in Fig. 4. 362

The *a priori* ellipsoidal height of the specular points as a 363 function of longitude is shown in Fig. 4. 364

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Fig. 4. Upper panel describes the elevation of each satellite as can be seen from the specular point. Note that only satellite PRN19 meets the 55° rule chosen in the Cox and Munk experiment [9]. The middle panel gives the ellipsoidal height of the geoid at the specular point as a function of longitude. The different curves correspond to the different satellites, and the lower panel presents the estimate of the ellipsoidal height of the aircraft during the experiment as a function of longitude. Each point represents 1-s measurement. The variations are on the order of meters, which prevented the accurate estimate of this quantity. Values based on the precise trajectory are computed by the Institut de Geomatica, Barcelona.

To geolocate the integrated waveforms, we also collected the following information:

- 367 1) *a priori* geoid grid in the WGS84 reference frame
 368 (see [20]);
- 2) precise platform trajectory (computed by the Institut de Geomatica, Barcelona, using a precise point positioning strategy);
- 372 3) platform orientation in the WGS84 reference frame, as373 measured with the aircraft inertial measurement unit;
- 4) the vector joining the up-looking and down-lookingantenna phase centers in the body frame;
- 376 5) antenna phase patterns in the body frame;
- 6) *a priori* sea state (mss_0) to linearize the forward model, taken from [19].

379 Fig. 5 shows, for the initial and final epochs when the 380 three GPSs were detected simultaneously, the first pulse-limited 381 footprint (isodelay line 1 C/A code chip away from the specular 382 delay), the antenna footprint (isogain line 3 dB down from 383 the maximum gain), and the glistening zone (iso- σ_o line 3 dB 384 down from the maximum bistatic radar cross section). Clearly, 385 the waveforms from PRN03 must be significantly affected by 386 the antenna pattern. It is also noted that the glistening area is 387 smaller than the pulse-limited footprint.



Fig. 5. For the initial and final epochs when the three GPSs were detected simultaneously, we include (discontinuous ellipses) the ISO range corresponding to 300 m for each satellite, (thin circle) the -3-dB gain line corresponding to the down-looking antenna, and (continuous ellipse) the glistening surface defined as the region where the normalized bistatic cross section of the sea surface reduces 3 dB.

TABLE II MAIN CHARACTERISTICS OF THE REFLECTION EVENTS FOR THE THREE SATELLITES SELECTED: PRN19, PRN22, AND PRN3. THE TYPICAL MSS DURING THE PART OF THE FLIGHT ANALYZED HERE WAS 0.020 (FROM[19]), WHICH WE HAVE USED AS THE A PRIORI VALUE IN OUR MODEL COMPUTATIONS. SoDIS FOR SECOND OF THE DAY

PRN	SoD	sin(e)	Az (deg)	$\frac{\partial \tau_{scatt}^{mod}}{\partial mss}(m)$	$\frac{\partial \tau_{spec}^{mod}}{\partial mss}(m)$
19	31749	0.967	237.8	2167	27
19	32649	0.968	210.1	2162	27
19	33649	0.948	185.1	2168	27
22	31749	0.827	89.5	2115	26
22	32649	0.796	78.6	2092	26
22	33649	0.737	68.9	2079	25
3	31749	0.769	156.4	2084	26
3	32550	0.701	154.8	2033	24
3	33649	0.594	153.7	1955	23

Table II summarizes the values of the parameters that are rel-888evant in the description of the reflection events at the beginning,899middle, and end of the chosen period.390



Fig. 6. (Top) Example of a 1-s integrated waveform and (bottom) its derivative collected during the experimental campaign and used in the analysis.

391 VII. DATA ANALYSIS

The aim of the data analysis is to confirm that the techniques
described in this paper perform at the expected uncertainty
level, limited in this case by the following: 1) lack of calibration
to type-B uncertainties and 2) the use of the C/A code of the
GPS signals.

397 The selected GOLD_RTR data set consisted of $3.8 \times 10^{+6}$ 398 1-ms complex waveforms. These waveforms were integrated 399 incoherently during 1-s intervals. This produced 1145 1-s in-400 tegrated waveforms for PRN 3, 1751 for PRN 19, and 1086 401 for PRN 22. Fig. 6 shows a 1-s integrated waveform and its 402 derivative. As discussed in [7] and [21], an estimate of the 403 expected uncertainty in the measurement σ_{τ} on the delay is 404 related to the integrated waveforms $w(\tau)$ and its derivative 405 $w'(\tau)$, for high SNR, through the following:

$$\sigma_{\tau} = \frac{w(\tau)}{w'(\tau)\sqrt{m}}$$

406 where m is the number of waveforms integrated. After the 407 integration of 1000 1-ms waveforms, we have σ_{τ} on the order 408 of 4 m.

409 The delay resolution of the waveforms was 1/20 of a mi-410 crosecond, or 14.99 m, which was too large compared with σ_{τ} : 411 The quantization noise was larger than the measurement noise. 412 To solve this issue, we have interpolated the waveforms using 413 Fourier transform algorithms to have an apparent sampling rate 414 that is eight times faster. An alternative procedure using finite 415 impulse response filters has been used in [13] to address a 416 similar problem.

417 From the integrated waveforms $w(\tau; t, sat)$, and its 418 derivatives $w'(\tau; t, sat)$, we have extracted the quantities 419 $\delta \tau_{\text{spec}}(t, sat)$ and $\delta \tau_{\text{scatt}}(t, sat)$, which are represented in 420 Figs. 7 and 8 as a function of longitude, respectively. Each dot 421 in these figures represents a 1-s delay sample. The number of 422 samples obtained for each PRN is given by N in Table III. 423 Tables III and IV provide the mean, the standard deviation 424 of the 1-s delay samples, and the estimated standard devia-425 tion for N-second delay samples. We also include in Fig. 9 426 the series of the amplitudes of the waveform. The variations 427 of the PRN3 amplitudes are consistent with the changes in 428 the gain shown in Fig. 5. The standard deviation of the 1-s 429 measured altimetric delays is consistent with those obtained



Fig. 7. Altimetric delays $\delta \tau_{\text{spec}}$ (in meters) as a function of the longitude of the corresponding specular point.



Fig. 8. Scatterometric delays $\delta \tau_{\text{scatt}}$ (in meters) as a function of the longitude of the corresponding specular point.

TABLE III Computed Mean, Standard Deviation of the 1-s Delay Samples, and the Estimated Standard Deviation for N-Second Delay Samples in the Series of Specular Delays $\delta \tau_{\rm spec}$ for the Three Satellites PRN19, PRN22, and PRN3. N is the Total Number of 1-s Samples for Each PRN

PRN	mean (m)	1-sec std. dev. (m)	N-sec std. dev. (m)	N
19	-4.55	2.51	0.06	1751
22	-4.58	2.25	0.07	1086
3	-6.21	2.06	0.06	1145

TABLE IV COMPUTED MEAN, STANDARD DEVIATION OF THE1-S DELAY SAMPLES, AND THE ESTIMATED STANDARD DEVIATION FOR N-Second Delay SAMPLES OF THE SERIES $\delta \tau_{scatt}$ FOR THE THREE SATELLITES PRN19, PRN22, AND PRN3 (N IS GIVEN IN Table III)

PRN	mean (m)	1-sec std. dev. (m)	N-sec std. dev. (m)
19	62.49	3.98	0.10
22	60.35	3.58	0.11
3	58.55	3.86	0.11

in other aircraft experiments [22], [23] and in agreement with 430 the theoretical expectations [7], [24], and the *N*-sample delay 431 standard deviation of the series, on the order of 10 cm, indicates 432 the potential of the technique to measure the sea-surface height, 433 if the type-B uncertainties are reduced below this quantity. 434

As shown in Fig. 8 and Table IV, the 1-s standard deviation of 435 the scatterometric delays $\delta \tau_{\rm scatt}$ is smaller than 4 m. According 436 to the values of the sensitivity of the scatterometric delay with 437 respect to the mss (see the fifth column in Table II), this will 438 represent uncertainties in the mss on the order of 0.002. Such 439 level of uncertainties in the sea-surface roughness will, in turn, 440



Fig. 9. Amplitude of the waveforms (in arbitrary units) as a function of the longitude of the corresponding specular point.

441 affect the specular delay (column six in Table II), producing 442 an rms dispersion of 5 cm in $\delta \tau_{\rm spec}$, which is negligible when 443 compared with the 2–3-m standard deviation of the 1-s delays. 444 As a consequence, we could ignore the contribution of Δmss 445 in (22)

$$\delta \tau_{\rm spec} = -2 \cdot \sin(e) \cdot \Delta h + b_{\rm spec}.$$
 (24)

446 In VLBI or GPS, this equation is used to separate the 447 geometrical part from a bias term. Even in the case where the 448 observation uncertainty is very small, a perfect separation is 449 difficult because the range of elevations, as well as other effects 450 like the atmospheric delay, which have similar signatures in the AQ13 451 delay, is limited. This is the reason why the VLBI and GPS 452 estimates of vertical components of baselines are less accurate 453 than the horizontal components. Examples of this approach in 454 the analysis of GNSS-R observations for altimetry could be 455 found in [8] and [22]. Separation would therefore require the 456 proper calibration of the instrumental biases, as well as the 457 ingestion of additional relevant information (tropospheric/space 458 weather), as it is done in the RA analysis.

459

VIII. DISCUSSION

This paper has first formalized the "retracking" of the GNSS 460 461 reflections by means of the derivative of their waveforms 462 (Sections II and III), which is a method that has the follow-463 ing advantages: 1) being computationally much faster than 464 the standard "retracking" procedure of fitting a model and 465 2) becoming formally independent of any model. The corre-466 sponding algorithms that produce real-time altimetric observ-467 ables to be used in dedicated digital signal processors have then 468 been presented (Sections IV and V). The overall technique is 469 finally implemented and tested with real data gathered in an 470 airborne experiment (work presented in Sections VI and VII) 471 to check that the uncertainties correspond to the theoretically 472 expected uncertainties and are at the same level with those in 473 similar experiments in which standard "retracking" techniques

A014 474 were used.

The geometry of this particular experiment hindered the 475 476 separation of some of the type-B uncertainties that typically can be retrieved with GNSS, as it is likely to happen in a 477 hypothetical orbital receiver. Similar problems are faced by RA 478 missions, which successfully solve them by ingestion of exter- 479 nal corrections. The results have confirmed that the altimetric 480 performance is not reduced by the technique but limited by the 481 bandwidth of the GNSS signal used (2.026 MHz) and the lack 482 of instrumental calibration. The theoretical range performance 483 improves drastically with a much wider bandwidth instrument 484 (4.5-cm instrument noise and speckle estimated in [25] for a 485 space GNSS-R interferometer), for which the described tech- 486 nique could provide a quick "retracking" method, independent 487 on modeling issues. 488

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496

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